

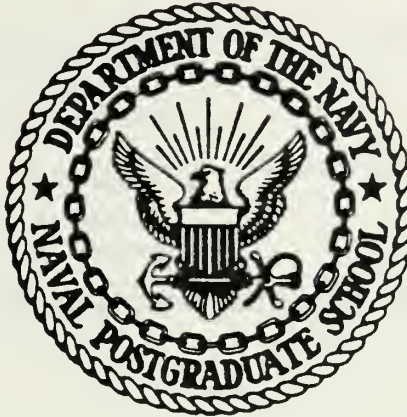
MEASUREMENT OF THE CALIFORNIA COUNTER-  
CURRENT

By Keith Coddington



# NAVAL POSTGRADUATE SCHOOL

## Monterey, California



# THESIS

MEASUREMENT OF THE  
CALIFORNIA COUNTERCURRENT

by

Keith Coddington

June 1979

Thesis Advisors:

J. B. Wickham  
S. P. Tucker

Approved for public release; distribution unlimited.

Thesis  
C53025

T191027

Thesis  
C 53025  
c.1

5901911

SECURITY CLASSIFICATION OF THIS PAGE (When Data Entered)

REPORT DOCUMENTATION PAGE		READ INSTRUCTIONS BEFORE COMPLETING FORM
1. REPORT NUMBER	2. GOVT ACCESSION NO.	3. RECIPIENT'S CATALOG NUMBER
4. TITLE (and Subtitle) Measurement of the California Countercurrent		5. TYPE OF REPORT & PERIOD COVERED Master's Thesis; June 1979
		6. PERFORMING ORG. REPORT NUMBER
7. AUTHOR(s) Keith Coddington		8. CONTRACT OR GRANT NUMBER(s)
9. PERFORMING ORGANIZATION NAME AND ADDRESS Naval Postgraduate School Monterey, California 93940		10. PROGRAM ELEMENT, PROJECT, TASK AREA & WORK UNIT NUMBERS
11. CONTROLLING OFFICE NAME AND ADDRESS Naval Postgraduate School Monterey, California 93940		12. REPORT DATE June 1979
		13. NUMBER OF PAGES 135
14. MONITORING AGENCY NAME & ADDRESS (if different from Controlling Office)		15. SECURITY CLASS. (of this report) Unclassified
		15a. DECLASSIFICATION/DOWNGRADING SCHEDULE
16. DISTRIBUTION STATEMENT (of this Report)  Approved for public release; distribution unlimited.		
17. DISTRIBUTION STATEMENT (of the abstract entered in Block 20, if different from Report)		
18. SUPPLEMENTARY NOTES		
19. KEY WORDS (Continue on reverse side if necessary and identify by block number) California Countercurrent California Undercurrent Davidson Current California Current		
20. ABSTRACT (Continue on reverse side if necessary and identify by block number)  Direct measurements by moored current meters and indirect measurements from geostrophy are compared and discussed for a region over the continental slope off central California during the Davidson Current period.  During that same period vertical temperature and salinity profiles were made at 23 stations on four separate cruises in the study area south of Monterey, California. These arrays of		



## #20 - ABSTRACT - CONTINUED

moored current meters simultaneously recorded the flow of the current at specified levels.

The California Countercurrent was found to be present in the region of study during the entire observation period. Its offshore position and extent, its intensity and its vertical location and extent varied in a way largely consistent with its reported behavior in other locations along the U.S. West Coast.







Measurement of the  
California Countercurrent

by

Keith Coddington  
Lieutenant, United States Coast Guard  
B.S., United States Coast Guard Academy, 1973

Submitted in partial fulfillment of the  
requirements for the degree of

MASTER OF SCIENCE IN OCEANOGRAPHY

from the

NAVAL POSTGRADUATE SCHOOL

June 1979



## ABSTRACT

Direct measurements by moored current meters and indirect measurements from geostrophy are compared and discussed for a region over the continental slope off central California during the Davidson Current period.

During that same period vertical temperature and salinity profiles were made at 23 stations on four separate cruises in the study area south of Monterey, California. These arrays of moored current meters simultaneously recorded the flow of the current at specified levels.

The California Countercurrent was found to be present in the region of study during the entire observation period. Its offshore position and extent, its intensity and its vertical location and extent varied in a way largely consistent with its reported behavior in other locations along the U.S. West Coast.



## TABLE OF CONTENTS

I.	INTRODUCTION -----	13
	A. THE CALIFORNIA CURRENT SYSTEM -----	13
	B. PREVIOUS STUDIES OF THE CALIFORNIA UNDERCURRENT -----	14
	C. STATEMENT OF THE PROBLEM -----	19
II.	AREA OF INVESTIGATION -----	24
III.	SALINITY-TEMPERATURE-DEPTH OBSERVATIONS -----	27
	A. INSTRUMENTATION AND DATA COLLECTION -----	27
	B. RESULTS -----	28
	C. DISCUSSION OF WATERMASS PROPERTIES AND GEOSTROPHY -----	29
IV.	DIRECT CURRENT OBSERVATIONS -----	54
	A. INSTRUMENTATION AND DATA COLLECTION -----	54
	B. DISCUSSION OF DIRECT CURRENT MEASUREMENTS -	58
V.	COMPARISON OF GEOSTROPHY WITH DIRECT CURRENT MEASUREMENTS -----	70
VI.	CONCLUSIONS -----	73
APPENDIX A:	NOYFB PROGRAM -----	76
APPENDIX B:	TEMPERATURE, SIGMA-T, SOUND SPEED, AND SIGMA-T AND SALINITY SUPERIMPOSED VERTICAL SECTIONS -----	99
BIBLIOGRAPHY	-----	131
INITIAL DISTRIBUTION LIST	-----	134



## LIST OF TABLES

TABLES	PAGE
I. Latitude, longitude and depth of stations --	26
II. Comparison of current meter and geostrophic current velocities on 27 November 1978 and 8 January 1979 -----	75





# LIST OF FIGURES

FIGURE		PAGE
1.	Temperature-salinity curves from selected stations (Sverdrup and Johnson, 1941) -----	21
2.	Diagram showing T-S curves defining percentage southern water (Sverdrup and Johnson, 1941) ----	21
3.	Mean dynamic topography of the sea surface reference 500 db during January off California and Baja California (Hickey, 1978) -----	22
4.	Mean dynamic topography of the 200 db surface reference 500 db during January off California and Baja California (Hickey, 1978) --	22
5.	Temperature-salinity relationships for selected stations in California undercurrent (Wooster and Jones, 1970) -----	23
6.	Chart indicating area of investigation and stations -----	25
7.	Salinity (‰) on a vertical section for the Cape San Martin line on 27-28 November 1978 ----	34
8.	Salinity (‰) on a vertical section for the Slate Rock line on 27-28 November 1978 -----	35
9.	Salinity (‰) on a vertical section for the Cape San Martin line on 8-9 January 1979 -----	36
10.	Salinity (‰) on a vertical section for the Slate Rock line on 8-9 January 1979 -----	37
11.	Salinity (‰) on a vertical section for the Cape San Martin line on 22-23 January 1979 -----	38
12.	Salinity (‰) on a vertical section for the Slate Rock line on 22-23 January 1979 -----	39
13.	Salinity (‰) on a vertical section for the Cape San Martin line on 21-22 February 1979 ----	40
14.	Salinity (‰) on a vertical section for the Slate Rock line on 21-22 February 1979 -----	41
15.	Dynamic topography of the 100/500 db surface on 27-28 November 1978 -----	42



16.	Dynamic topography of the 200/500 db surface on 27-28 November 1978 -----	42
17.	Dynamic topography of the 300/500 db surface on 27-28 November 1978 -----	42
18.	Dynamic topography of the 100/500 db surface on 8-9 January 1979 -----	43
19.	Dynamic topography of the 200/500 db surface on 8-9 January 1979 -----	43
20.	Dynamic topography of the 300/500 db surface on 8-9 January 1979 -----	43
21.	Dynamic topography of the 100/500 db surface on 22-23 January 1979 -----	44
22.	Dynamic topography of the 200/500 db surface on 22-23 January 1979 -----	44
23.	Dynamic topography of the 300/500 db surface on 22-23 January 1979 -----	44
24.	Dynamic topography of the 100/500 db surface on 21-22 February 1979 -----	45
25.	Dynamic topography of the 200/500 db surface on 21-22 February 1979 -----	45
26.	Dynamic topography of the 300/500 db surface on 21-22 February 1979 -----	45
27.	Vertical section of the normal component of geostrophic velocity for the Cape San Martin line on 27-28 November 1978 -----	46
28.	Vertical section of the normal component of geostrophic velocity for the Slate Rock line on 27-28 November 1978 -----	47
29.	Vertical section of the normal component of geostrophic velocity for the Cape San Martin line on 8-9 January 1979 -----	48
30.	Vertical section of the normal component of geostrophic velocity for the Slate Rock line on 8-9 January 1979 -----	49



## FIGURE

## PAGE

31.	Vertical section of the normal component of geostrophic velocity for the Cape San Martin line on 22-23 January 1979 -----	50
32.	Vertical section of the normal component of geostrophic velocity for the Slate Rock line on 22-23 January 1979 -----	51
33.	Vertical section of the normal component of geostrophic velocity for the Cape San Martin line on 21-22 February 1979 -----	52
34.	Vertical section of the normal component of geostrophic velocity for the Slate Rock line on 21-22 February 1979 -----	54
35.	Array configuration -----	57
36.	Progressive vector diagram for the current meter at station 2 at 220 meters depth from 25 July 1978 to 28 August 1978 -----	63
37.	Progressive vector diagram for the current meter at station 2 at 190 meters depth from 20 September 1978 to 27 November 1978 -----	64
38.	Progressive vector diagram for the current meter at station 2 at 100 meters depth from 27 November 1978 to 22 January 1979 -----	65
39.	Progressive vector diagram for the current meter at station 2 at 175 meters depth from 27 November 1978 to 22 January 1979 -----	66
40.	Progressive vector diagram for the current meter at station 2 at 300 meters depth from 27 November 1978 to 22 January 1979 -----	67
41.	Progressive vector diagram for the current meter at station 5 at 140 meters depth from 27 November 1978 to 22 January 1979 -----	68
42.	Progressive vector diagram for the current meter at station 5 at 215 meters depth from 27 November 1978 to 22 January 1979 -----	69
43.	Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Cape San Martin line on 27-28 November 1978 -	99
44.	Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Slate Rock line on 27-28 November 1978 -----	100





FIGURE		PAGE
45.	Temperature (°C) on a vertical section for the Cape San Martin line on 8-9 January 1979 -----	101
46.	Temperature (°C) on a vertical section for the Slate Rock line on 8-9 January 1979 -----	102
47.	Temperature (°C) on a vertical section for the Cape San Martin line on 22-23 January 1979 -----	103
48.	Temperature (°C) on a vertical section for the Slate Rock line on 22-23 January 1979 -----	104
49.	Temperature (°C) on a vertical section for the Cape San Martin line on 21-22 February 1979 -----	105
50.	Temperature (°C) on a vertical section for the Slate Rock line on 21-22 February 1979 -----	106
51.	Sigma-t on a vertical section for the Cape San Martin line on 27-28 November 1978 -----	107
52.	Sigma-t on a vertical section for the Slate Rock line on 27-28 November 1978 -----	108
53.	Sigma-t on a vertical section for the Cape San Martin line on 8-9 January 1979 -----	109
54.	Sigma-t on a vertical section for the Slate Rock line on 8-9 January 1979 -----	110
55.	Sigma-t on a vertical section for the Cape San Martin line on 22-23 January 1979 -----	111
56.	Sigma-t on a vertical section for the Slate Rock line on 22-23 January 1979 -----	112
57.	Sigma-t on a vertical section for the Cape San Martin line on 21-22 February 1979 -----	113
58.	Sigma-t on a vertical section for the Slate Rock line on 21-22 February 1979 -----	114
59.	Sound speed (m/sec) on a vertical section for the Cape San Martin line on 27-28 November 1978 -	115
60.	Sound speed (m/sec) on a vertical section for the Slate Rock line on 27-28 November 1978 -----	116
61.	Sound speed (m/sec) on a vertical section for the Cape San Martin line on 8-9 January 1979 ----	117



62.	Sound speed (m/sec) on a vertical section for the Slate Rock line on 8-9 January 1979 -----	118
63.	Sound speed (m/sec) on a vertical section for the Cape San Martin line on 22-23 January 1979 --	119
64.	Sound speed (m/sec) on a vertical section for the Slate Rock line on 22-23 January 1979 -----	120
65.	Sound speed (m/sec) on a vertical section for the Cape San Martin line on 21-22 February 1979 -	121
66.	Sound speed (m/sec) on a vertical section for the Slate Rock line on 21-22 February 1979 -----	122
67.	Sigma-t and salinity (‰) superimposed on a vertical section for the Cape San Martin line on 27-28 November 1978 -----	123
68.	Sigma-t and salinity (‰) superimposed on a vertical section for the Slate Rock line on 27-28 November 1978 -----	124
69.	Sigma-t and salinity (‰) superimposed on a vertical section for the Cape San Martin line on 8-9 January 1979 -----	125
70.	Sigma-t and salinity (‰) superimposed on a vertical section for the Slate Rock line on 8-9 January 1979 -----	126
71.	Sigma-t and salinity (‰) superimposed on a vertical section for the Cape San Martin line on 22-23 January 1979 -----	127
72.	Sigma-t and salinity (‰) superimposed on a vertical section for the Slate Rock line on 22-23 January 1979 -----	128
73.	Sigma-t and salinity (‰) superimposed on a vertical section for the Cape San Martin line on 21-22 February 1979 -----	129
74.	Sigma-t and salinity (‰) superimposed on a vertical section for the Slate Rock line on 21-22 February 1979 -----	130



## ACKNOWLEDGMENT

The author would like to express his sincere appreciation to his thesis advisors, Professor S.P. Tucker and Professor J.B. Wickham. Their knowledge and expertise in this project was only surpassed by their willingness to freely devote their time.

Special appreciation is extended to the Captain, W.W. Reynolds, and crew of the R/V ACANIA. Their highly skilled seamanship provided for smooth array installation and retrieval. Appreciation is also extended to Mr. Tim Stanton for his assistance in translating the current meter tapes.

Finally, a devoted thanks to my wife, Michelle, for her patience and understanding during this time.



## I. INTRODUCTION

### A. THE CALIFORNIA CURRENT SYSTEM

In the past decade eastern boundary currents and coastal upwelling have come under considerable scrutiny. This is due in part to their influence on the fishing industry and various other economic enterprises. Examples of eastern boundary current systems are the California, the Peru, the Benguela and the Canary Current Systems.

One eastern boundary current system, the California Current system is made up of an equatorward surface flow which extends along the entire west coast of the United States and Baja California and a counterflow, poleward, sometimes beneath this, at others on the surface shoreward of it. The equatorward surface flow is fed by the North Pacific Current, i.e. the northern limb of the North Pacific Gyre, and is known as the California Current. The poleward flow may be submerged, when it is then termed the California Undercurrent, or, at certain times of the year it may appear at the surface, when it is called the Davidson Current.

Reid, Roden and Wyllie (1958) apply the term California Current to all southward flow in the North Pacific Gyre. It is common to define the boundary of the California Current at a distance 1000 kilometers from shore (Hickey, 1978). High velocities are generally not encountered in this cold water mass.





The California Undercurrent is a poleward flow of water, the temperature and salinity of which is slightly higher than that of the surrounding water and which is usually found shoreward of the southward flowing California Current. It is much narrower and has a maximum of northward flow at intermediate depth. It is uncertain whether the Davidson Current is superimposed on the California Undercurrent, suppressing the core to great depths, or whether the Davidson Current is actually the expression of the undercurrent at the surface (Hickey, 1978).

The California Undercurrent is present year round and, with the onset of north-northwest winds and upwelling along California, it is found predominately below 200 meters. It is characterized by relatively high temperature, salinity and phosphate and low dissolved oxygen concentrations because of its southern origin. The undercurrent has been intensely studied off Oregon, Washington, and Southern California, but direct current measurements are lacking for the Central California region.

#### B. PREVIOUS STUDIES OF THE CALIFORNIA UNDERCURRENT

The California Undercurrent was first discussed by Sverdrup and Fleming (1941). During their cruises in 1937, they defined "northern water" on a T-S curve which showed an increase in salinity with decreasing temperature (Figure 1, curve C131). The T-S curve for "southern water" showed salinity relatively constant as temperature decreased (Figure 1,



curves 5.3 and B III, 31). They constructed a chart, defining percentage of southern water for given T-S pairs (Figure 2). Using these parameters they traced southern water as far north as Cape Mendocino. They also found that the southern water was close to the coast and was concentrated in the northward flowing current. They also showed the existence of the northward flowing undercurrent by means of dynamic heights.

Reid, Roden, and Wyllie (1958) expanded on Sverdrup and Fleming (1941) and Sverdrup, et al. (1942). They concluded the evidence for the undercurrent was of two sorts: (1) The warm or more saline subsurface water of low oxygen content suggested southern origin; and (2) geostrophic flow at the 200 decibar surface with respect to both the 500 and 1000 decibar surfaces indicated a northward flowing current, 30-60 miles in width near the coast north of 30°N and somewhat wider to the south.

Direct measurements of the undercurrent was made by Reid (1962, 1963) and Reid and Schwartzlose (1962) using drogues. Their results indicated the existence of a northward flowing current at 200 meters depth off Monterey, California, and Baja California. During the winter they found that a northward flow existed at the surface.

As a part of a California Cooperative Oceanic Fisheries Investigation (CALCOFI) study Wyllie (1966) showed the existence of a northward flowing undercurrent on the basis of dynamic topography. Wyllie's charts of mean monthly dynamic



topography on the 200 decibar surface relative to 500 decibars provide the best description of the flow of the undercurrent in January south of Cape Mendocine (Figures 3 and 4).

From Point Conception northward stronger subsurface flow was observed in winter than in summer. The weakest flow occurred from March to May. Pavlova (1966) found that during the spring, when the undercurrent appears to be absent at the usual depth of about 200 meters, it may be present at depths exceeding 500 meters. Pavlova also concluded that the undercurrent actually reaches the surface during late fall and winter, when it is known as the Davidson Current. Wyllie's data supported these conclusions.

Wooster and Jones (1970) found that a characteristic of the undercurrent was a relatively high salinity bulge centered at  $\sigma_t$  equal 26.54 (150 cl/t) on the T-S diagram (Figure 5). They also gave some evidence for an inter-annual variation in the northward extent of a given isohaline. They pointed out that a coastal deepening of isotherms and isopycnals and rising of isohalines are characteristics of the poleward undercurrent.

In the last ten years the study of the undercurrent has been concentrated north of Cape Mendocino and to some extent in the vicinity of Monterey, California. Mooers, Collins and Smith (1976) in their study of upwelling off the Oregon coast found a northward flow along the continental slope between 300 and 1000 meters. They suggested that it may exist





at greater depths and may extend from the continental slope to perhaps 500 kilometers from shore. Their primary observations were conducted during the same year and season as those of Wooster and Jones (1970). Mooers, et al, (1976) found that during July 1975, the near surface flow was predominately southward, and the near bottom flow alternated between northward and southward. In August and September 1965 and 1966 the near bottom flow was predominately northward.

Huyer and Smith (1976) and Halpern, Smith and Reed (1978) used direct measurements of current on the slope and shelf to describe the seasonal developments of the undercurrent off the coast of Oregon. Huyer and Smith's (1976) data suggest that the northward flow is present at depths greater than 400 meters in the spring but increases in speed and vertical extent as the season progresses. That the undercurrent was found at the shelf edge by summer is consistent with Pavlova's (1966) findings for northern California. Halpern, Smith, and Reed's (1978) current meter results support those of Huyer and Smith (1976). The data of both these studies suggest that the shelf and slope undercurrents were portions of the same flow.

Eddies are frequently observed off the coast of Vancouver Island, B.C. Mysak (1977) in conjunction with his study of the undercurrent suggests that the eddies are produced by baroclinic instability of the California Undercurrent. For the undercurrent he found a northward flow along the



continental slope but a southward flow farther offshore. Thus, off Vancouver Island the northward flowing California Undercurrent is essentially confined to the continental slope. The main core of the current at that latitude occurs around 300 meters.

Off Monterey, California, the undercurrent has been studied by Molnar (1972), Hughes (1975), Greer (1975), and Wickham (1975). Wickham (1975) used drogues and a continuous measuring salinity-temperature-depth profiler (STD). For August three main results were found in conjunction with the undercurrent: (1) At both 50 meters and 200 meters geostrophy and drogues both indicate that there is a narrow band of poleward flowing water near the shelf edge; (2) both drogues and geostrophy also indicate that there is a complex flow farther west which seems to split the poleward flow into two branches; and (3) there is a broader poleward flow still farther west which is centered at 40-50 kilometers from the shelf edge.

An analytical model by McCreary (1977) indicated that, due to local wind forcing, the pycnocline tilts alongshore to balance the meridional component of the wind and results in an alongshore flow. This disturbance is not confined to the coast but propagates offshore and northward as a Kelvin-Rossby wave carrying along with it both the pycnocline deformation and the alongshore flow.

Hickey (1978) has examined most of the data to date. She found that the northward subsurface flow is generally



found off the west coast of North America over the continental slope. The flow on the 200 decibar surface is most continuous alongshore and strongest (south of Point Conception only) in summer and early fall. It is weakest and least continuous in the spring. North of Point Conception, the flow on the 200 decibar surface is stronger during winter than during summer and fall. She found that the depth of the high-speed core varied seasonally and that the flow appeared to have a jet-like structure, both vertically and horizontally and appeared to extend to the bottom over the slope. This is in agreement with McCreary (1977) who called this jet-like flow, quasi-geostrophic. In support of Wooster and Jones (1970), Hickey (1978) found that the salinity and temperature at the core of the undercurrent generally decreased from about 34.6‰ and 9.5°C off Baja California to about 33.9‰ and 7°C off Vancouver Island.

The flow from the surface to a depth of about 500 meters is confined to the continental slope, but the overall width of the region of northward flow has not been firmly established. The relationship between the undercurrent jet that occurs over the upper slope and the slower broader flow that occurs deeper in the water column farther offshore is uncertain.

#### C. STATEMENT OF THE PROBLEM

The presence of southern water can be inferred from isohalines, isotherms, and isopleths of sound speed; and the geostrophic current can be inferred from isopycnals. The



current can also be found through direct measurements by means of current meters. Wickham (1975) noted that a comparison of geostrophic observations with direct measurements of current would test the utility of geostrophy to describe flow in areas of complexity, such as off Monterey, California.

The initial objective of this study was to collect data by both means. This involved setting up stations where salinity, temperature and depth measurements could be taken and moored current meter arrays could be maintained. The thesis addresses the problems associated with the assembly and maintenance of the moored current arrays, the collection of salinity, temperature and depth data, and the analysis and interpretation of the direct and indirect current observations.







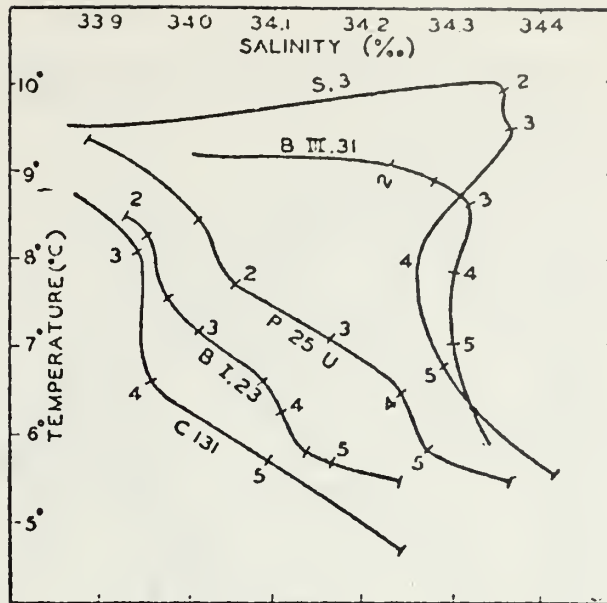


Figure 1. Temperature-salinity curves selected stations (Sverdrup and Johnson, 1941).

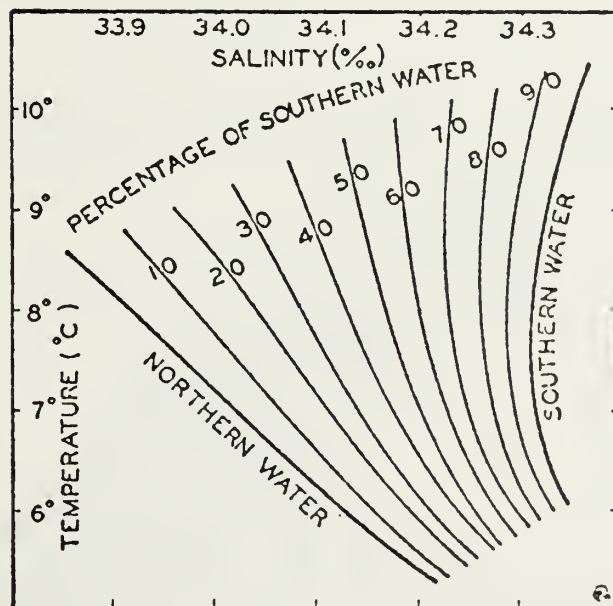


Figure 2. Diagram showing T-S curves defining percentage southern water (Sverdrup and Johnson, 1941).



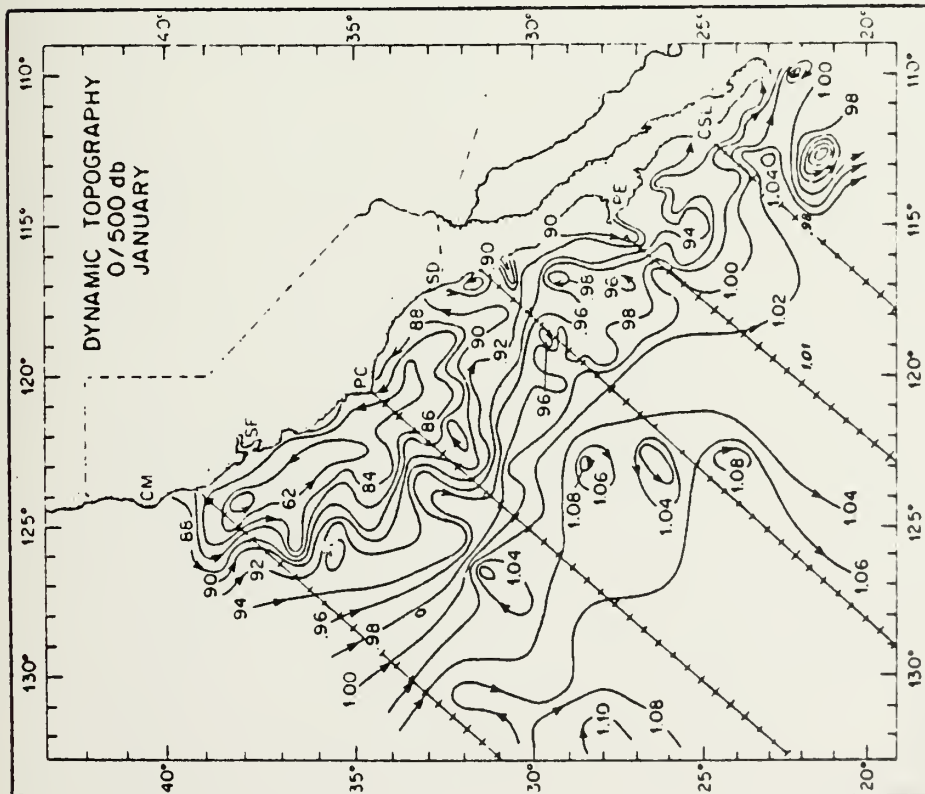


Figure 3. Mean dynamic topography of the sea surface off California and Baja California relative to 500 db during January, contoured from data given by Wyllie (1966). Contour interval is 0.02 dynamic meters (Hickey, 1978).

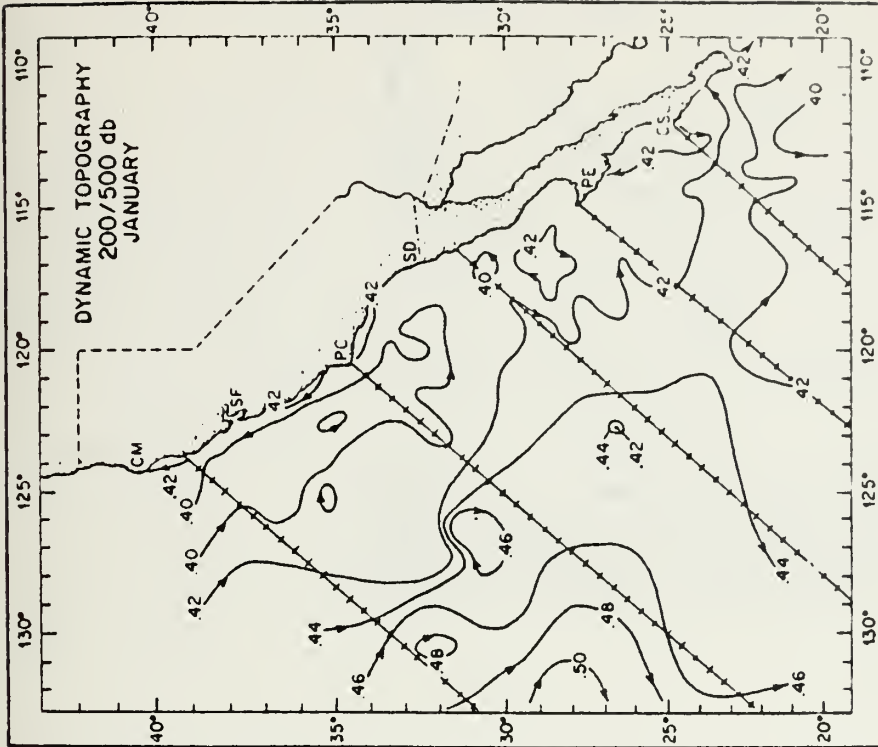


Figure 4. Mean dynamic topography of the 200 db surface relative to 500 db off California and Baja California during January, contoured from data given by Wyllie (1966). Contour interval is 0.02 dynamic meters (Hickey, 1978).



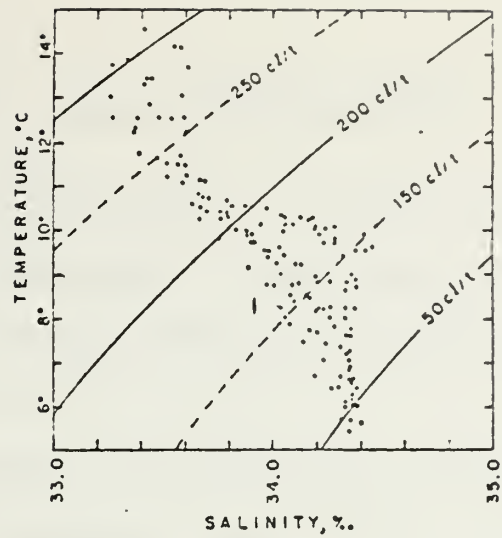


Figure 5. Temperature-salinity relationships for selected stations in California under-current (Wooster and Jones, 1970).



## II. AREA OF INVESTIGATION

The area of investigation is south of Monterey, California, as shown in Figure 6. Two lines of stations were established. The station locations and water depths are listed in Table I.

One of the reasons for positioning the stations on this part of the California coast is the relative simplicity of the bathymetric features. The depth contours run approximately parallel to the coast, and the shelf break is close to the coast. Another, but crucial, reason for using this part of the California coast is that it is less heavily fished than the areas immediately to the north and to the south. The current meter arrays are entirely subsurface with no surface markers. The presence of fishing activity increases the possibility of array damage or loss which we have tried to minimize through our selection of the study area.





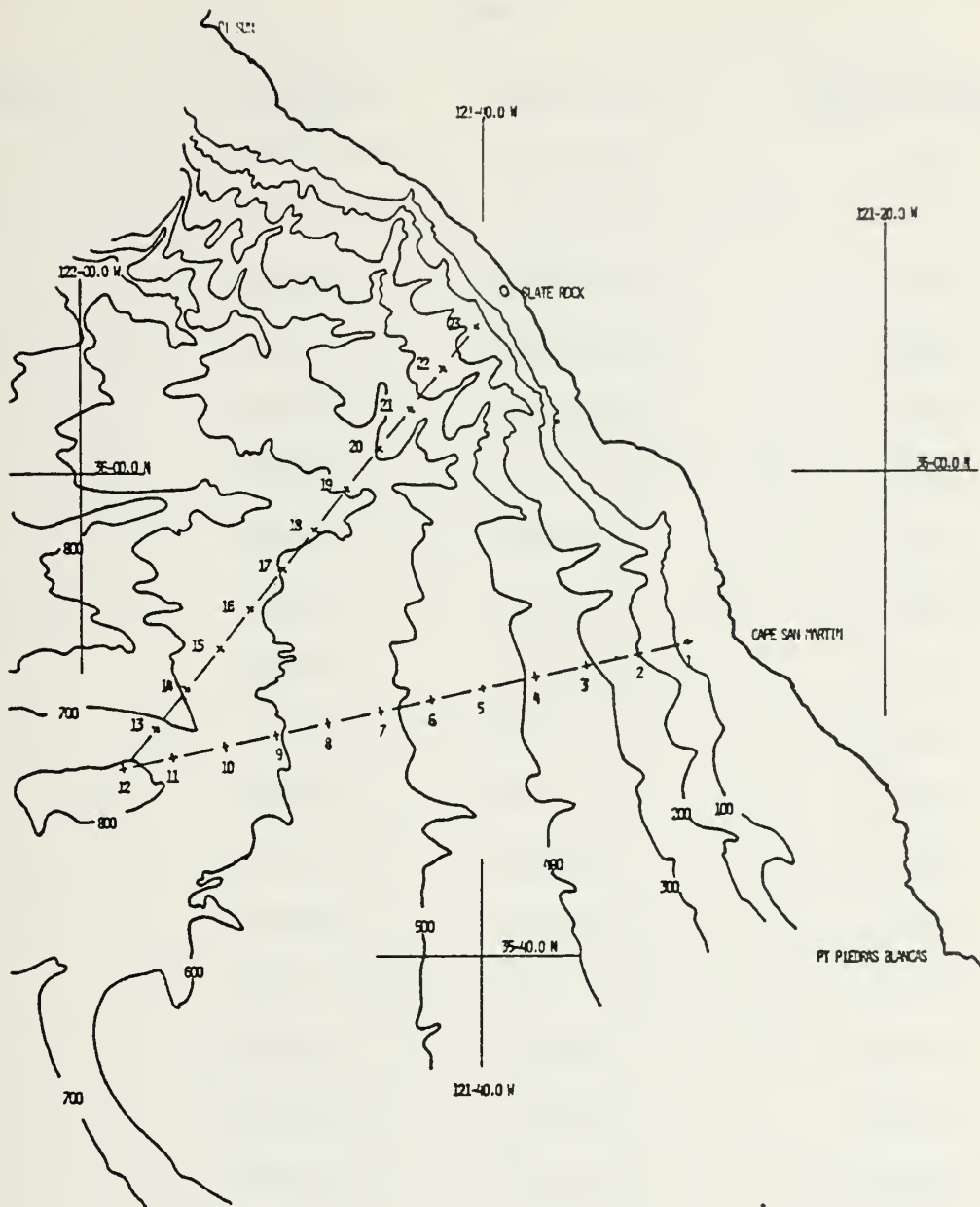


Figure 6. Chart indicating area of investigation and stations. Depth contours in fathoms.



TABLE I

<u>STATION</u>	<u>LATITUDE</u>	<u>LONGITUDE</u>	<u>DEPTH (METERS)</u>
1	35-53.0	121-29.8	100
2	35-52.5	121-32.3	357
3	35-52.1	121-34.7	520
4	35-51.6	121-37.4	668
5	35-51.1	121-39.9	759
6	35-50.7	121-42.4	833
7	35-50.2	121-45.0	915
8	35-49.7	121-47.6	988
9	35-49.2	121-50.1	1061
10	35-48.7	121-52.7	1150
11	35-48.3	121-55.4	1182
12	35-47.8	121-57.7	1044
13	35-49.5	121-56.1	1274
14	35-51.1	121-54.6	1183
15	35-52.7	121-53.0	1146
16	35-54.3	121-51.3	1089
17	35-55.9	121-49.9	1080
18	35-57.6	121-48.3	1098
19	35-59.3	121-47.8	997
20	36-00.9	121-45.2	842
21	36-02.5	121-43.5	732
22	36-04.3	121-42.0	560
23	36-05.9	121-40.2	350



### III. SALINITY-TEMPERATURE-DEPTH OBSERVATIONS

#### A. INSTRUMENTATION AND DATA COLLECTION

Watermass data was collected on four separate cruises of the R/V ACANIA, the oceanographic vessel of the Naval Postgraduate School. The cruises were on 27-28 November 1978, 8-9 January 1979, 22-23 January 1979, and 21-22 February 1979. During each cruise all 23 stations were occupied. Sampling of watermass properties was done where possible to 500 meters, the reference level used by CALCOFI (Wyllie, 1966).

For each cruise the primary instrument, a Bisset Berman Model 9006 STD, was used for delineating the vertical distributions of temperature and salinity. Nansen bottles and reversing thermometers provided independent measurements to check the calibration of the STD. Expendible bathythermograph (XBT) drops and surface temperature observations were made also at each station.

The analog recorder usually used with the Model 9006 STD was replaced on these cruises. The three separated signal frequencies were sent through a Hewlett-Packard Model 57307A VHF Switch with a 20 milisecond settling time to a Hewlett-Packard Model 5328A Universal Counter. The resulting binary coded decimal output was then read into the random access memory of a Hewlett-Packard 9831A desktop computer. After each complete profile the data was transferred to magnetic



tape. Profiles were recorded for both the descent and ascent of the STD.

Spikes in the salinity trace are known to be caused by a poorly matched time constant between the conductivity sensor and the platinum resistance thermometer used in the STD to correct the conductivity measurements to salinity. Salinity spikes were eliminated by comparison of descent and ascent profiles and XBT profiles. This was done by hand during examination of the profiles.

## B. RESULTS

The results for each cruise are shown as vertical sections for both the Cape San Martin line and the Slate Rock line. Vertical sections are drawn for: salinity, temperature, sigma-t, sound speed, geostrophic currents, and sigma-t and salinity together. The dynamic topography of the 100 db, 200 db, and 300 db surfaces relative to 500 decibars is also given for the four separate cruises.

Salinity and temperature sections were contoured using data stored on the magnetic tapes from the HP 9831A computer after the salinity spikes had been removed. Sigma-t, geostrophic shear, and dynamic heights were calculated using the library computer program HYDRO, available at the Naval Postgraduate School for the IBM 360 computer. Sound speed was calculated using Wilson's equations (Wilson, 1960), available in the same program.

From the calculation of geostrophic shear mean values for four station intervals were found. This was done to





reduce the non-geostrophic contributions to the calculations which are inversely proportional to the distance over which geostrophic shear is averaged. Thus, resolution is diminished in order to give a more accurate picture of the larger scale circulation features.

Dynamic heights were smoothed in a similar manner. Values were averaged over three station heights. For station 12 the mean value was found using stations 11, 12, and 13. For the coastal stations 1 and 23 only two stations were used in the averaging, i.e., stations 1 and 2 and stations 22 and 23. For the coastal stations with less than 500 meters of water the stations were treated as if the depth were 500 meters and dynamic topographies were extrapolated from the first station seaward with such depth.

There are no results for stations 21, 22, and 23 during the cruise of 21-22 February 1979, as these stations were lost due to failure of the data recording equipment. On the same cruise there were four stations at five nautical mile intervals added to the west of station 12. This allowed computation of geostrophic shear out to station 12.

#### C. DISCUSSION OF WATERMASS PROPERTIES AND GEOSTROPHY

The presence of souther water is indicated by the distributions in two cross-sections of salinity, temperature, sigma-t, and sound speed on the series of four cruises. The current structure for the same series is deduced from geostrophic current sections and from dynamic topography at



three different levels, both currents and topographies being referred to 500 decibars. Direct current measurements are discussed later in Section IV.

Water with relatively high southern watermass properties is present below the pycnocline on the first cruise on 27-28 November 1978. This is particularly evident on the salinity sections, Figures 7 and 8. Both show a bulge of high salinity water below 200 meters, on Figure 7 for example between stations 3 and 10. The associated temperature and sound speed sections (Appendix B) also indicate the presence of southern water in this region. On this cruise the southern water characteristics appear below 200 meters and from about 4 kilometers to 38 kilometers offshore over the continental slope.

Geostrophy, Figures 27 and 28, indicates northward flow in the upper layers with a surface maximum of 25 to 30 cm/sec. The current appears to have two branches with weaker southward flow between them. Wickham (1975) found similar indications of branched flow farther offshore for his August data in the latitude of Monterey, and these were confirmed by drogue drifts. This branched northward flow is further shown at each of the three levels of contoured dynamic topography (100, 200, and 300 decibars, Figures 15, 16, and 17). Some southward flow appears at all three levels from stations 15 to 18. This is a small-scale feature and might not be real since small scales are not well resolved by geostrophy.



Comparison will be made in Section V between geostrophic and direct current measurements.

The cruise of 8-9 January 1979 also showed southern water; but, as Figures 9 and 10 show, the bulge of high salinity occurs farther west. The associated temperature and sound speed sections (Appendix B) also show this westward displacement of the southern water. Below 200 meters this southern water is found 15 kilometers farther offshore than on the last cruise. Although Pavlova (1966) and Hickey (1978) indicated the countercurrent moves offshore in the spring, our observations were made during the winter season. McCreary's (1977) view of the current's variations as manifestations of baroclinic Kelvin-like waves is consistent with this offshore movement.

Geostrophy in the cross-sections for 8-9 January 1979 (Figures 29 and 30) also shows the northward flow farther offshore. There is still a maximum at the surface, but with an increase in velocity to 70 cm/sec normal to the Cape San Martin line and 35 cm/sec normal to the Slate Rock line. There is an indication that this may be the shoreward branch of the northward flow found during the first cruise, as southward flow appears on the offshore edge of both sections. Dynamic topography (Figures 18, 19, and 20) shows this same pattern. At the 100 db level (Figure 18) the flow is intense and northward between stations 3 and 9. To the west of station 9 the flow intensity drops off sharply and southward flow appears between stations 13 and 15. To the east of





station 3 the dynamic topography is generally flat with northward flow indicated. At the 200 db level (Figure 19) the flow is weaker but is still northward between stations 3 and 9. Southward flow is now more evident between stations 13 and 15. At the 300 db level (Figure 20) the flow is southward with only a trace of northward flow between stations 15 and 17.

The cruise on 22-23 January 1979 shows reductions in southern water characteristics. All indicators, i.e., isopleths of salinity, temperature, sigma-t, and sound speed, are nearly parallel with only small horizontal gradients. Note that the 34.20 ‰ isohaline which in November lay in places higher in the water column than 200 meters is now at a depth of 300 meters, except within a few kilometers of the slope. This may indicate that southern water has moved seaward or deeper beyond the range of observations.

Geostrophy shows slight westward propagation of the northward flow (Figures 31 and 32, also, 21, 22, and 23). The flow appears slower, 20 cm/sec, and more diffuse. At the 200 db level (Figure 22) southward flow now exists shoreward of station 4. At the 300 db level (Figure 23) the flow has become more diffuse and the southern flow is now shoreward of stations 5 and 20.

The observations for the cruise of 21-22 February 1979 indicate an increase in salinity below 200 meters (Figures 13 and 14), the 34.20 ‰ isohaline having risen to a depth of 250 meters over most of both sections. The dynamic topography





(Figures 24, 25, and 26) now indicates a northward flow at stations 2 and 3 at the 200 db and 300 db levels. The geostrophic sections do not show this since the station averaging interval used in their construction does not permit calculations shoreward of station 4. This flow at the eastern stations and below 200 meters may be the start of the undercurrent.

An immediate observation must be made: The regions with indications of southern water and the regions where northward flow is indicated by geostrophy do not exactly coincide. For all the cruises considered, geostrophy shows a northward surface flow, in some instances with flow as great as 70 cm/sec, even though the watermass characteristics in some regions are not southern. This is not too surprising, since near the boundaries between watermasses, eddies and entrainment of anomalous water is common. The observed variations in salinity and velocity may also have alternative explanations. Passing eddies or meanders in the countercurrent might give results similar to those just discussed.

In the following section the currents inferred from geostrophy are compared to those measured directly by moored current meter arrays along the Cape San Martin line.



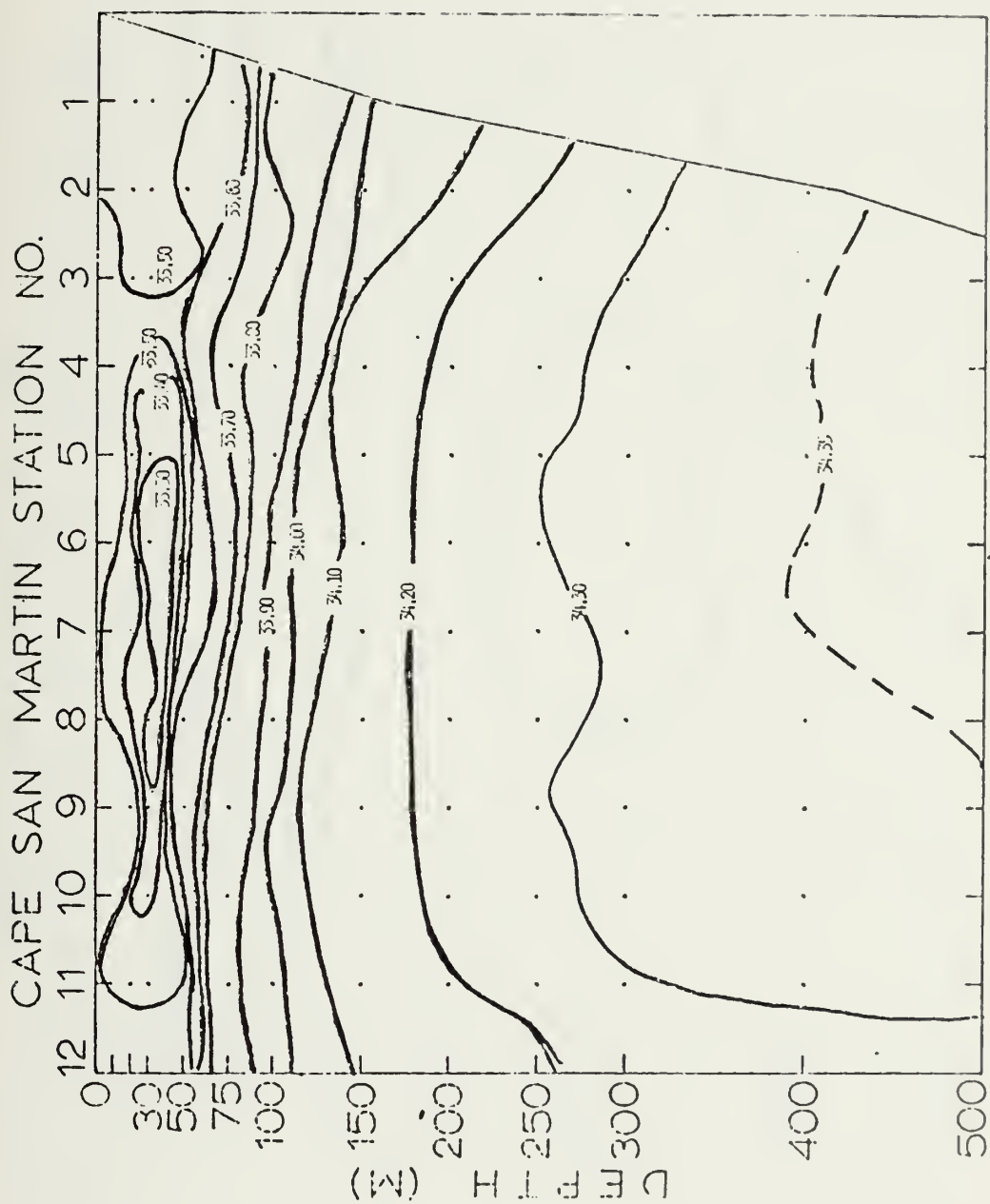


Figure 7. Salinity (‰) on a vertical section for the Cape San Martin line on 27-28 November 1978.



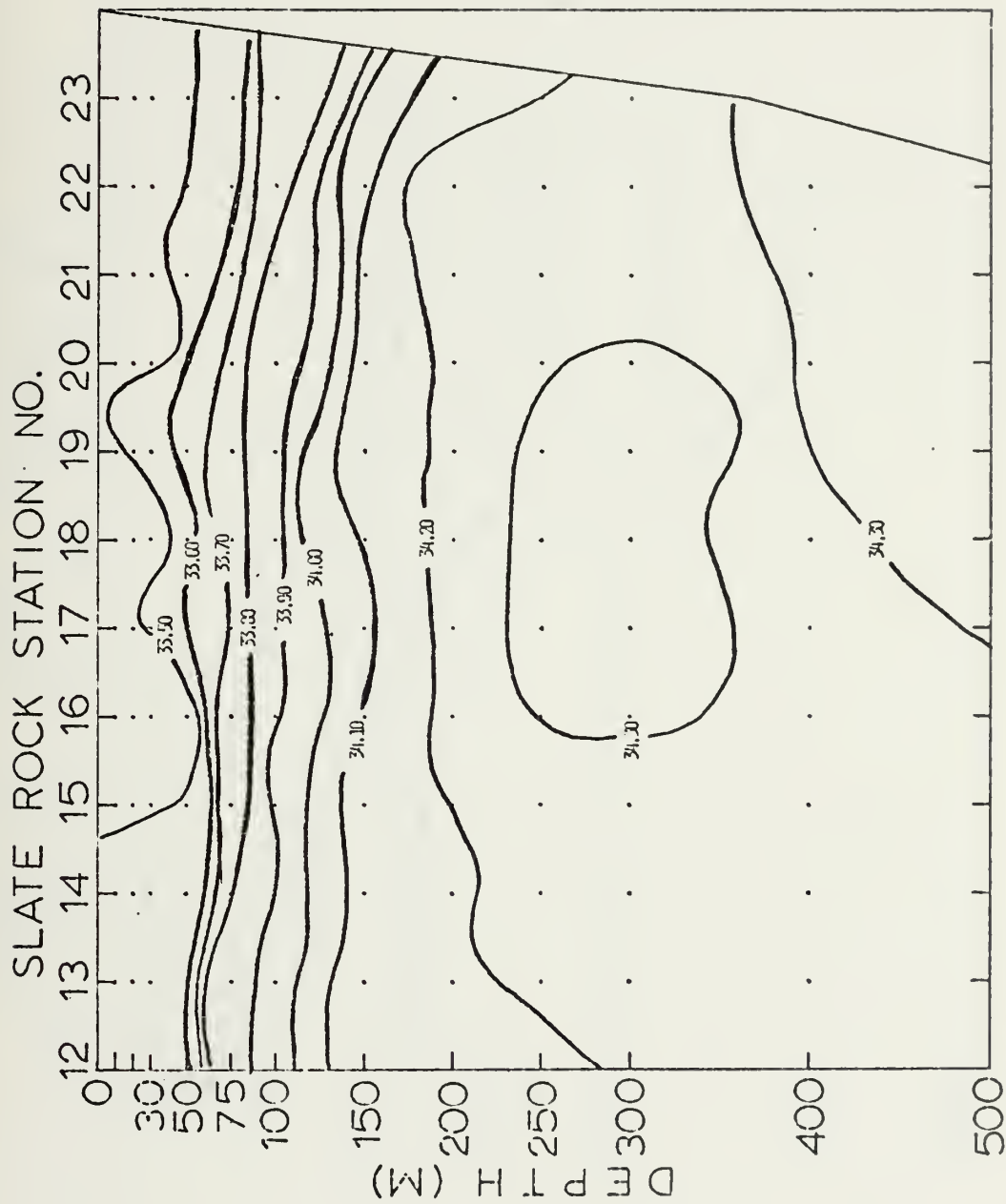


Figure 8. Salinity (‰) on a vertical section for the Slate Rock line on 27-28 November 1978.



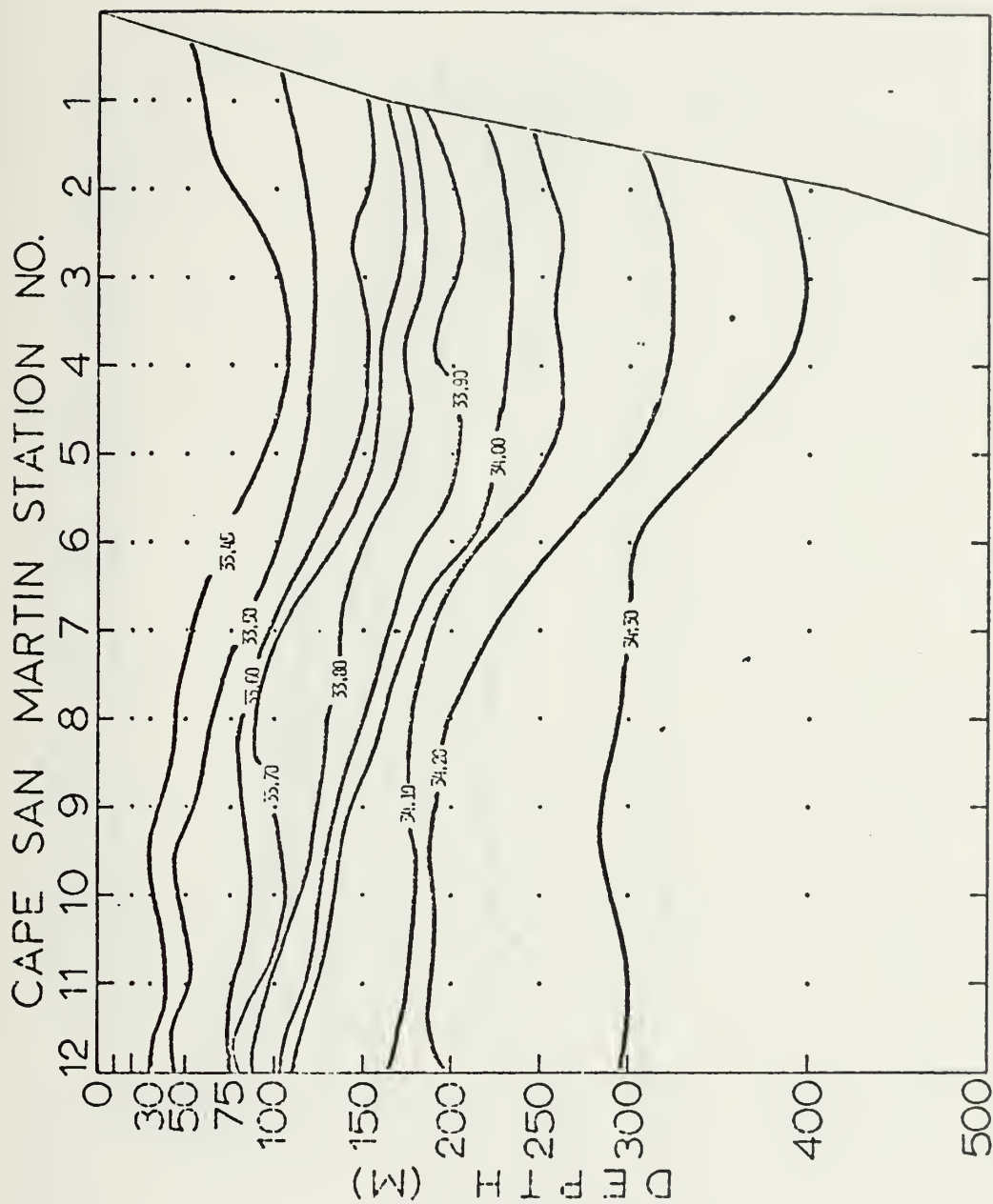


Figure 9. Salinity ( $\text{‰}$ ) on a vertical section for the Cape San Martin line on 8-9 January 1979.





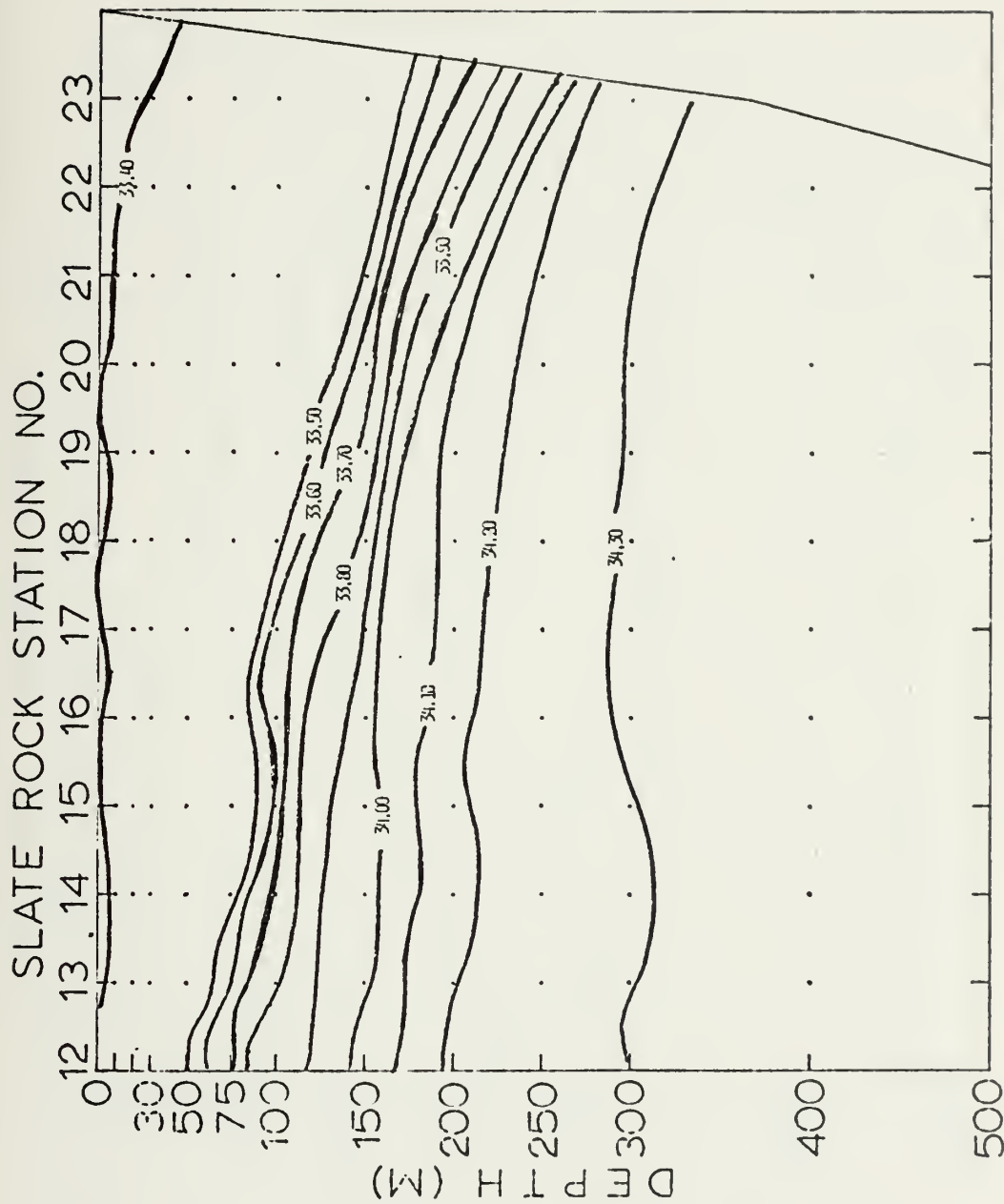


Figure 10. Salinity (‰) on a vertical section for the Slate Rock line on 8-9 January 1979.



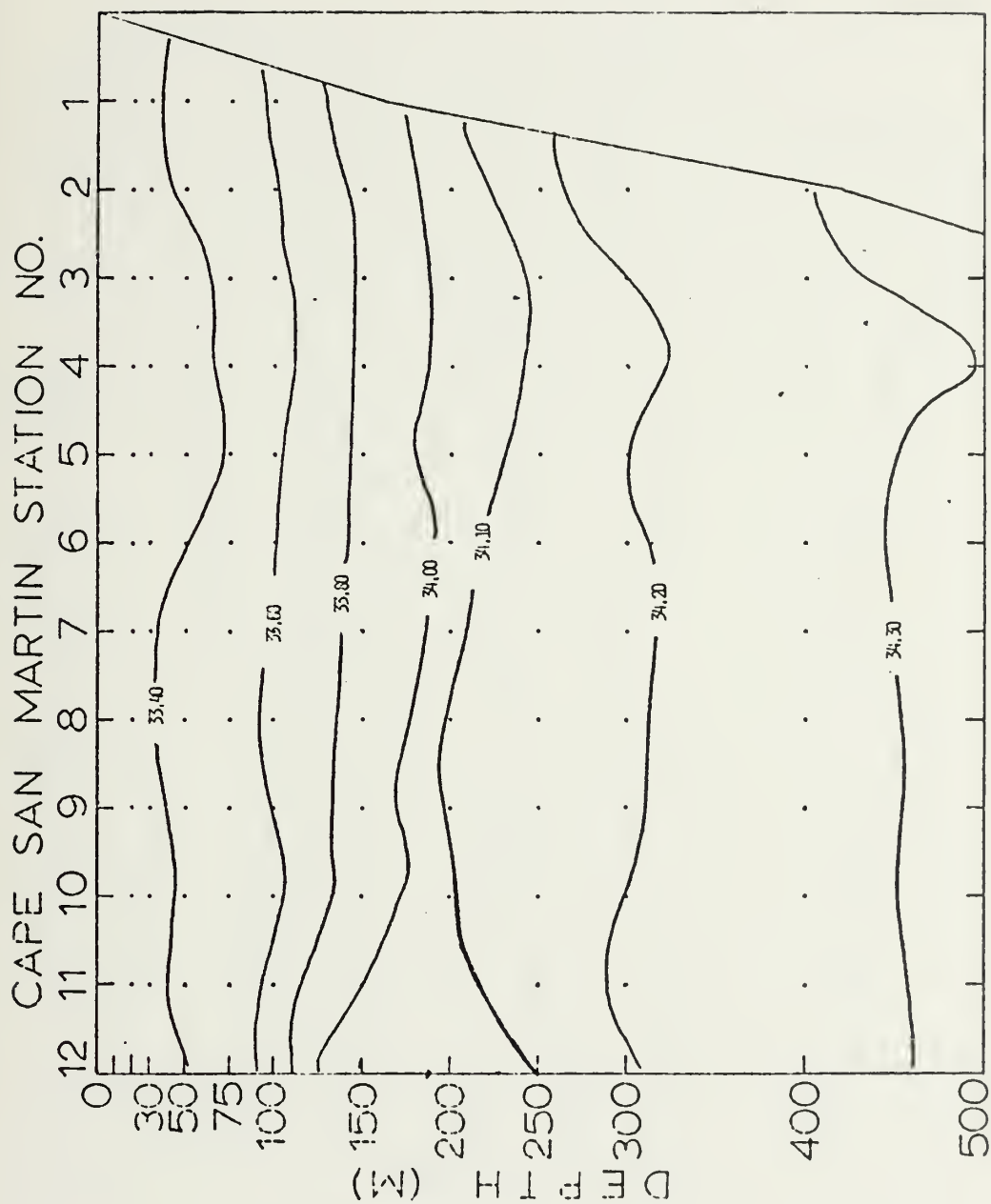


Figure 11. Salinity (‰) on a vertical section for the Cape San Martin line on 22-23 January 1979.



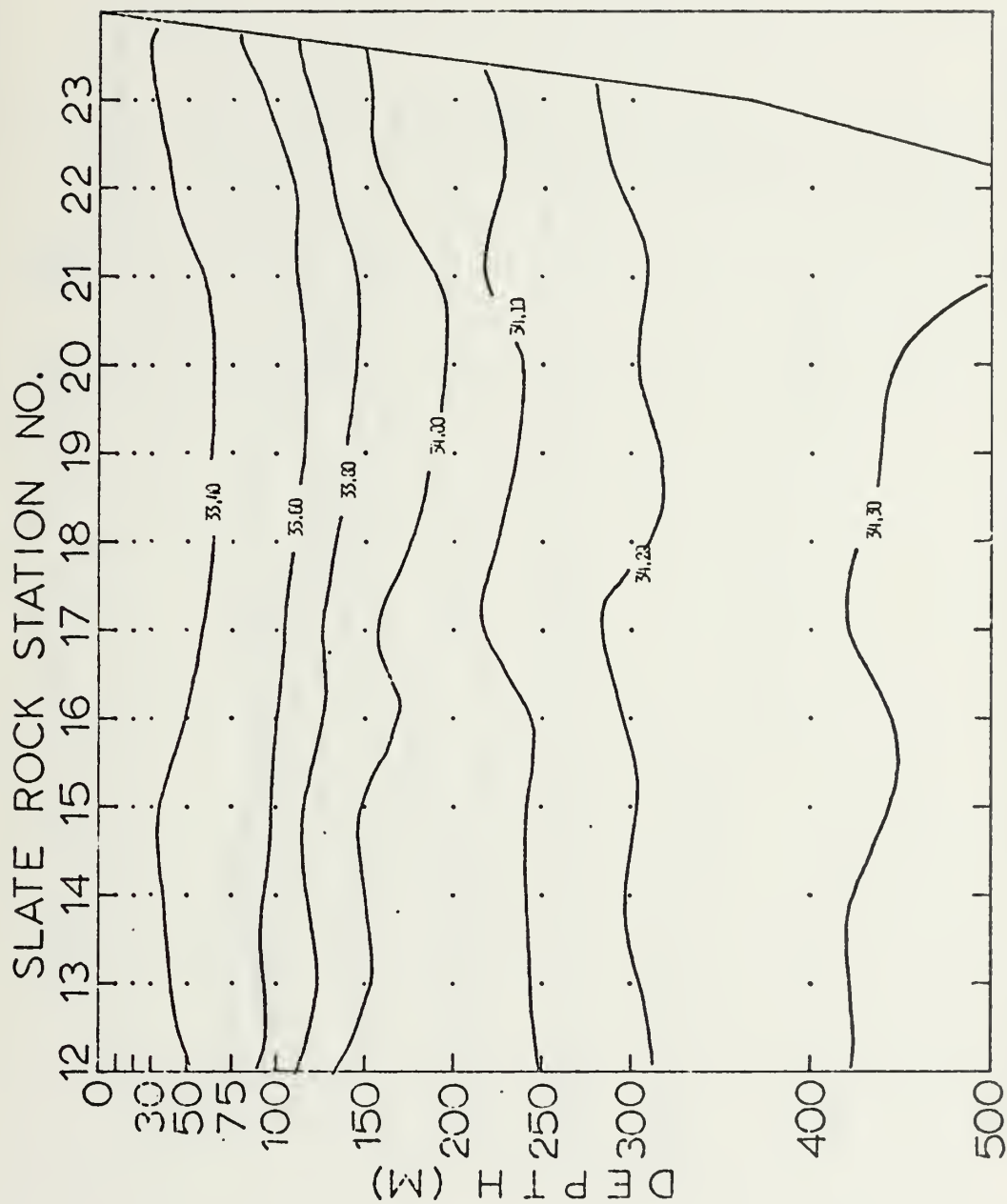


Figure 12. Salinity ( $\text{‰}$ ) on a vertical section for the Slate Rock line on 22-23 January 1979.



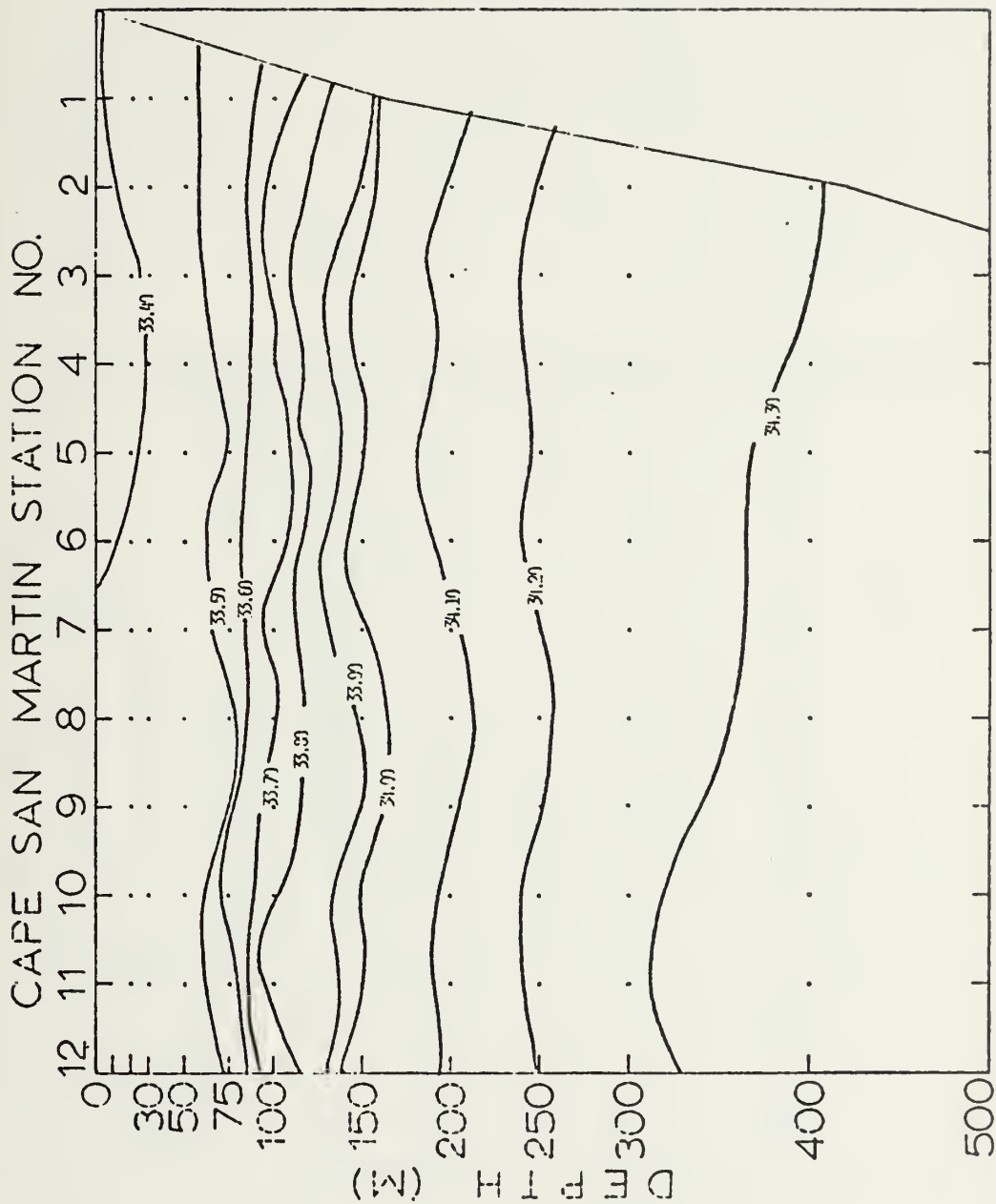


Figure 13. Salinity ( $\text{‰}$ ) on a vertical section for the Cape San Martin line on 21-22 February 1979.





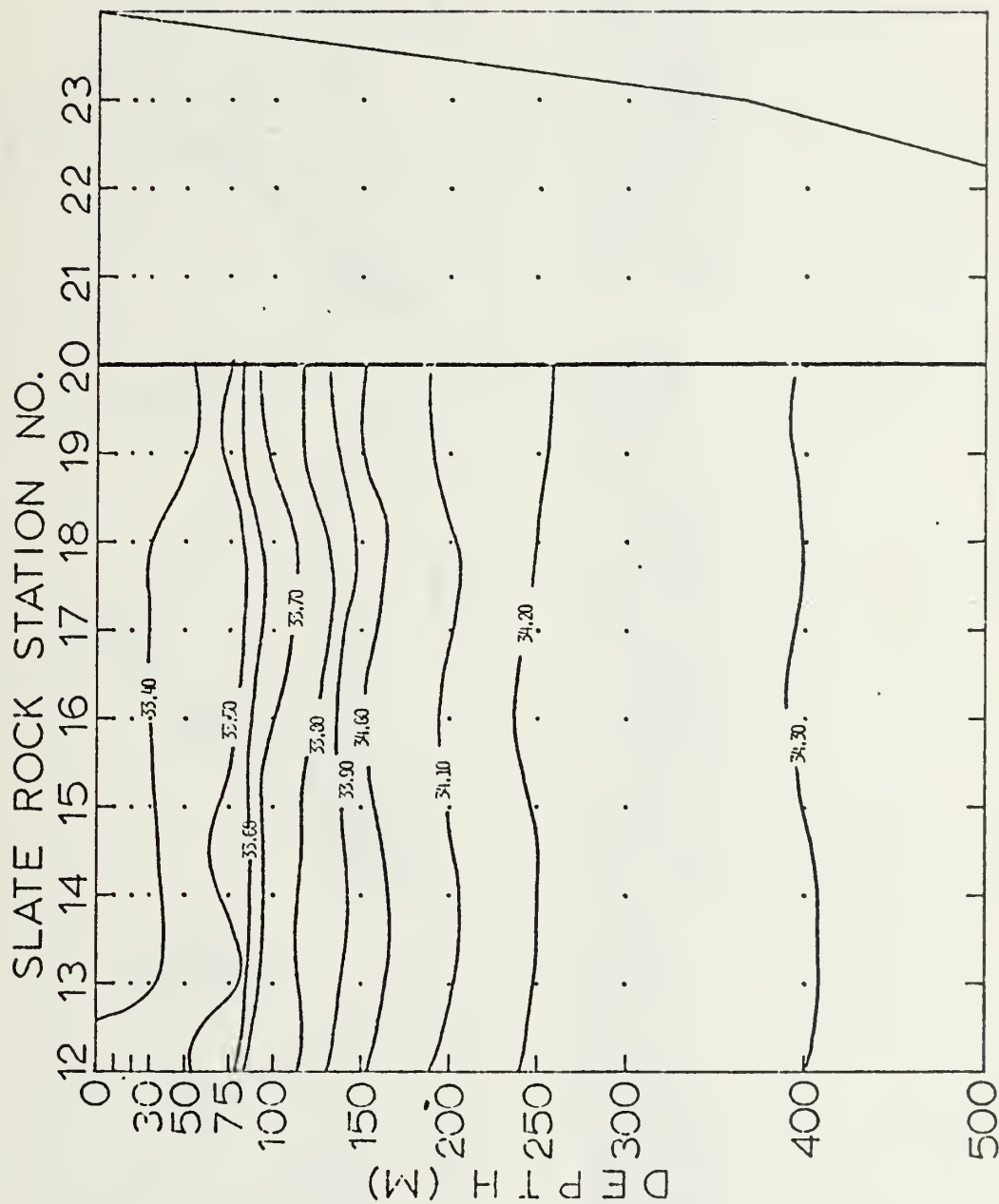


Figure 14. Salinity (‰) on a vertical section for the Slate Rock line on 21-22 February 1979.







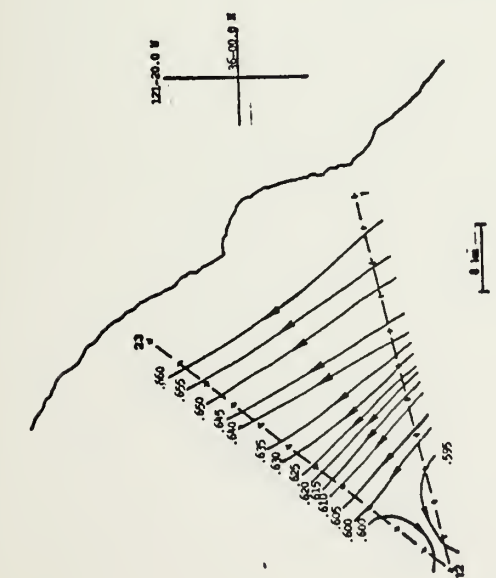


Figure 18. Dynamic topography of the 100/500 decibar surface on 8-9 January 1979. Units are in dynamic meters.

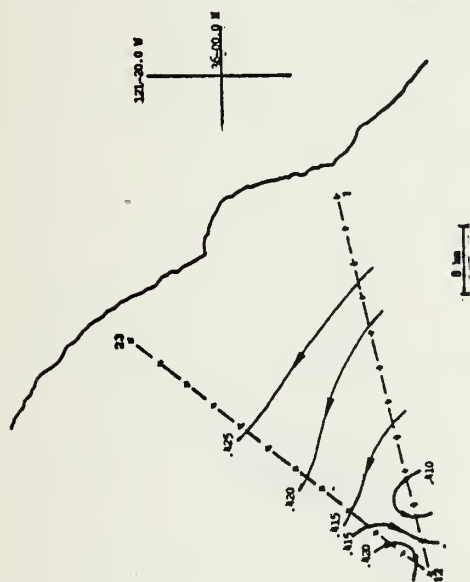


Figure 19. Dynamic topography of the 200/500 decibar surface on 8-9 January 1979. Units are in dynamic meters.

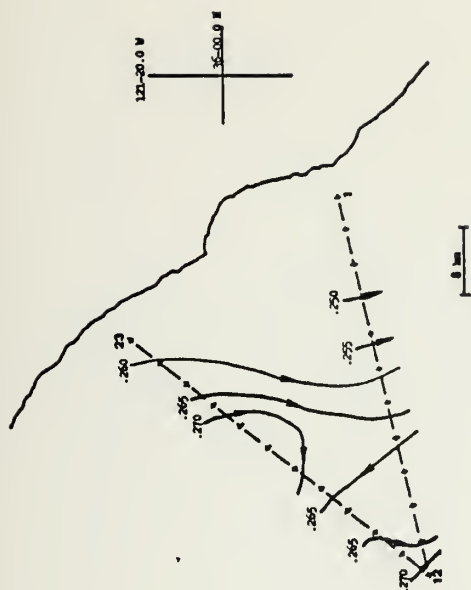


Figure 20. Dynamic topography of the 300/500 decibar surface on 8-9 January 1979. Units are in dynamic meters.









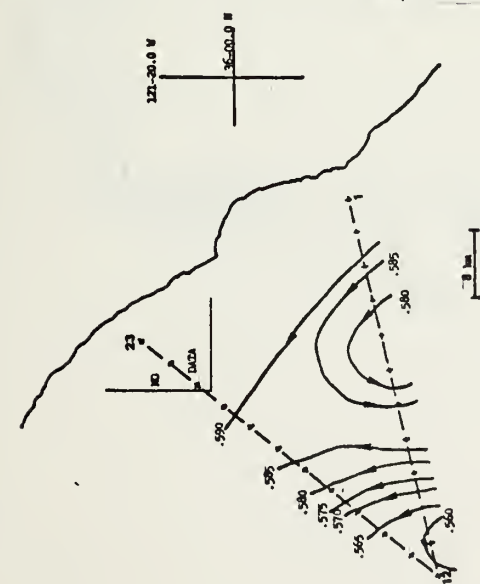


Figure 24. Dynamic topography of the 100/500 decibar surface on 21-22 February 1979. Units are in dynamic meters.

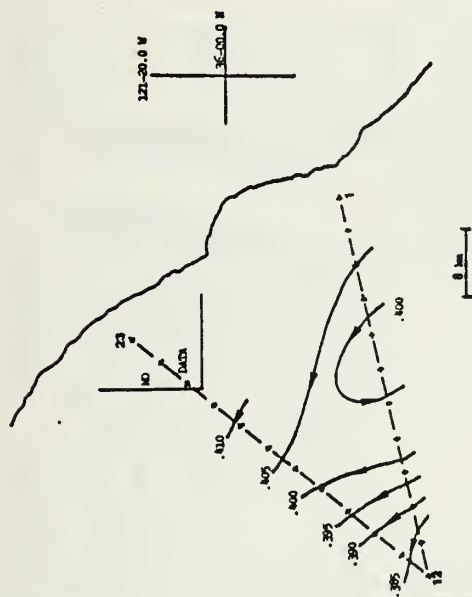


Figure 25. Dynamic topography of the 200/500 decibar surface on 21-22 February 1979. Units are in dynamic meters.

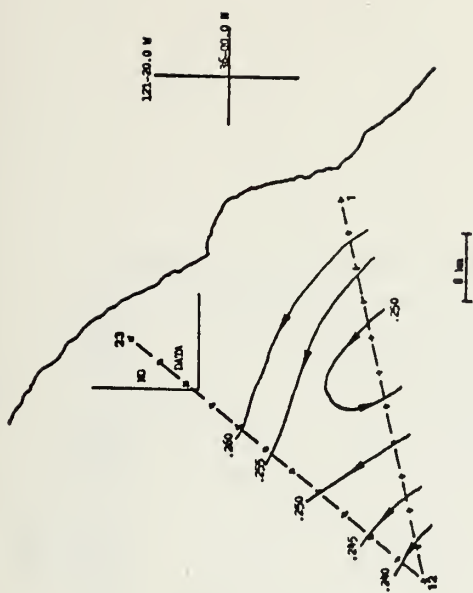


Figure 26. Dynamic topography of the 300/500 decibar surface on 21-22 February 1979. Units are in dynamic meters.



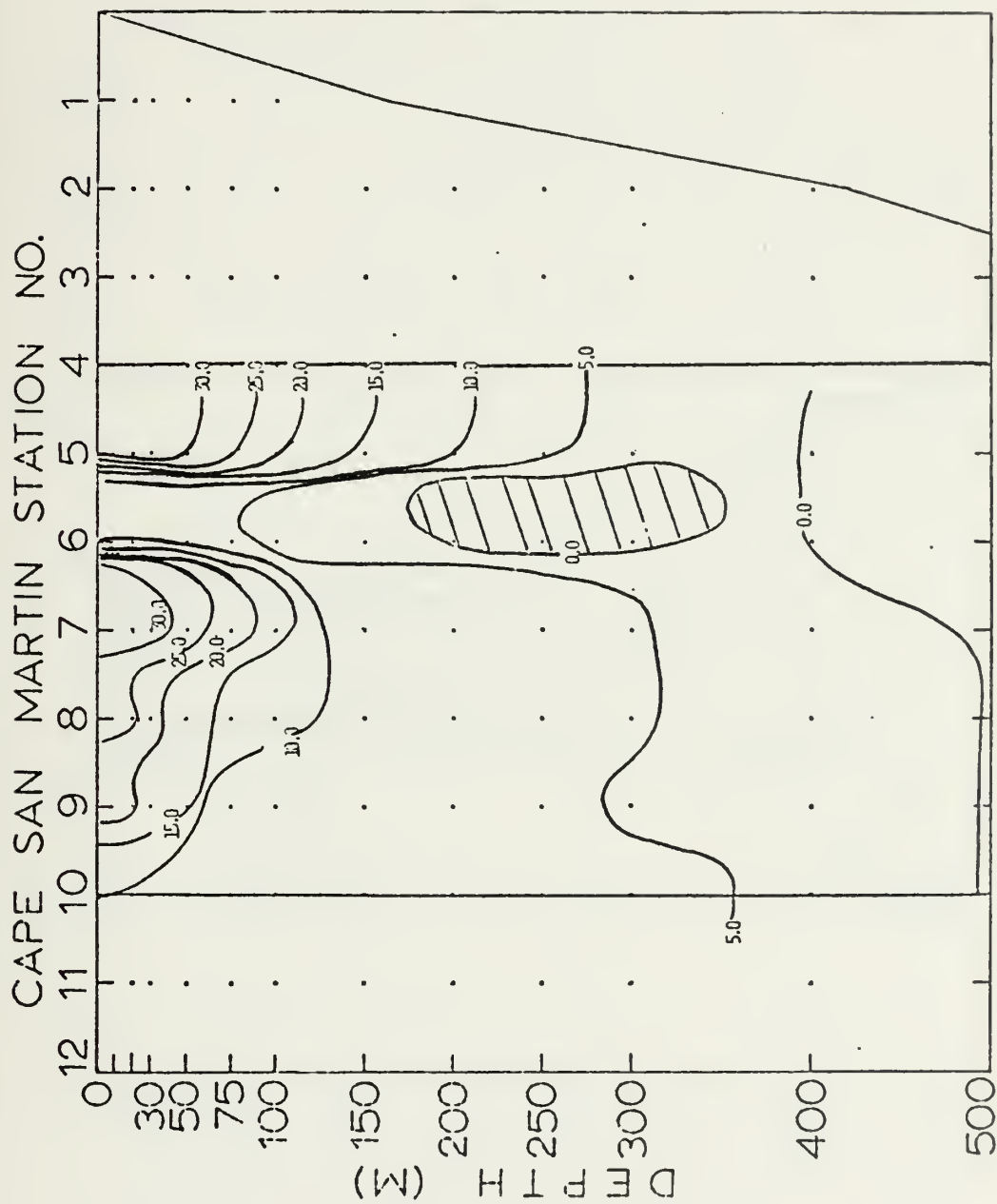


Figure 27. Vertical section of the normal component of geostrophic velocity in cm/sec for the Cape San Martin line on 27-28 November 1978. Southward flow indicated by cross hatched area.



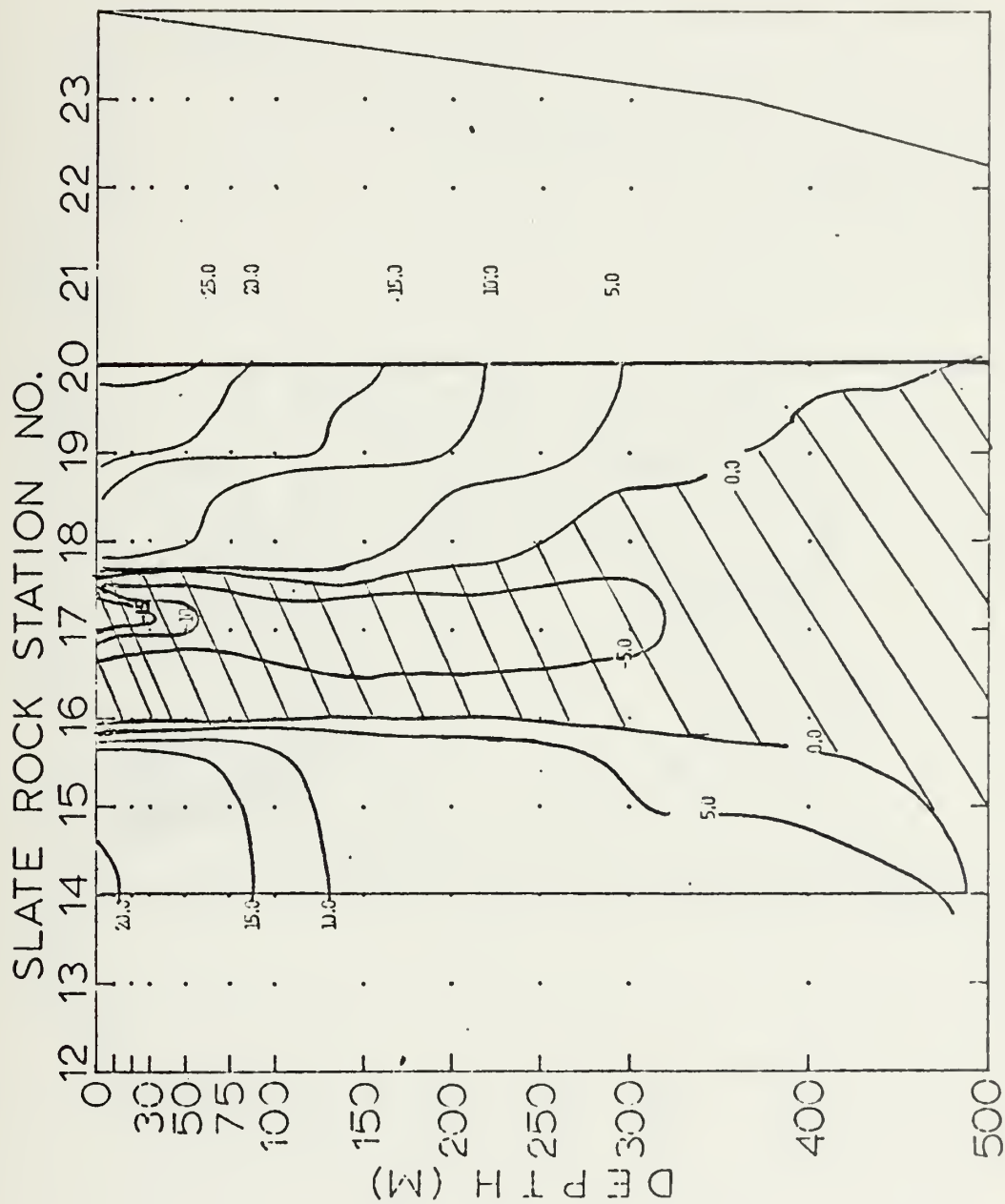


Figure 28. Vertical section of the normal component of geostrophic velocity in cm/sec for the Slate Rock line on 27-28 November 1978. Southward flow indicated by cross hatched area.



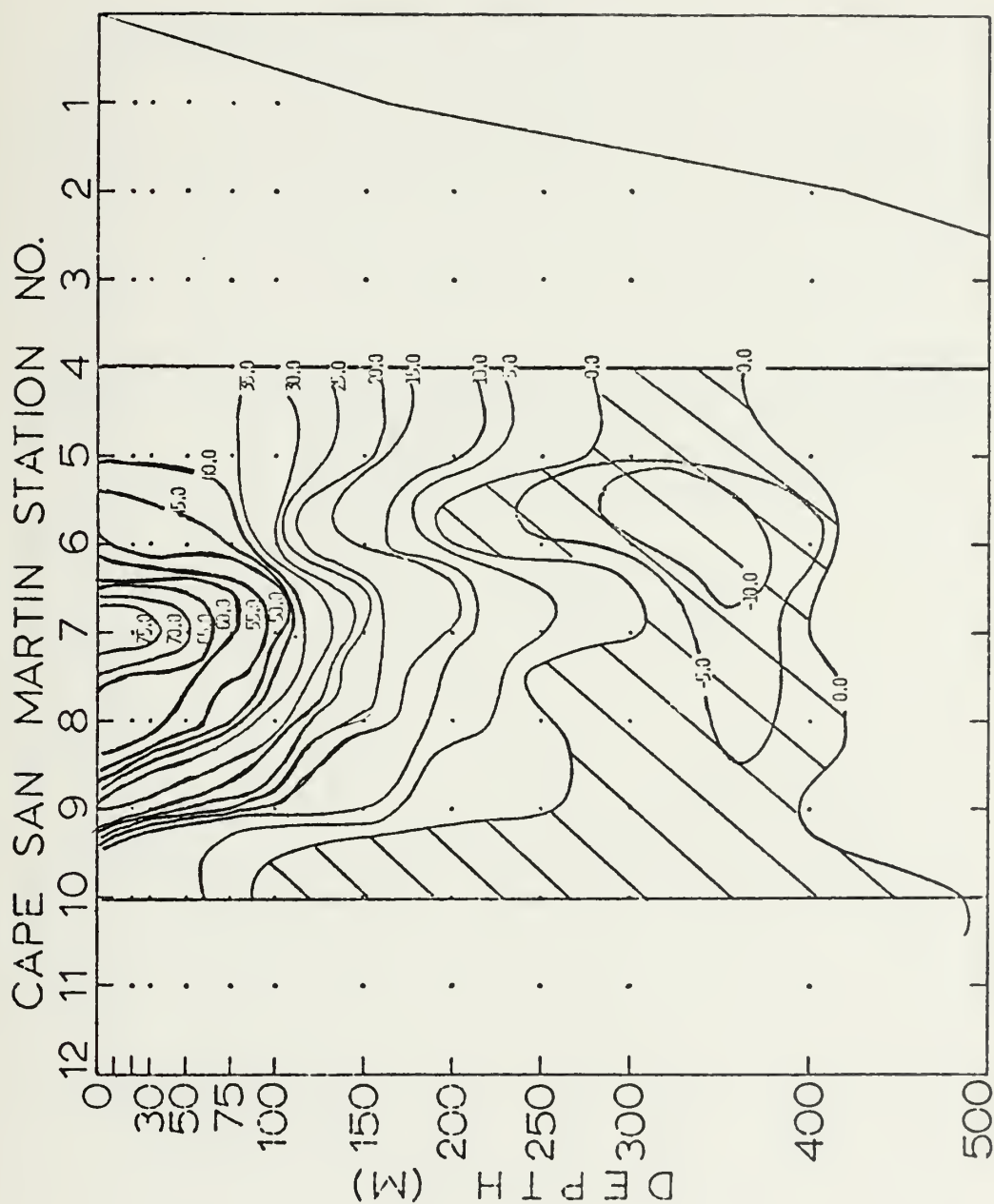


Figure 29. Vertical section of the normal component of geostrophic velocity in cm/sec for the Cape San Martin line on 8-9 January 1979. Southward flow indicated by cross hatched area.





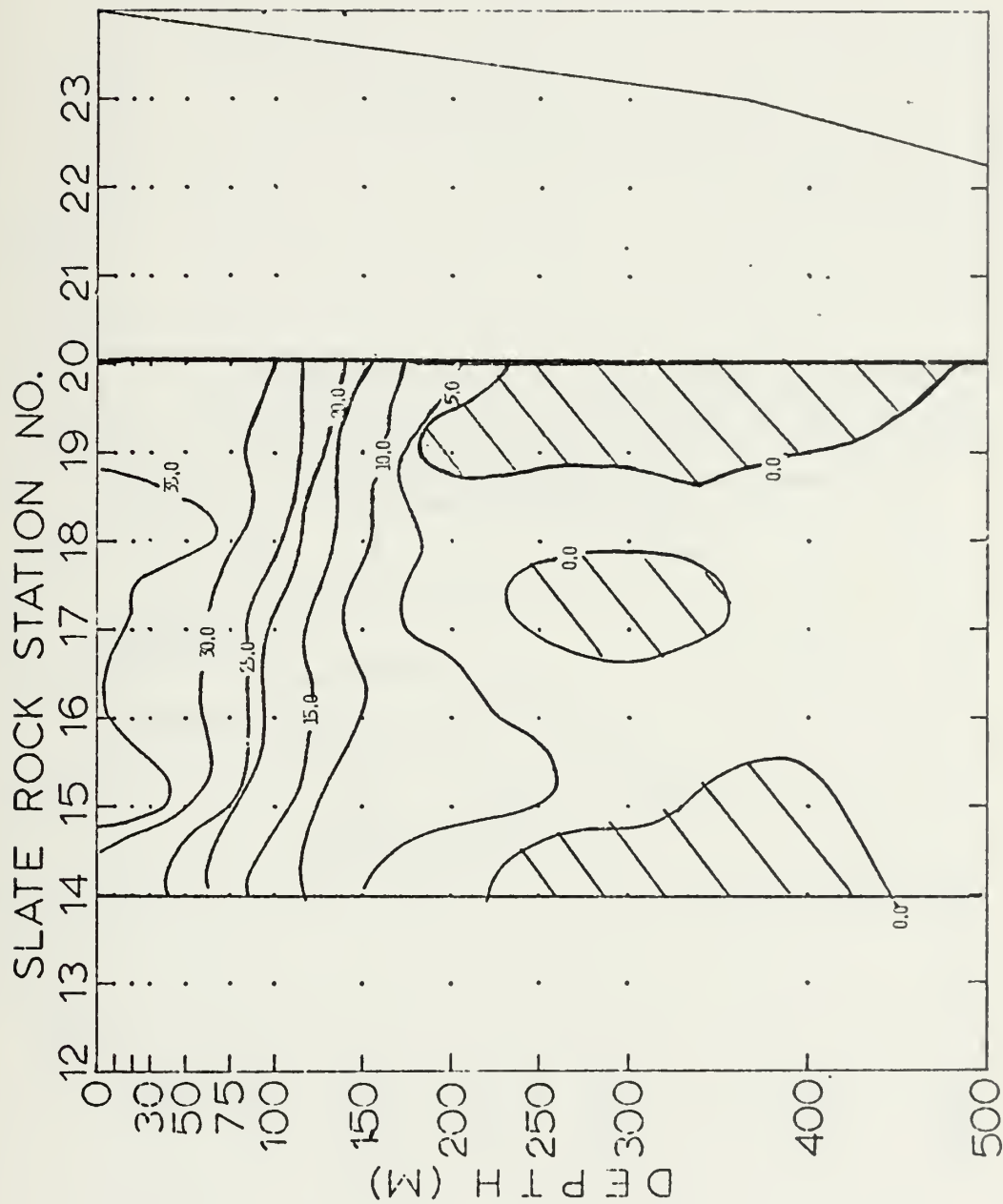


Figure 30. Vertical section of the normal component of geostrophic velocity in cm/sec for the Slate Rock line on 8-9 January 1979. Southward flow indicated by cross hatched area.



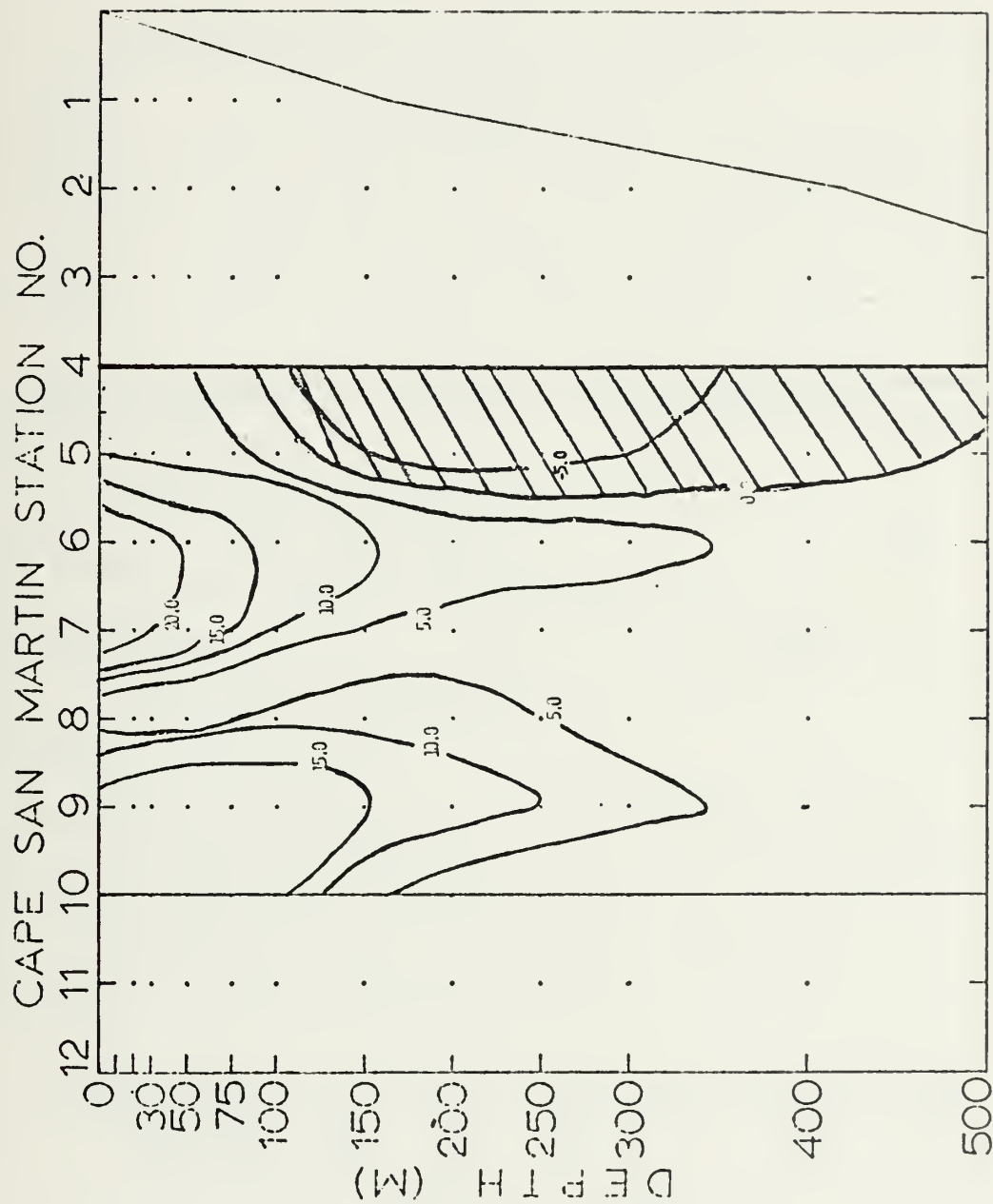


Figure 31. Vertical section of the normal component of geostrophic velocity in cm/sec for the Cape San Martin line on 22-23 January 1979. Southward flow indicated by cross hatched area.



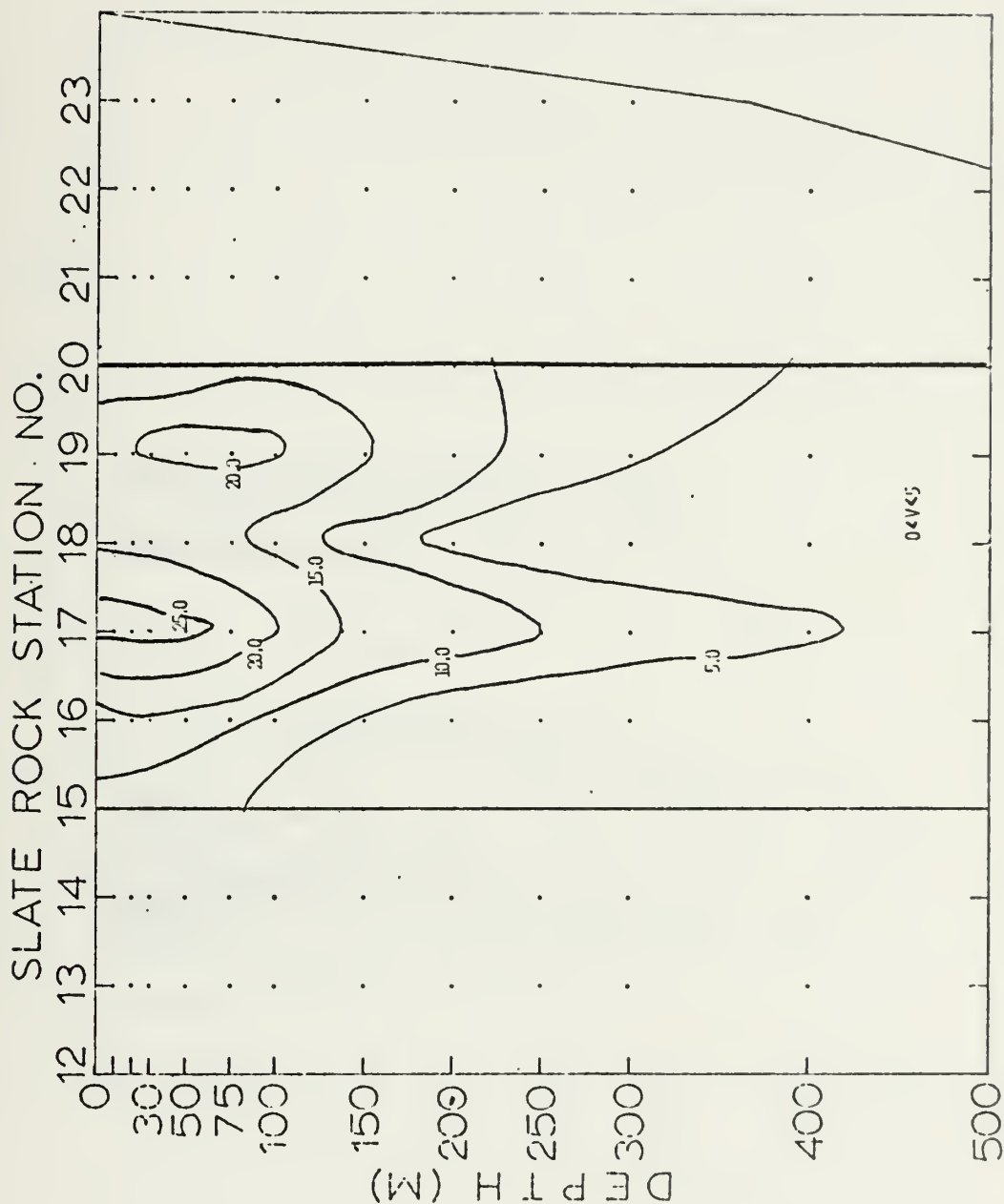


Figure 32. Vertical section of the normal component of geostrophic velocity in cm/sec for the Slate Rock line on 22-23 January 1979. Southward flow indicated by cross hatched area.



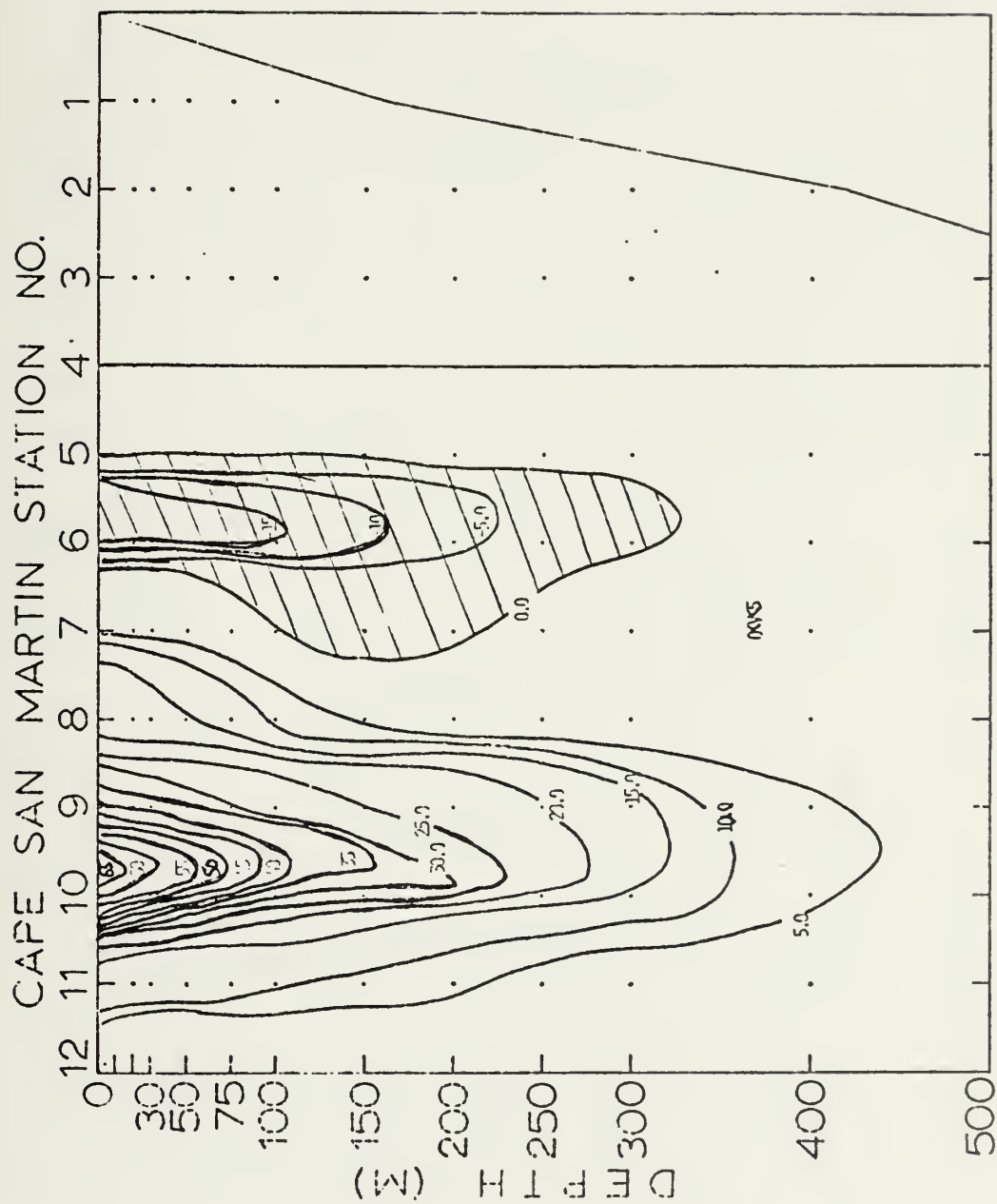


Figure 33. Vertical section of the normal component of geostrophic velocity in cm/sec for the Cape San Martin line on 21-22 February 1979. Southward flow indicated by cross hatched area.





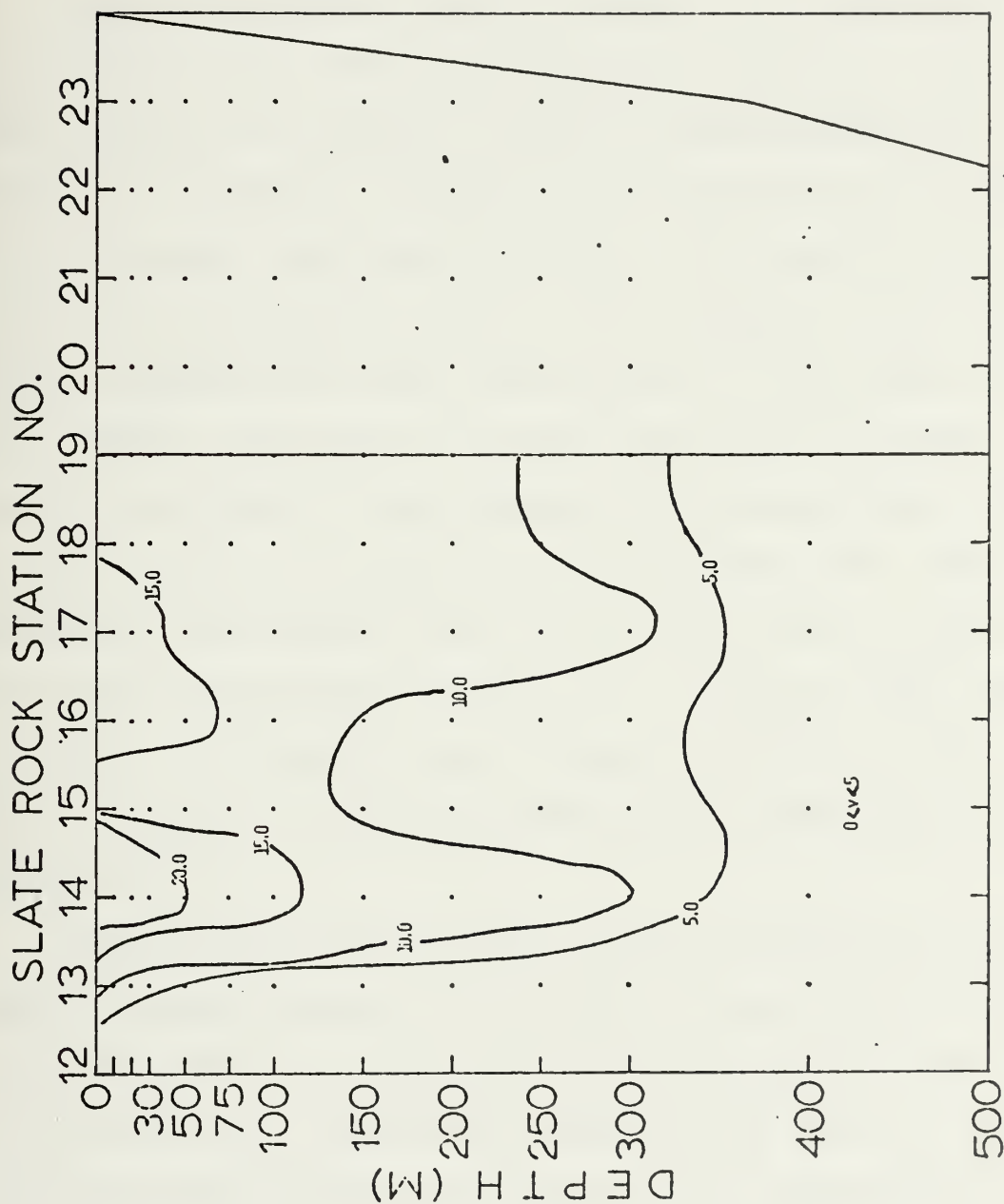


Figure 34. Vertical section of the normal component of geostrophic velocity in cm/sec for the Slate Rock line on 21-22 February 1979. Southward flow indicated by cross hatched area.



#### IV. DIRECT CURRENT OBSERVATIONS

##### A. INSTRUMENTATION AND DATA COLLECTION

Data collection was begun with the first array installed on 25 July 1978. This array contained one meter. The array was placed in 371 meters of water at approximately the position of station 2 (Figure 6) and remained in the water for one month. The second array was also placed near station 2 in 354 meters of water on 20 September 1978. The array contained one meter and remained in the water for two months. The meters from both the first and second array were placed at approximately 200 meters depth. Two larger arrays were installed on 27 November 1978. They contained three meters each and were installed in water of 353 meters depth near station 2 and 777 meters depth near station 5. The meters were situated at depths of 100, 175, and 300 meters. Both arrays remained in place until 22 January 1979.

Organization of each array was done by means of the NOYFB computer program. The original program was written by Donald Moller of the Woods Hole Oceanographic Institution in Fortran II for the Hewlett-Packard 2100 series. The program has been rewritten in Fortran IV for compatibility with the IBM 360 computer and is given in Appendix A. Through the use of on line terminals the program gives the operator a description of the mooring and its performance from an



operational point of view. Given environmental data and mooring components, the program presents the operator with mooring behavior information for evaluation of collected data as well as array design modifications. Step by step instructions for program operation are displayed to the user with the sequence determined by his selection of options. Standard mooring component characteristics (buoyancy, cross-sectional areas, elastic properties, etc.) are written into the program. These components are those used by Woods Hole Oceanographic Institution. Components which are non-standard with respect to the program are added during initialization or may be changed by an option.

The current meter used in these arrays is the Aanderaa Recording Current Meter Model 4 (RCM4). It is a self-contained instrument for recording speed, direction, and temperature of ocean currents with conductivity and pressure options. The RCM4 has a depth capability of 2000 meters. The RCM4 uses a rotor type current speed sensor, a magnetic compass, and a thermistor. An electro-mechanical encoder (analog to digital converter) samples and converts the measurements to binary digital signals which are then recorded on 1/4-inch magnetic tape. A sampling interval of 10 minutes was chosen for our meters. Input parameters for the RCM4 into the NOYFB program are:  $W(I) = -64.66$  (buoyancy per meter length) and  $A(I) = +0.065$  (area of component in square meters per meter length).



The array is arranged as in Figure 35 with no surface markers in order to keep all array elements out of the region of strong surface wave action. An acoustic release is used for retrieval of the arrays. It is an AMF Acoustic Release/Pinger Model 242, which is activated acoustically and also has a reply pinger used for interrogation and, if necessary, as a locating beacon.

Buoyancy for the array is provided by Benthos glass spheres. The glass spheres are 17-inches in diameter, housed in plastic hard hats and are capable to a depth of 6700 meters. They provide 55 pounds of buoyancy each and are connected together in pairs with 3/8-inch galvanized chain.

The wire finally used was 5/32-inch, 7x7 stainless steel wire, breaking strength about 2400 pounds, and stainless Nicopress fillings around plastic thimbles. Although the initial arrays used 1/4-inch galvanized wire with copper Nicopress fittings for terminations. Zinc anodes were attached to all Nicopress fittings, following a technique developed at Oregon State University (OSU). After one month of immersion the zincs on some of the wire fittings had almost totally disappeared. It was feared that a longer use would have led to corrosion of the copper Nicopress fittings and possibly to the loss of the array. It was then decided to use stainless steel wire. Again, following the technique developed at OSU, plastic 1/2-inch thimbles are used at all eye terminations. Galvanized anchor shackles with stainless steel cotterpins were used as connectors for





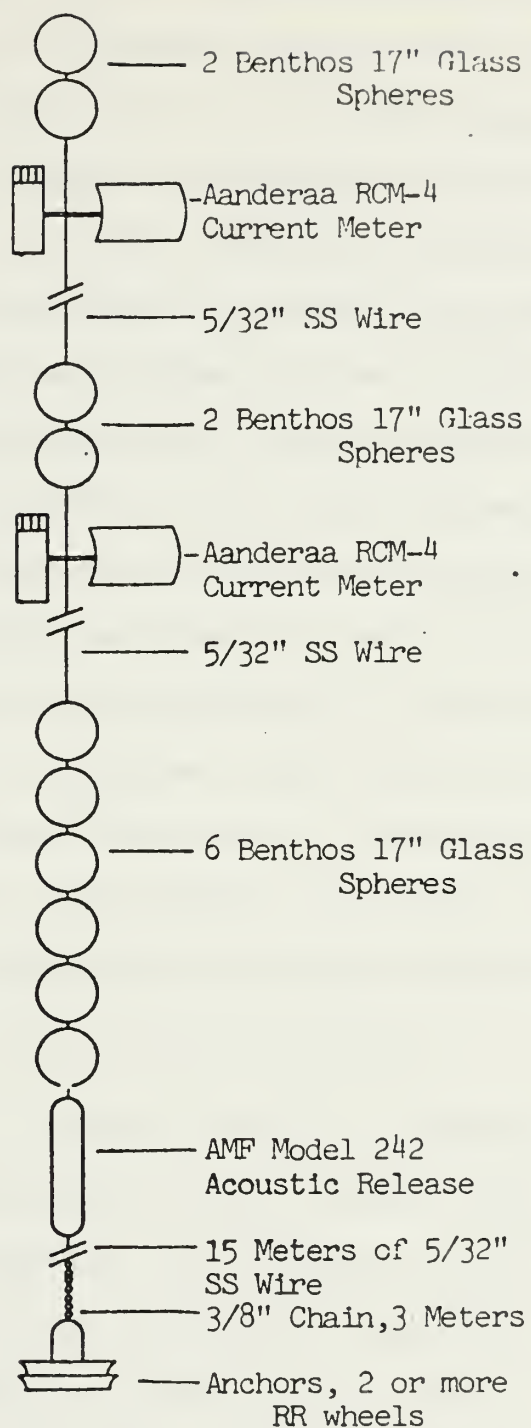


Figure 35. Array configuration.



the various components. Zincs are maintained on the meters and release as a protection against their corrosion. Input parameters for the wire into the NOYFB program are:

$W(I) = -0.1225$  (weight of component in pounds per meter length),  $A(I) = +0.00397$  (diameter in meters),  $RBS(I) = +2400.0$  (rated breaking strength in pounds), and  $AW(I) = +0.065$  (cross sectional metal area of wire in square inches).

Below each release is a 15-meter section of wire, 3-meter section of 3/8 inch chain and the weights. Weight is provided by scrap railroad wheels which have a weight of about 600 pounds per wheel in water. Two wheels are used for the shallow arrays and three for the deeper arrays.

Placement of arrays was done from the R/V ACANIA by the method of stringing the array out behind the ship, upper-most components first, and dropping the anchor last. As an added precaution a 7-foot cargo parachute was attached to the anchors to slow their rate of descent.

## B. DISCUSSION OF DIRECT CURRENT MEASUREMENTS

All data obtained with the current meters was converted to speed and direction using Aanderaa calibration constants and consolidated on a single IBM compatible tape.

Current speed and direction were used to construct progressive vector diagrams. Vertical and horizontal scales are equal in these diagrams. The vertical axis is zero degrees, magnetic north. The coast in the study area is aligned approximately in a  $305^\circ$  magnetic direction.



The current at station 2 during the period 25 July to 28 August 1978 is depicted by the progressive vector diagram shown in Figure 36. This meter was at 220 meters depth. During its operation the average overall current was  $336^{\circ}$  magnetic at 7.6 cm/sec. With respect to the alignment of the coast this shows a slight onshore component.

The semidiurnal components in the currents are visible upon close examination (the crosses on the diagram appear at two-day intervals). Two reversals of current which last four-days each can be seen at the top of the figure.

The second meter was installed in September (Figure 37) at the same location, station 2, but at a depth of 190 meters. It was in operation from 20 September to 27 November 1978. For this period the progressive vector diagram is shown in Figure 37, and the crosses are located at three-day intervals. The mean flow was  $318^{\circ}$  magnetic at 2.9 cm/sec. During the middle of the observation period numerous reversals of flow appear. These oscillations suggest that the meter was near a frontal zone and that its position alternated between sides of that zone, perhaps as a meander or eddy in the counter-current moved through the station's position. The mean flow over the seventy days is dominated by the countercurrent flowing nearly parallel to the coast.

Figures 38 through 42 are progressive vector diagrams for the arrays in operation from 27 November 1978 to 22 January 1979 on moorings at stations 2 and 5. The crosses



in all these figures are at three-day intervals. Semi-diurnal components are again visible.

Figures 38, 39, and 40 are for the array at station 2. The upper meter was at 100 meters depth (Figure 38). The mean flow at this level was toward  $337^{\circ}$  magnetic at 12.8 cm/sec, a moderately strong onshore and northward flow. On the 8th of January the flow changes from northward to southward for three days. The next meter was at 175 meters depth (Figure 39). The mean flow was toward  $321^{\circ}$  magnetic at 3.2 cm/sec. This small mean value is deceiving; it results from a period of mainly southward flow during the first nine days and oscillations during the last part of the period. The speed during the middle of the period was at a maximum of 25 cm/sec during a three-day period. Thus, although the oscillations result in a small value for the local time-averaged flow, weakening of the countercurrent is not necessarily implied. The oscillations may simply imply a spatially fluctuating front in the region.

The bottom meter on this array was at 300 meters (Figure 40). The mean flow at this level was towards  $151^{\circ}$  magnetic at 1.8 cm/sec. During most of the period the flow appears weak and fluctuating. The oscillations are mainly in two directions, northward and southward parallel to the coast. The meter appears to be astride the front separating the countercurrent and the southward flow. As the countercurrent meanders the meter moves from one flow to the other. In summary the current is moderately strong and northward at





the surface and weak at mid-depth with the exception of the strong northward flow during the middle of the period. The oscillations during the first and last days and at the lower meter indicate a moving shear zone between the countercurrent and the southward flow.

Figures 41 and 42 are for the array installed at station 5 and operating in the interval 27 November 1978 through 22 January 1979. The upper meter was at 140 meters depth (Figure 41). The mean flow was toward  $319^{\circ}$  magnetic at 12.3 cm/sec. As can be seen, the flow is almost exclusively northward, parallel to the coast, during the entire period. The exception is a three-day interval at the start with southward flow.

At 215 meters (Figure 42) the mean flow was toward  $308^{\circ}$  magnetic at 6.3 cm/sec. This reduction of the mean from the upper meter is largely due to the southward flows during the first nine days and an eight-day period beginning about 1 January. During these reversals the flow maintains its new direction and intensity without multiple oscillations. The sharp, clear change in flow indicates reversal from countercurrent to southward flow with no lingering of the front in the region of the meter. This pattern is similar to that at station 2 with a weakening of flow with depth and the appearance of southward flow at the deeper meters.

In summary the mean flow at the surface during the period of 27 November 1978 through 22 January 1979 is northward with an onshore component and a surface maximum of speed. At mid-depth the current has maintained its direction



with reduced speed. At the deeper meters the flow is variable in direction and speed, indications of a frontal region separating northward and southward flow.



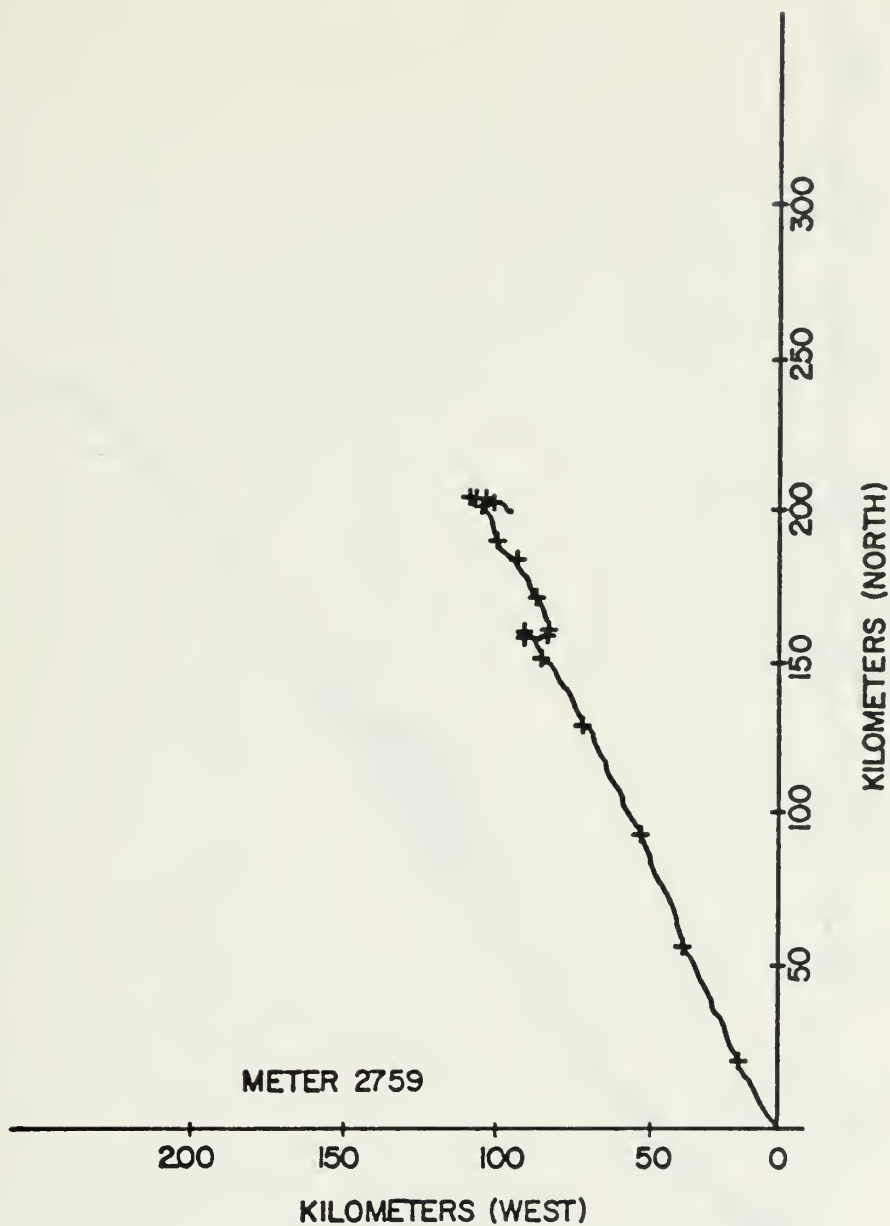


Figure 36. Progressive vector diagram for the current meter at 220 meters depth at station 2 from 25 July 1978 to 28 August 1978. Crosses are positioned at 2 day intervals. Vertical axis indicates Magnetic North.



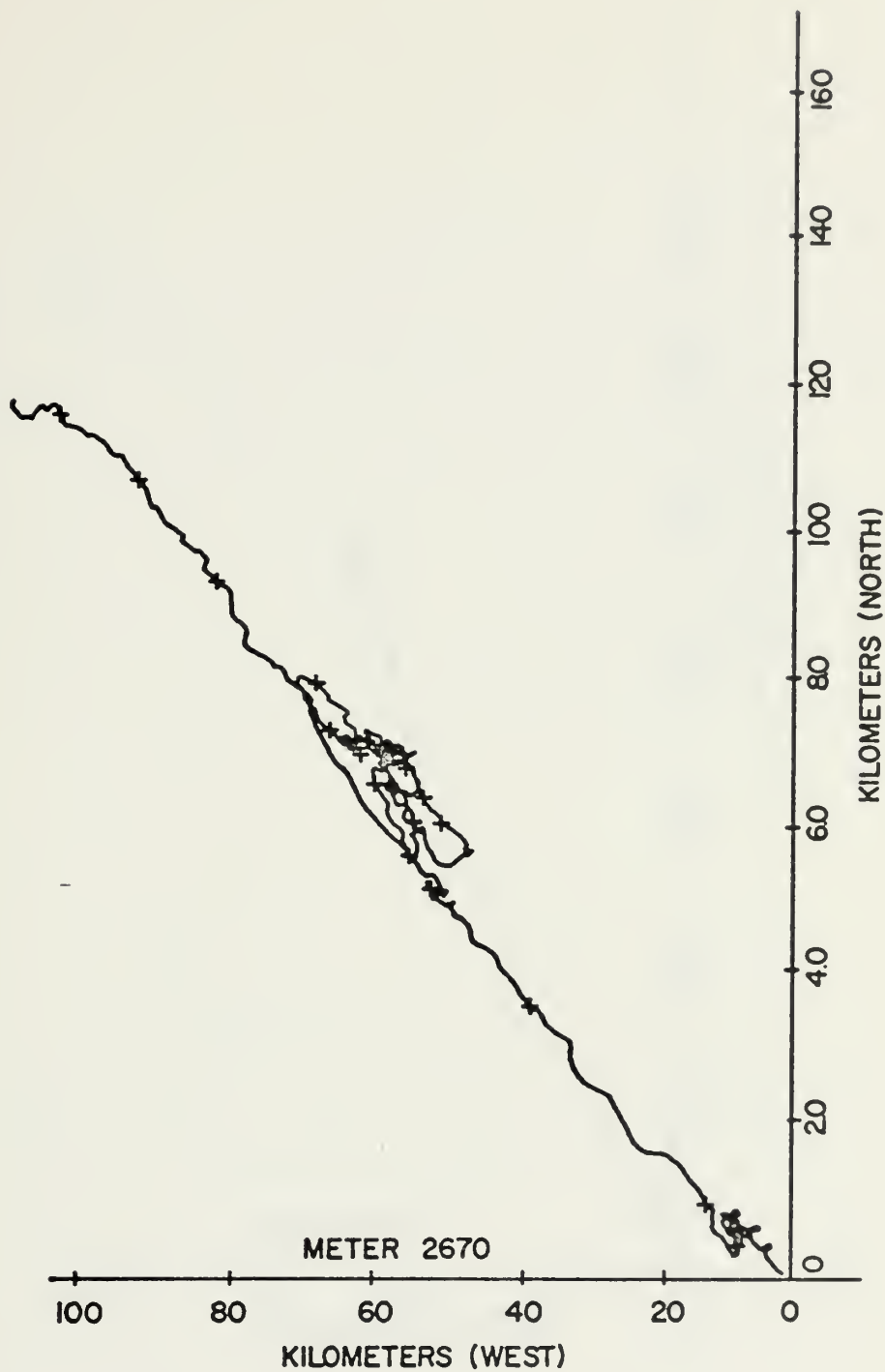


Figure 37. Progressive vector diagram for the current meter at 190 meters depth at station 2 from 20 September 1978 to 27 November 1978. Crosses are positioned at 3 day intervals. Vertical axis indicates Magnetic North.





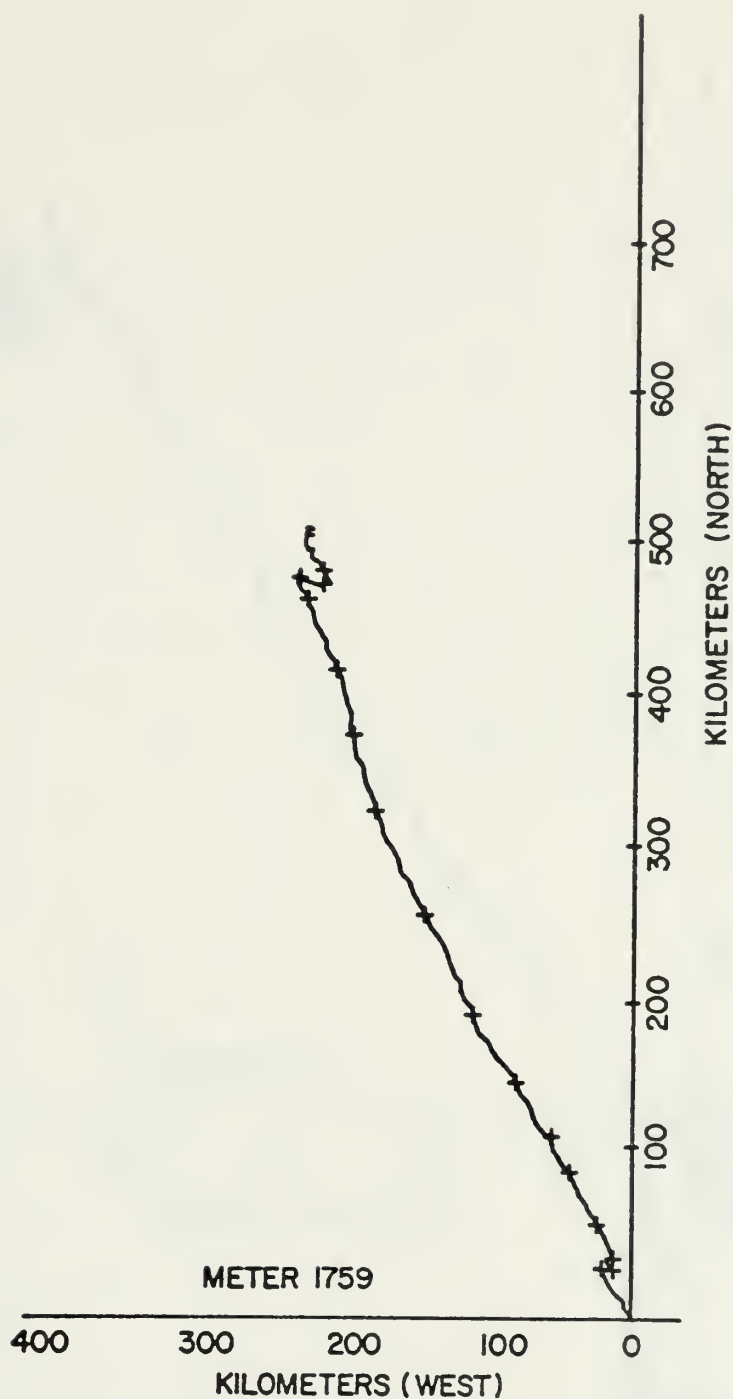


Figure 38. Progressive vector diagram for the current meter at 100 meters depth at station 2 from 27 November 1978 to 22 January 1979. Crosses are positioned at 3 day intervals. Vertical axis indicates Magnetic North.



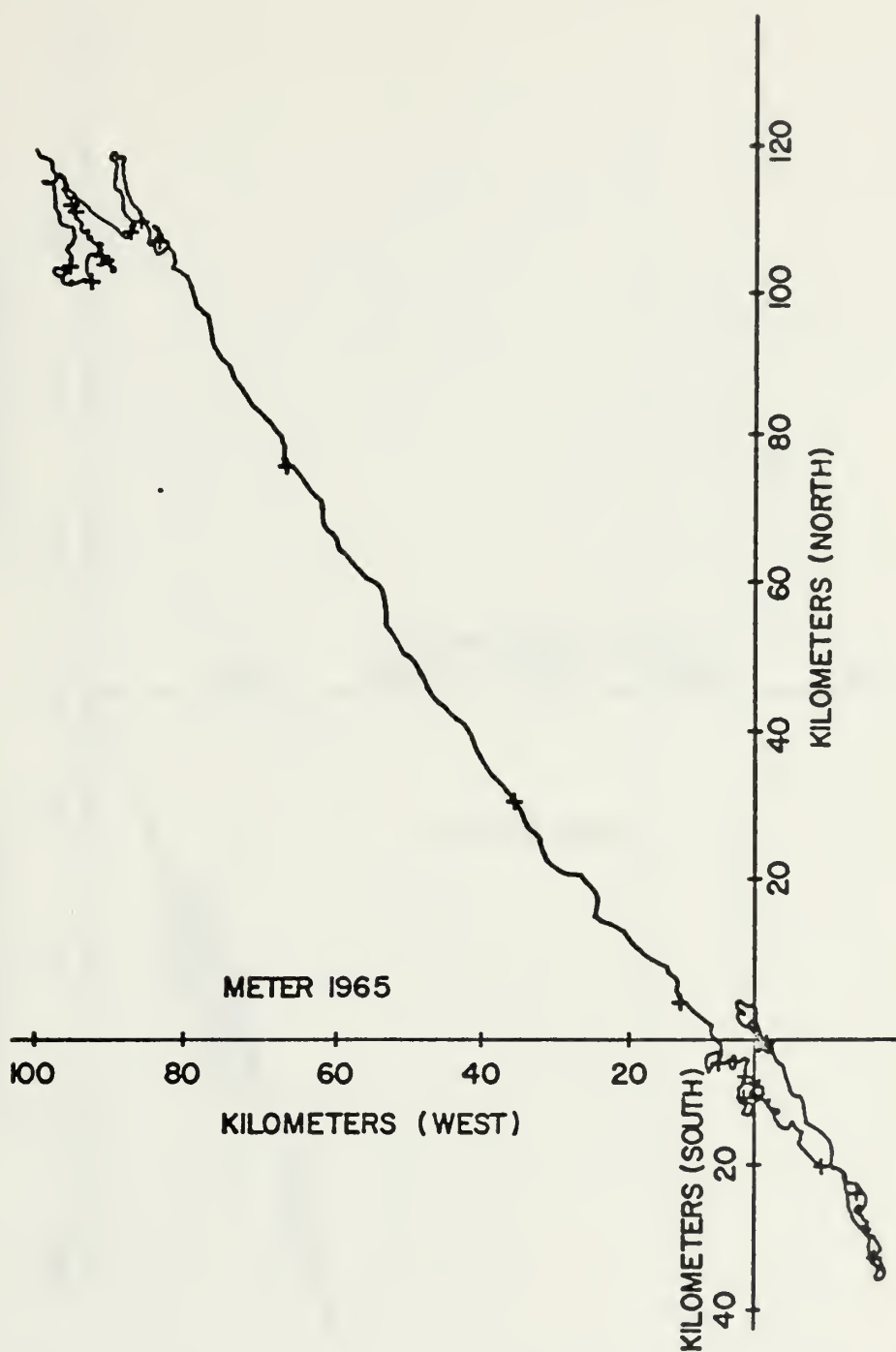


Figure 39. Progressive vector diagram for the current meter at 175 meters depth at station 2 from 27 November 1978 to 22 January 1979. Crosses are positioned at 3 day intervals. Vertical axis indicates Magnetic North.



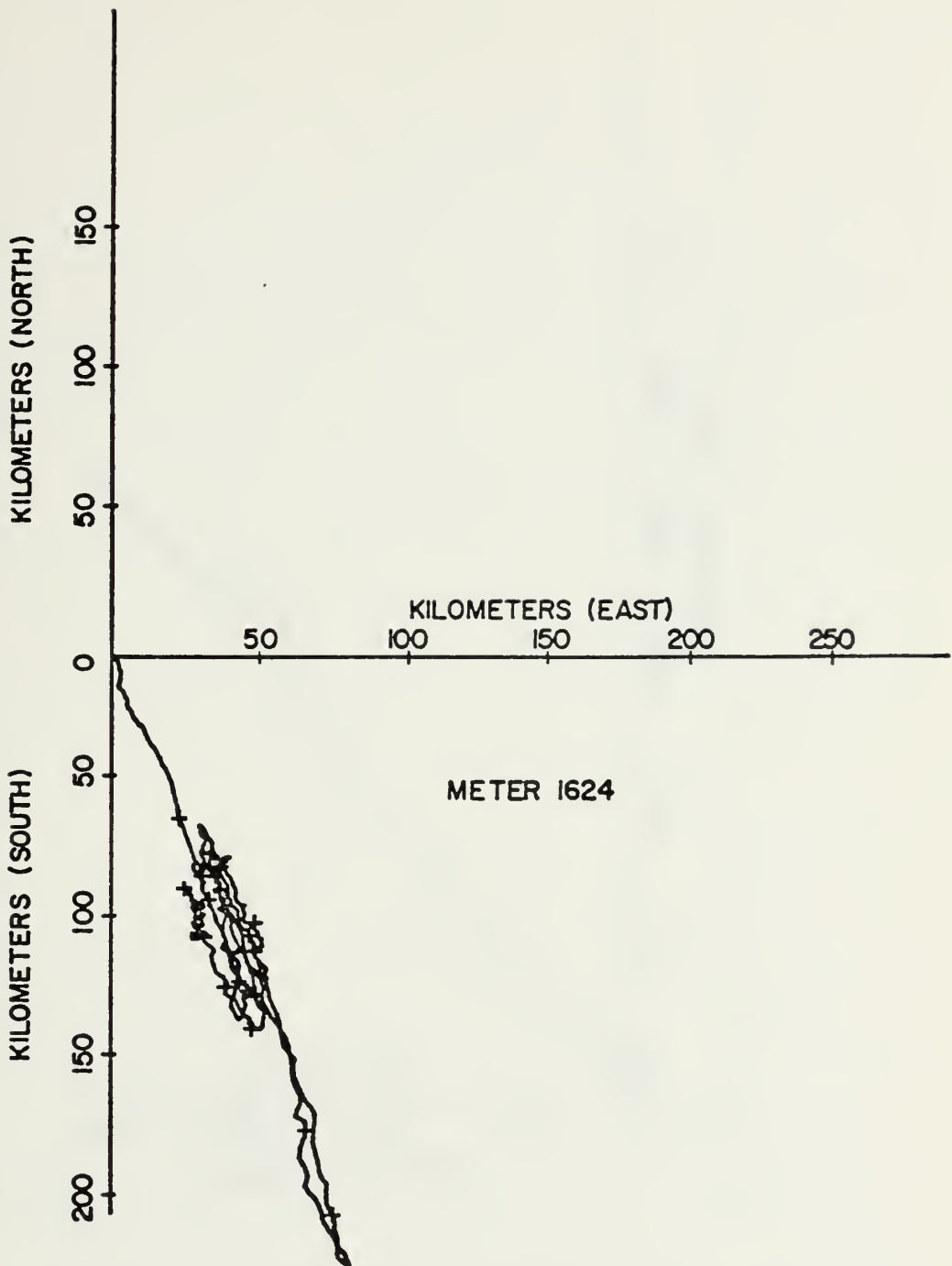


Figure 40. Progressive vector diagram for the current meter at 300 meters depth at station 2 from 27 November 1978 to 22 January 1979. Crosses are positioned at 3 day intervals. Vertical axis indicates Magnetic North.



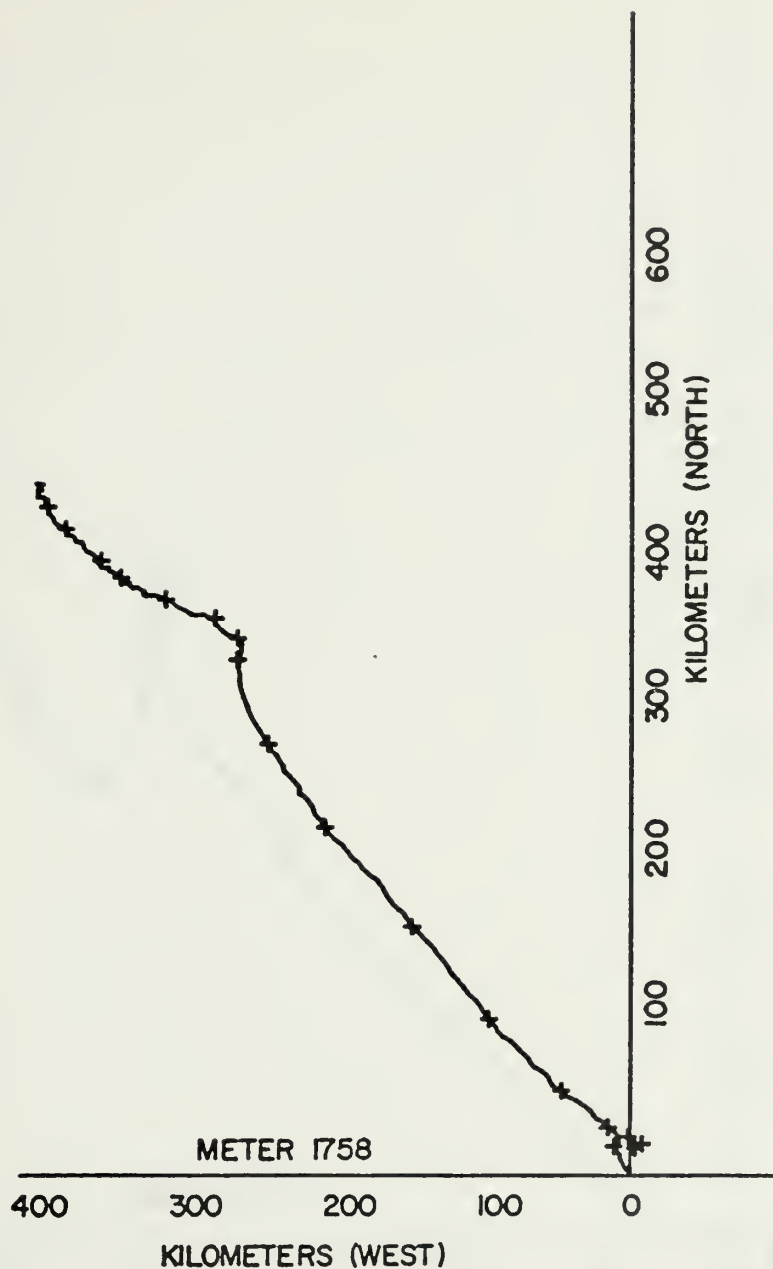


Figure 41. Progressive vector diagram for the current meter at 140 meters depth at station 5 from 27 November 1978 to 22 January 1979. Crosses are positioned at 3 day intervals. Vertical axis indicates Magnetic North.





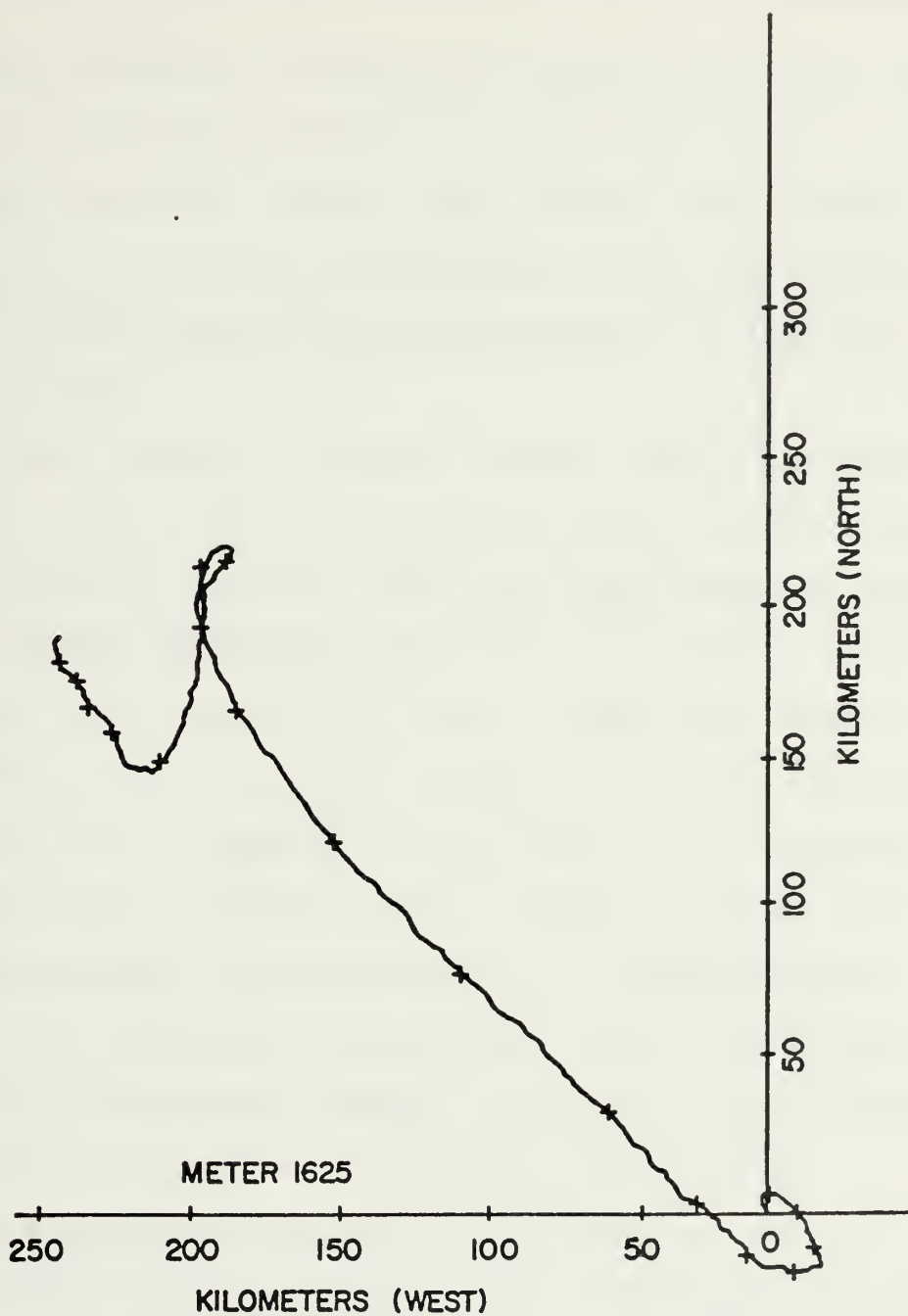


Figure 42. Progressive vector diagram for the current meter at 215 meters depth at station 5 from 27 November 1978 to 22 January 1979. Crosses are positioned at 3 day intervals. Vertical axis indicates Magnetic North.



## V. COMPARISON OF GEOSTROPHY WITH DIRECT CURRENT MEASUREMENTS

The two methods of inferring currents, indirect (geostrophy) and direct (current metering) coincided on 27 November 1978 and 8 January 1979. There were vertical profiling for watermass properties on the Cape San Martin line and simultaneous direct measurements at stations 2 and 5 of that line.

The mean velocity from the current meters was found for a 24-hour period bracketing the time during which watermass properties were sampled. This was accomplished by computing:  $\vec{V}(t) = \frac{\vec{R}(t+12) - \vec{R}(t-12)}{24}$  where  $\vec{R}$  is the position vector at the given time taken from a list of the values used to produce the progressive vector diagram. This 24-hour averaging was done so as to smooth the high frequency non-geostrophic contributions to the flow while retaining larger scales. It must be expected that this averaging process does not smooth out all non-geostrophic components. Those which remain contribute to differences between the direct (current meter) and indirect (geostrophic) velocity measurements.

The mean metered velocities at stations 2 and 5 on both 27 November 1978 and 8 January 1979 are compared to those velocities found by geostrophy in Table II.

As seen in Table II, on 27 November 1978 at station 2, there is general agreement on the direct and indirect measurement of velocities at 100 and 175 meters. At 300 meters the



methods of measurement indicate opposite flow directions. Dynamic topography indicates the region to be fairly flat with adjacent northward flow. The current meter shows a moderate southward flow of 10 cm/sec. This bottom meter is the meter which showed obvious non-geostrophic components, frequent oscillations between northward and southward flow. On the same data at station 5, the table shows strong agreement between the two methods of measurement. Geostrophy indicates a slightly stronger flow and the current meters show more of an onshore flow.

On 8 January 1979 at station 2, there is poor correlation of the current from the two measurement methods. Geostrophy indicates the region, at mid and lower depths, to be fairly flat with weak flow. Direction can only be inferred from that at station 4 (Figures 19 and 20). At the 100 meter level the current meter shows an onshore southward flow while geostrophy indicates northward flow parallel to the coast. At station 5 one can see reasonable agreement between the two methods of measurement (Table II). Current meters show uniform speed with depth and geostrophy indicates weakening of the flow with depth.

In summary there is general agreement between the two methods of measurement. It must be remembered that geostrophic calculations for station 2 were based on extrapolations and to that extent are uncertain. Further, the averaging of the current over 24 hours does not remove all non-geostrophic components, especially in the periods of



strong fluctuations. The geostrophic current, as expected, is a more useful estimate of the flow in this variable regime during the relatively quiet (i.e., steady) flow periods.





## VI. CONCLUSIONS

In summarizing the current and watermass variations, the time interval between cruises is used for an initial breakdown. Currents are described by geostrophy at the time of each cruise and by current meter measurements during the intervals. Watermass properties are known only at cruise times.

From 27 November 1978 to 8 January 1979, near the level of 100 meters, there is an increased halocline slope downward toward the coast and an intense northward flow spreading westward. This spreading flow also occurs at 200 and 300 meters. Below 300 meters, southward flow develops. A salinity maximum at 300 meters occurs near the center of the station lines in November and at the western edge in early January.

From 8 January 1979 to 22 January 1979 the halocline levels off and becomes less intense. At 100 meters the northward flow has weakened and at 200 meters remains unchanged. From 300 to 500 meters there is a reduction in salinity due to horizontal advection, vice vertical, of low salinity water. There is also reduced southward flow below 300 meters during this interval.

These results by indirect measurements are consistent with direct (current meters) methods. At 100 meters northward flow spreads westward from 27 November 1978 to 22



January 1979. At about 1 January this flow shows an increased westward component, i.e. it crosses contours into deeper water. Vertical extent of the northward flow also increases during this period. Southward flow is present at mid-depth until 6 December when it turns northward at both stations 2 and 5. At 300 meters, a frontal region separating northward and southward flow, the flow is southward until early December when much oscillating dominates the rest of the period.

Indications are that the flow is branched and northward at the surface, consistent with Wickham's (1975) findings farther north. Presence of a surface maximum, Davidson Current, during the winter is almost universally consistent with the observations of other investigators. Westward propagation of the flow tends to support McCreary's (1977) model.

From 22 January 1979 to 21 February 1979 there is an increase in salinity at all levels. The northward jet shows further westward movement.

Measurement of the currents is still in progress. A 3-meter array was installed farther west at station 7 and in April 1979 was increased to a 4-meter array. Arrays are also being maintained at stations 2 and 5. Along with current meter measurements, simultaneous measurements of watermass properties are made at regular intervals.



TABLE II

COMPARISON OF CURRENT METER AND GEOSTROPHIC CURRENT  
VELOCITIES ON 27 NOVEMBER 1978 AND 8 JANUARY 1979

Date	Station	Depth (meters)	Current Meters		Geostrophy	
			$\theta$ (1) direction (ref. Mag N)	V(1) (cm/sec)	$\theta$ (2) direction (ref. Mag N)	V(3) (cm/sec)
27 November	2	100	331°	16.2	310°	strong
		175	352°	2.9	305°	moderate
		300	112°	10.0	335°	weak
27 November	5	140	337°	9.5	310°	15.1
		215	339°	4.7	305°	9.1
8 January	2	100	076°	4.8	310°	strong
		175	249°	6.8	variable	weak
		300	332°	38.3	variable	weak
8 January	5	140	300°	16.6	310°	25.1
		215	240°	16.9	285°	weak

NOTES

- (1)  $\theta$  and V (current Meters) found using  $\frac{\vec{R}(t+12) - \vec{R}(t-12)}{24}$  where  $\vec{R}$  is the position vector at the given time taken from a list of values used to produce the progressive vector diagram.
- (2)  $\theta$  (geostrophy) found from dynamic topographies.
- (3) V (geostrophy)  $V = \frac{V_n}{\cos \theta}$  where  $V_n$  is the normal component of velocity from geostrophic cross sections and  $\theta$  is the angle from the normal.



# APPENDIX A

```

PROGRAM NOYFB
PROGRAM NYOFB, REVISED VERSION OF:
REVISION 9.1, MARCH 1975
REVISED OCTOBER 1978 BY K. CODDINGTON
FOR USE WITH IBM 360 COMPUTER
WRITTEN IN FORTRAN IV LANGUAGE
REQUIRES SUBROUTINES CALC,CONST,TRANS,PTIO
REQUIRES FUNCTION JCD

DETERMINES THE STATIC CONFIGURATION OF SUB-SURFACE
MOORINGS WHEN ACTED UPON BY NON-COPLANAR CURRENT PROFILES

MAIN PROGRAM CONTROLS I/O AND OPTION SELECTION
SENSE SWITCH OPTIONS ARE INDICATED BY ISSW(I) WHERE 1=ON AND 2=OFF
S.S. - 1 - ON - LIST INPUT PARAMETERS
      - 2 - ON - OUTPUT SEGMENT STATS
      - 3 - ON - OUTPUT SUPPLEMENTAL STATS
      - 4 - ON - OUTPUT SUMMARY MOORING STATISTICS
      - 5 - ON - OUTPUT COMPONENT CHARACTERISTICS
      -10 - ON - OUTPUT TO SOFT COPY DEVICE
      -10 - OFF- OUTPUT TO HARD COPY DEVICE
      -14 - ON - ABORT RUN - GO TO 'PAUSE'

COMMON W(42), A(42), RBS(24), AW(5), E(16)
COMMON IT(65), XL(65), TW(66), CDN(5), CDT(5)
COMMON DCP(20), CP(20), RCP(20), IHDG(36)
COMMON T1(65), T2(65), T3(65), T4(65), T5(65), T6(65), T7(65)
COMMON IDA(10), IDM(10), DB(10)
COMMON DPT, QN200, D200, FPM, L11, LC, LC3, NEW, TERMW, TERML
COMMON RD, PI, TL, IC, IV, FUDG, RC, SEG, DONE
COMMON TCLIN, TTENS, TROT, QBKUP, TVELD, ANCR, ANCR4
COMMON ISSW(14)
ISSW(1)=1
ISSW(2)=1
ISSW(3)=1
ISSW(4)=1
ISSW(5)=1
ISSW(6)=1
ISSW(7)=1
ISSW(8)=1
ISSW(9)=1
ISSW(10)=1
ISSW(11)=0
ISSW(12)=0
ISSW(13)=0
ISSW(14)=0

```

CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC





```

DATA IBLANK/' '
LL = 0
IC = 0
NEW = 15
ISLEW = IBLANK

C SET STANDARD WHOI COMPONENT CONSTANTS
C
C CALL CONST
C
C ENTER I/O UNIT REFERENCE NUMBERS
C STANDARD U.R.N. FOR WHOI/BUOY COMPUTER ARE LISTED
C
C WRITE (6,1)
C FORMAT (10,ENTER FIVE I/O DEVICES',/, ' SOFT COPY,HARD COPY,PAPER PUNY
1 INCH,KEYBOARD,PAPER READER',/, ' STANDARD URN ARE: 6,8,2,5,2')
600 READ (5,600) L01, L02, L03, LI, LI1
C FORMAT (511)
C WRITE (L01,46)
46 FORMAT (10,INITIAL RUN FROM P.T.?: 1-YES, 0-NO')
601 READ (LI, 601) I
C FORMAT (12)
C IF (I-1) 59,55,55
C
C INPUT MOORING COMPONENTS AND CONSTANTS FROM PAPER
C TAPE OR OTHER SOURCE
C
55 CALL PTIO
51 WRITE (L01,51) DPT
C FORMAT (F6.1)
C LL = 0
C GO TO 105
C
C THE FOLLOWING SECTION PERMITS INPUT OF VARIABLE
C PARAMETERS FOR THE INITIAL AND SUBSEQUENT RUNS
C FORMAT STATEMENTS DESCRIBE THE OPERATIONS
C
59 CONTINUE
60 WRITE (L01,2)
2 FORMAT (10,CHANGE COMP. CONSTANTS?: 1-YES,0-NO')
602 READ (LI, 602) I
C FORMAT (12)
602 IF (I-1) 75,61,61
C WRITE (L01,3)
61 READ (LI, 603) I,J,X
62 FORMAT (10,ENTER CODING:1-W(I),2-A(I),3-RBS(I),4-AW(I)',/, ' THEN
3 IPE COMPONENT CODE NO. AND NEW VALUE')
603 FORMAT (I2, I1, I2, F10.5)

```







NY001450  
 NY001460  
 NY001470  
 NY001480  
 NY001490  
 NY001500  
 NY001510  
 NY001520  
 NY001530  
 NY001540  
 NY001550  
 NY001560  
 NY001570  
 NY001580  
 NY001590  
 NY001600  
 NY001610  
 NY001620  
 NY001630  
 NY001640  
 NY001650  
 NY001660  
 NY001670  
 NY001680  
 NY001690  
 NY001700  
 NY001710  
 NY001720  
 NY001730  
 NY001740  
 NY001750  
 NY001760  
 NY001770  
 NY001780  
 NY001790  
 NY001800  
 NY001810  
 NY001820  
 NY001830  
 NY001840  
 NY001850  
 NY001860  
 NY001870  
 NY001880  
 NY001890  
 NY001900  
 NY001910  
 NY001920

```

IT(I1) = J
XL(I1) = X
GO TO 195

CHANGE COMPONENT

I1=I+I2
IT(I1)=J
XL(I1)=X
GO TO 810

INSERT COMPONENT

I1 = IC
180 IF (I1-(I+I2)) 186,186,184
182 XL(I1+1) = XL(I1)
184 IT(I1+1) = IT(I1)
I1 = I1-1
GO TO 182
186 XL(I1+1) = X
IT(I1+1) = J
IC = IC+1
I2 = I2+1
GO TO 810

DELETE COMPONENT

I3 = I+I2
I4 = IC-1
DO 192 I1 = I3,I4
XL(I1) = XL(I1+1)
IT(I1) = IT(I1+1)
192 CONTINUE
IC = IC-1
I2 = I2-1
WRITE (LO1,811)
FORMAT ('ONEXT OR 04')
810 GO TO 153
195 WRITE (LO1,4)
GO TO 153
197 IF (LL-1) 92,340,340
92 WRITE (LO1,10)
10 FORMAT ('OLINE MEASURED AT 200(D)SQR?: 1-YES, 0-NO')
609 READ (LI,609) I
FORMAT (I2)
IF (I-1) 95,94,94
94 ON200 = 1.0
D200 = 1.0
  
```



```

95 GO TJ 96
QN200 = 1.042
D200 = 1.032
IF (LL-1) 98,340,340
96 WRITE (LO1,11)
98 FORMAT ('OENTER ANCHOR WT.(+LBS), AREA((M)SQR)')
11 READ (LI,610) ANCR, ANCR
610 FORMAT (F7.2, F5.2)
IF (LL-1) 215,340,340
215 WRITE (LO1,19)
19 FORMAT ('OENTER COMMENTS - 1 LINE MAX.')
C
C
C SET 'COMMENT' ARRAY TO BLANKS
DO 217 K = 1,36
IHDG(K) = IBLANK
217 CONTINUE
35 READ (LI,35) (IHDG(K),K = 1,36)
35 FORMAT (36(A2))
IF (LL-1) 100,340,340
100 WRITE (LO1,41)
41 FORMAT ('OENTER: DEPTH OF WATER (METERS)')
611 READ (LI,611) DPT
105 FORMAT (F6.1)
110 CONTINUE
IV = 1
WRITE (LO1,12)
12 FORMAT ('OINPUT CURRENT PROFILE--DEPTH(METERS), SPEED(CM/SEC), DIRECTION(DEGR)')
112 READ (LI,612) DCP(IV), CP(IV), RCP(IV)
612 FORMAT (F6.1, IX, F4.1, IX, F5.1)
IV = IV + 1
IF (DCP(IV-1)-DPT) 112,115,115
115 IV = IV - 1
IF (LL-1) 120,340,340
120 WRITE (LO1,13)
113 FORMAT ('OCHANGE STANDARD CD?: 1-YES, 0-NO')
613 READ (LI,613) I
122 FORMAT (I2)
IF (I-1) 130,122,122
14 WRITE (LO1,14)
14 FORMAT ('OENTER: 1-WIRE,2-LINE,3-INSTR,4-BALLS,5-UNSPEC. THEN CD(
IN),CD(T),)
123 READ (LI,614) I,X,Y
614 FORMAT (I2, IX, F3.1, IX, F6.3)
IF (I-99) 124,127,127
124 CCN(I) = X
CDT(I) = Y

```

NY 001930  
 NY 001940  
 NY 001950  
 NY 001960  
 NY 001970  
 NY 001980  
 NY 001990  
 NY 002000  
 NY 002010  
 NY 002020  
 NY 002030  
 NY 002040  
 NY 002050  
 NY 002060  
 NY 002070  
 NY 002080  
 NY 002090  
 NY 002100  
 NY 002110  
 NY 002120  
 NY 002130  
 NY 002140  
 NY 002150  
 NY 002160  
 NY 002170  
 NY 002180  
 NY 002190  
 NY 002200  
 NY 002210  
 NY 002220  
 NY 002230  
 NY 002240  
 NY 002250  
 NY 002260  
 NY 002270  
 NY 002280  
 NY 002290  
 NY 002300  
 NY 002310  
 NY 002320  
 NY 002330  
 NY 002340  
 NY 002350  
 NY 002360  
 NY 002370  
 NY 002380  
 NY 002390  
 NY 002400





```

C      IF (I-4) 126,125,126
C      Y/PI FOR CDT OF SPHERES ALLOWS USE OF THE
C      EQUATION FOR THE TANGENTIAL DRAG OF CYLINDERS
C
125 CDT(I) = Y/PI
126 WRITE (LOI,4)
    GO TO 123
127 IF (LL-1) 130,340,340
130 WRITE (LOI,15)
15  FORMAT ('0SEGMENT LENGTH(METERS)?')
    READ (LI,615) SEG
615  FORMAT (F7.2)
    IF (LL-1) 200,340,340
200 WRITE (LOI,16)
16  FORMAT ('0AUTO LENGTH ADJUST?:-0 TO 10 COMPCNTS')
    READ (LI,616) IDZ
616  FORMAT (I2)
    IF (IDZ.EQ.0) GO TO 206
202 WRITE (LOI,17)
17  FORMAT ('0ENTER: CRITICAL COMP., DESIRED DEPTH, ADJUST. COMP.')
    DO 204 I = 1, IDZ
617  READ (LI,617) IDA(I), DB(I), IDM(I)
    FORMAT (I2,I,X,F6.1,I,X,I2)
    CONTINUE
204 IF (LL-1) 210,340,340
206 WRITE (LOI,18)
210 FORMAT ('0ENTER MAGNITUDE, INCLINATION, AZIMUTH OF P.F.')
    READ (LI,618) TTENS, TCCLIN, TROTN
618  FORMAT (3F5.1)
    IF (LL-1) 211,340,340
211 WRITE (LOI,42)
42  FORMAT ('0CHANGE IN TERMINATION CONSTANTS?: 1=YES, 0=NO')
619  READ (LI,619) I
    FORMAT (I2)
    IF (I-1) 213,212,212
212 WRITE (LOI,43)
43  FORMAT ('0ENTER TERM. LENGTH(METER), WT.(LBS)')
    READ (LI,620) TERML, TERMW
620  FORMAT (F6.3,I,X,F6.3)
    SET LL=2: INDICATES INITIALIZATION COMPLETE
213 LL=1
    GO TO 340
C      START NEW COMPUTATION
C      PESET CRITICAL VARIABLES TO 0.0
C
220 CONTINUE

```







```

253 IF (ISSW(I)) 254,269,254
254 WRITE (LO,20) ISLEW
      WRITE (LO,35) (IH DG(K),K = 1,36)
      WRITE (LO,33)
33  FORMAT (//28X,18HSEGMENT STATISTICS )
      WRITE (LO,34)
34  FORMAT (//, COMP TYPE LENGTH INCL XEXC. YEXC. C.SPD C.DIR M.AZI UD
      1RAG VDRAG TDRAG,')
C
C
C
      SUBROUTINE CALC PERFORMS ALL BASIC CALCULATIONS
C
255 CALL CALC
      IF (ISSW(14)) 320,257,320
C
C
C
      CHECK DIFFERENCE BETWEEN ASSUMED DEPTH(RC) AND THE
      CALCULATED DEPTH OF THE TOP COMPONENT. LESS THAN 2.0 M.?
      IF YES - CHECK AUTO ADJUST STATUS,
      IF NO - CALC NEW FUDG, SET NEW RC.
      (X-RC)*0.7 TO HASTEN CONVERGENCE IN LARGE CURRENT SHEER
C
257 IF (DONE-1.0) 258,269,269
258 X = DPT-T4(IC)
      IF (ABS(RC-X)-2.0) 260,259,259
259 FUDG = FUDG+(X-RC)*0.7
      RC = (DPT-TL)+FUDG
      GO TO 255
C
C
C
      AUTO ADJUST EVALUATION - CHECK SPECIFIED INSTRUMENT
      DEPTHS, IF AT DESIRED DEPTHS - SET (DONE=2.0) AND
      GO TO OUTPUT, IF DEPTHS ARE INCORRECT - ADJUST
      LENGTH OF (IDM) AND GO TO (220)
C
260 IF (IDZ-1) 266,261,261
261 DO 264 K = 1,IDZ
C
C
C
      CHECK AND ADJUST LOWEST ELEMENTS FIRST
C
      J = (IDZ-K)+1
      I = IDA(J)
      Z = ((DPT-T4(IC))+T4(I))-DB(J)
      IF (ABS(Z)-1.0) 264,262,262
262 I = IDM(J)
C
C
C
      99% OF Z TO PARTIALLY ALLOW FOR STRETCH
C
      XL(I) = XL(I)+Z*.99
      GO TO 220
264 CONTINUE

```





```

266 DONE = 2.0
GO TO 245
C
C
C
IF SS-4 ON OUTPUT SUMMARY OF MOOR. STATS.
269 IF (ISSW(4)) 271,281,271
271 WRITE (LO,20) ISLEW
272 WRITE (LO,35) (IHDG(K),K = 1,36)
273 WRITE (LO,25)
25 FORMAT (22X,28HMOORING STATISTICS - SUMMARY
1//, COMP TYPE LENGTH WEIGHT DEPTH INCLIN TENSION EXCUR DRAG
2ACK-UP, )
IBKUP = 0
BKUP = QBKUP
DO 280 K = 1,IC
I1 = IT(K)
IF (IBKUP) 274,274,273
273 BKUP = 0.0
GO TO 276
274 BKUP = BKUP-(XL(K)*W(I1)+TERMW)
IF (IT(K)-35) 276,275,276
275 IBKUP = 2
276 QW = XL(K)*W(I1)
QX = (DPT-T4(IC))+T4(K)
QY = SQRT((T5(IC)-T5(K))**2+(T6(IC)-T6(K))**2)
WRITE (LO,26) K, IT(K), XL(K), QW, QX, T2(K),
IT3(K), QY, T1(K), BKUP
26 FORMAT (I4,3X,I2,2X,F6.1,1X,F7.1,2X,F6.1,1X,
1F5.1,3X,F6.1,1X,F6.1,1X,F5.1,1X,F7.1)
IF (ISSW(14)) 320,280,320
280 CONTINUE
C
C
C
IF SS-3 ON OUTPUT SUPPLEMENTAL STATISTICS
281 IF (ISSW(3)) 282,291,282
282 WRITE (LO,20) ISLEW
283 WRITE (LO,35) (IHDG(K),K = 1,36)
273 WRITE (LO,27)
27 FORMAT (/24X,23HSUPPLEMENTAL STATISTICS
1//391 COMP TYPE CD(N) AREA STR.LT PERC.STR
232H S.F. XEXCUR YEXCUR LAUNCH TENS )
DO 290 K = 1, IC
I1 = IT(K)
J1 = JCD(I1)
IF (IT(K)-25) 283,284,284
283 QR = RBS(I1)/T3(K-1)
284 QR = 0.0

```

NY003850  
 NY003860  
 NY003870  
 NY003880  
 NY003890  
 NY003900  
 NY003910  
 NY003920  
 NY003930  
 NY003940  
 NY003950  
 NY003960  
 NY003970  
 NY003980  
 NY003990  
 NY004000  
 NY004010  
 NY004020  
 NY004030  
 NY004040  
 NY004050  
 NY004060  
 NY004070  
 NY004080  
 NY004090  
 NY004100  
 NY004110  
 NY004120  
 NY004130  
 NY004140  
 NY004150  
 NY004160  
 NY004170  
 NY004180  
 NY004190  
 NY004200  
 NY004210  
 NY004220  
 NY004230  
 NY004240  
 NY004250  
 NY004260  
 NY004270  
 NY004280  
 NY004290  
 NY004300  
 NY004310  
 NY004320













NY 005290  
 NY 005300  
 NY 005310  
 NY 005320  
 NY 005330  
 NY 005340  
 NY 005350  
 NY 005360  
 NY 005370  
 NY 005380  
 NY 005390  
 NY 005400  
 NY 005410  
 NY 005420  
 NY 005430  
 NY 005440  
 NY 005450  
 NY 005460  
 NY 005470  
 NY 005480  
 NY 005490  
 NY 005500  
 NY 005510  
 NY 005520  
 NY 005530  
 NY 005540  
 NY 005550  
 NY 005560  
 NY 005570  
 NY 005580  
 NY 005590  
 NY 005600  
 NY 005610  
 NY 005620  
 NY 005630  
 NY 005640  
 NY 005650  
 NY 005660  
 NY 005670  
 NY 005680  
 NY 005690  
 NY 005700  
 NY 005710  
 NY 005720  
 NY 005730  
 NY 005740  
 NY 005750  
 NY 005760

```

PCLIN = CLIN
TENS = TTENS
ROTN = TROT N*RD
TXEX = 0.0
TYEX = 0.0
CPK = 1.076391E-01

PRIMARY LOOP
DO 570 I = 1,IC
  I1 = IT(I)
  DETERMINE DRAG COEFF. SUBSCRIPT (FUNCTION JCD)
  J1 = JCD(I1)
  DETERMINE NO. OF SEGMENTS IN COMPONENT (I)
  500 L = 1+IFIX(XL(I)/SEG)
  SECONDARY LOOP FOR SEGMENT CALC.
  505 DC 562 K = 1,L
  SET SEGMENT LENGTH (CLI) AND BUOYANCY (WT)
  CALC. CURRENT: - MEAN DEPTH (DN), SPEED(VN),
  DIRECTION(ROT), FOR EACH SEGMENT
  IF (K-L) 509,507,507
  507 CLI = (XL(I)-SEG*FLOAT(K-1))+TERML
  WT = W(I1)*(CLI-TERML)+TERMW
  GO TO 510
  509 CLI = SEG
  WT = W(I1)*SEG
  510 DN = TLD+(CLI/2.0)+FUDG*((DPT-(TLD+CLI/2.0))/TL)
  ASSIGN DEEPEST CURRENT VALUES IF DN EXCEEDS DPT
  IF (DN-DPT) 514,512,512
  512 VN = CP(IV)*CPK
  ROT = RCP(IV)*RD
  GO TO 519
  514 IF (DN-DCP(JC)) 518,516,516
  516 JC = JC+1
  GO TO 514
  518 VN = (CP(JC-1)+((DN-DCP(JC-1))/(DCP(JC)-DCP(JC-1))))
  1*(CP(JC)-CP(JC-1))*CPK
  ROT = (RCP(JC-1)+((DN-DCP(JC-1))/(DCP(JC)-DCP(JC-1))))

```





```

1*(RCP(JC)-RCP(JC-1)))#RD
SET ROTN=ROT WHEN TENSION = 0
519 IF (TENS) 522,522,524
522 ROTN = ROT
CALC. DRAG FORCES, TENSION, INCLINATION AND AZIMUTH OF SEG.
ITERATE UNTIL CHANGE LESS THAN 0.1 DEGREES FOR
INCLINATION - LOOP 526 TO 530
MOORING AZIMUTH - LOOP 526 TO 527
524 X = ROTN
525 Y = ROTN
526 Z = TENS*SIN(CLIN)
QCLIN = PCLIN
ANINC = VN*CCS(ROT-ROTN)
UTVN = VN*SIN(ROT-ROTN)
DRGT = A(I1)*PI*CLI*CDT(J1)*(SIN(QCLIN)*UTVN)*
1(ABS(SIN(QCLIN)*UTVN))
UDRGN = A(I1)*CLI*CDN(J1)*(COS(QCLIN)*UTVN)*
1(ABS(COS(QCLIN)*UTVN))
VDRGN = A(I1)*CLI*CDN(J1)*VTVN*ABS(VTVN)
WTT = WT*CCS(QCLIN)
WTN = WT*SIN(QCLIN)
QTENS = TENS*CCS(ANINC)+WTT+DRGT
(-WTN) CHANGE SIGN OF WTN FOR CORRECT SENSE
PCLIN = ATAN ((UDRGN-WTN-TENS*SIN(ANINC))/QTENS)+QCLIN
AVOID /0.0 IN STATEMENT 523
IF (QTENS*SIN(PCLIN)) 523,529,523
523 ROTN = ATAN((VDRGN-2*SIN(X-Y))/(QTENS*SIN(PCLIN)))+X
529 IF (ISSW(14)) 570,525,570
ASSURE +PCLIN
525 IF (PCLIN) 531,530,527
RATE OF CHANGE IN AZIMUTH < 0.1 DEGREE?
IF (ABS(ABS(X)-ABS(ROTN))-(0.1*RD)) 530,528,528
528 X = ROTN
GO TO 526

```

```

NY005770
NY005780
NY005790
NY005800
NY005810
NY005820
NY005830
NY005840
NY005850
NY005860
NY005870
NY005880
NY005890
NY005900
NY005910
NY005920
NY005930
NY005940
NY005950
NY005960
NY005970
NY005980
NY005990
NY006000
NY006010
NY006020
NY006030
NY006040
NY006050
NY006060
NY006070
NY006080
NY006090
NY006100
NY006110
NY006120
NY006130
NY006140
NY006150
NY006160
NY006170
NY006180
NY006190
NY006200
NY006210
NY006220
NY006230
NY006240

```





```

C      RATE OF CHANGE IN INCLINATION < 0.1 DEGREE?
C
530    IF (ABS(QCLIN-PCLIN)-(0.1*RD)) 532,526,526
C
C      HANDLES 180 DEG. ROTATION OF AZIMUTH WHEN VDRGN = 0.0
C
531    FCLIN = -PCLIN
      RCTN = ROTN+PI
      GO TO 524
532    TENS = COS(PCLIN-CLIN)*TENS+WT*COS(PCLIN)+DRGT
      CLIN = PCLIN
C
C      STRETCH CALC.: - J2 = 1 FOR WIRE (538)
C                      2 FOR DACRON (544)
C                      3 FOR NYLON (544)
C                      4 FOR UNSPECIFIED (544)
C
      STRL = CL I
      ELT = 0.0
534    IF (I1-21) 536,554,554
536    J2 = ((I1-1)/5)+1
      IF (J2-2) 538,540,540
538    ELT = ((TENS/RBS(I1))*E(1))+((TENS/(E(2)*AW(I1))))
      GO TO 546
C
C      INTERPOLATION BETWEEN TW(I)&TW(I-1) FOR LAUNCH TRANSIENTS
C      FOR PERMANENT ELONGATION CALC.
C      ELT = PERMANENT ELONGATION + ELASTIC ELONGATION
C
540    X = TW(I-1)+((TW(I)-TW(I-1))/XL(I))*(SEG*FLOAT(K-1)+CL I))
542    IF (X-TENS) 542,544,544
544    X = TENS
      N = (J2-1)*4+1
      ELT = (((X/(E(N)*A(I1)**2))*E(N+1))+((TENS/
1(E(N+2)*A(I1)**2))*E(N+3)))/100.0
C
C      CALC. SEGMENT STRETCHED LENGTH(STRL)
C      LENGTH * PERCENT STRETCH * 200(D)SCR MEAS. K
C
546    GO TO (548,550,552,548), J2
548    STRL = CL I*(1.0+ELT)
      GO TO 554
550    STRL = CL I*(1.0+ELT)*D200
      GO TO 554
552    STRL = CL I*(1.0+ELT)*QN200
C
C      CALC. HORIZ. EXCUR. (XEX&YEX) AND VERT. HEIGHT(VHT) OF SEG.
C

```

```

NY006250
NY006260
NY006270
NY006280
NY006290
NY006300
NY006310
NY006320
NY006330
NY006340
NY006350
NY006360
NY006370
NY006380
NY006390
NY006400
NY006410
NY006420
NY006430
NY006440
NY006450
NY006460
NY006470
NY006480
NY006490
NY006500
NY006510
NY006520
NY006530
NY006540
NY006550
NY006560
NY006570
NY006580
NY006590
NY006600
NY006610
NY006620
NY006630
NY006640
NY006650
NY006660
NY006670
NY006680
NY006690
NY006700
NY006710
NY006720

```



```

554 XEX = (STRL*SIN(CLIN))*COS(ROTN)
    YEX = (STRL*SIN(CLIN))*SIN(ROTN)
    VHT = STRL*COS(CLIN)
C
C
C
    OUTPUT SEGMENT STATISTICS
Z = CLIN/RD
557 IF (DONE-1.0) 560,558,558
558 X = ROTN/RD
    YN = VN/CPK
    WRITE (LO,32) I,IL,CLI,Z,XEX,YEX,VN,X,Y,UDRGN,VDRGN,DRGT
32 FORMAT (I3,3X,I2,3X,F5.1,1X,F4.1,2(1X,F5.1),1X,
13(1X,F5.1),2(1X,F5.2),1X,F8.5)
C
C
C
    SUM SEGMENT STATS. WITH COMPONENT TOTAL
560 DRAG = DRAG+UDRGN
    TVHT = TVHT+VHT
    TXEX = TXEX+XEX
    TYEX = TYEX+YEX
    TSTR = TSTR+(STRL-CLI)
    TLD = TLD+CLI
562 CONTINUE
C
C
C
    TRANSFER COMPONENT STATS. INTO T ARRAYS
T1(I) = DRAG
T2(I) = Z
T3(I) = TENS
T4(I) = TVHT
T5(I) = TXEX
T6(I) = TYEX
T7(I) = TSTR
TSTR = 0.0
570 CONTINUE
    RETURN
END
SUBROUTINE CONST - FOR USE WITH PRG. NOYFB, REV. 9.1
SUBROUTINE CONST WHOI BUOY COMPONENT CHARACTERISTICS
INPUT STANDARD AND STRETCH CHARACTERISTICS AND DRAG COEFFICIENTS
C
C
C
COMMON W(42),A(42),RBS(24),AW(5),E(16)
COMMON IT(65),XL(65),TW(66),CDN(5),CDT(5)
COMMON DCP(20),CP(20),RCP(20),IHDG(36)
COMMON T1(65),T2(65),T3(65),T4(65),T5(65),T6(65),T7(65)
COMMON IDA(10),IDM(10),MDB(10)

```



COMMON DPT, CN200, D200, FPM, L11, LO, LO3, NEW, TERMW, TERML  
COMMON RD, PI, TL, IC, IV, FUDG, RC, SEG, DONE  
COMMON TCLIN, TTENS, TROTN, QBKUP, TVELO, ANCR, ANCR

TERMW = -2.19  
TERML = 0.203  
RD = 0.017453293  
PI = 3.141592654  
FPM = 3.28084

A(I) - AREA OF COMP. IN SOR. METER PER METER LENGTH  
(LINE, WIRE, CHAIN = DIAMETER IN METERS)  
W(I) - WEIGHT OF COMP. IN POUNDS PER METER LENGTH  
(+= POSITIVE BUOYANCY, -= NEGATIVE BUOYANCY)  
RBS(I) - RATED BREAKING STRENGTH IN POUNDS  
AW(I) - CROSS SECTIONAL METAL AREA OF WIRE IN SOR. INCHES

WIRE CONSTANTS: (1)-3/16', (2)-1/4', (3)-5/16', (4)-3/8', (5)-?  
U.S.S., TORQUE BALANCED JACKETED 3X19 WIRE

A(1) = 6.5786E-03  
W(1) = -0.154  
RBS(1) = 4000.0  
AW(1) = 0.01611  
A(2) = 8.3566E-03  
W(2) = -0.266  
RBS(2) = 6750.0  
AW(2) = 0.02738  
A(3) = 9.9568E-03  
W(3) = -0.41  
RBS(3) = 10300.0  
AW(3) = 0.04206  
A(4) = 1.15824E-02  
W(4) = -0.594  
RBS(4) = 14800.0  
AW(4) = 0.06015  
A(5) = 0.0  
W(5) = 0.0  
RBS(5) = 1.0  
AW(5) = 1.0

SAMPSON, SINGLE BRAID DACRON, WHOI SPECS.  
DACRON LINE CONSTANTS: (6)-3/8', (7)-7/16', (8)-1/2',  
(9)-9/16', (10)-5/8',  
DIAMETER IS 93.5% OF NOMINAL SIZE

NY007210  
NY007220  
NY007230  
NY007240  
NY007250  
NY007260  
NY007270  
NY007280  
NY007290  
NY007300  
NY007310  
NY007320  
NY007330  
NY007340  
NY007350  
NY007360  
NY007370  
NY007380  
NY007390  
NY007400  
NY007410  
NY007420  
NY007430  
NY007440  
NY007450  
NY007460  
NY007470  
NY007480  
NY007490  
NY007500  
NY007510  
NY007520  
NY007530  
NY007540  
NY007550  
NY007560  
NY007570  
NY007580  
NY007590  
NY007600  
NY007610  
NY007620  
NY007630  
NY007640  
NY007650  
NY007660  
NY007670  
NY007680



A(6) = 8.90588E-03  
 W(6) = -0.0375  
 RBS(6) = 5700.0  
 A(7) = 1.03902E-02  
 W(7) = -0.0516  
 RBS(7) = 7000.0  
 A(8) = 1.18745E-02  
 W(8) = -0.0667  
 RBS(8) = 9000.0  
 A(9) = 1.33588E-02  
 W(9) = -0.0851  
 RBS(9) = 11200.0  
 A(10) = 1.48431E-02  
 W(10) = -0.1082  
 RBS(10) = 14000.0

C  
C  
C  
C  
C  
C

COLUMBIA, SINGLE BRAID PLAITED NYLON, WHOI SPECS  
 NYLON LINE CONSTANTS: (11)-3/8', (12)-1/2', (13)-9/16',  
 (14)-5/8', (15)-3/4',  
 DIAMETER IS 93.5% OF NOMINAL SIZE

A(11) = 8.90588E-03  
 W(11) = -0.011  
 RBS(11) = 3700.0  
 A(12) = 1.18745E-02  
 W(12) = -0.0204  
 RBS(12) = 6400.0  
 A(13) = 1.33588E-02  
 W(13) = -0.0261  
 RBS(13) = 8200.0  
 A(14) = 1.48431E-02  
 W(14) = -0.033  
 RBS(14) = 10400.0  
 A(15) = 1.78118E-02  
 W(15) = -0.0457  
 RBS(15) = 14200.0

C  
C  
C

UNSPECIFIED SYNTHETIC LINE : (16)-? THROUGH (20)-?

A(16) = 0.0  
 W(16) = 0.0  
 RBS(16) = 1.0  
 A(17) = 0.0  
 W(17) = 0.0  
 RBS(17) = 1.0  
 A(18) = 0.0  
 W(18) = 0.0  
 RBS(18) = 1.0

NY 007690  
 NY 007700  
 NY 007710  
 NY 007720  
 NY 007730  
 NY 007740  
 NY 007750  
 NY 007760  
 NY 007770  
 NY 007780  
 NY 007790  
 NY 007800  
 NY 007810  
 NY 007820  
 NY 007830  
 NY 007840  
 NY 007850  
 NY 007860  
 NY 007870  
 NY 007880  
 NY 007890  
 NY 007900  
 NY 007910  
 NY 007920  
 NY 007930  
 NY 007940  
 NY 007950  
 NY 007960  
 NY 007970  
 NY 007980  
 NY 007990  
 NY 008000  
 NY 008010  
 NY 008020  
 NY 008030  
 NY 008040  
 NY 008050  
 NY 008060  
 NY 008070  
 NY 008080  
 NY 008090  
 NY 008100  
 NY 008110  
 NY 008120  
 NY 008130  
 NY 008140  
 NY 008150  
 NY 008160







A(19) = 0.0  
W(19) = 0.0  
RES(19) = 1.0  
A(20) = 0.0  
W(20) = 0.0  
RES(20) = 1.0

CHAIN: (21)-1/4', (22)-3/8', (23)-1/2', (24)-3/4'

A(21) = 2.74E-02  
W(21) = -2.33  
RBS(21) = 5400.0  
A(22) = 3.7E-02  
W(22) = -5.12  
RBS(22) = 12150.0  
A(23) = 4.8E-02  
W(23) = -9.02  
RBS(23) = 21600.0  
A(24) = 6.85E-02  
W(24) = -19.52  
RBS(24) = 48.6E+03

CYLINDRICAL INSTRUMENTS:  
DIVISOR IS THE LENGTH OF THE INSTRUMENT (METERS)  
(25)-VACM, (26)-850(LT), (27)-850(HEAVY),  
(28)-ENG.CM, (29)-INCLINOMETER, (30)-DEPTH REC.,  
(31)-TENSION REC., (32)-TENSAC  
(33)-?, (34)-?, (35)-RELEASE, AMF TRANSPONDING

A(25) = 0.1579  
W(25) = -75.0/1.9  
A(26) = 0.17917  
W(26) = -40.0/1.8  
A(27) = 0.17917  
W(27) = -50.0/1.8  
A(28) = 0.16043  
W(28) = -40.0/0.8  
A(29) = 0.16043  
W(29) = -40.0/0.8  
A(30) = 0.16043  
W(30) = -40.0/0.8  
A(31) = 0.16043  
W(31) = -40.0/0.8  
A(32) = 0.16875  
W(32) = -70.0/1.6  
A(33) = 0.0  
W(33) = 0.0  
A(34) = 0.0

NY008170  
NY008180  
NY008190  
NY008200  
NY008210  
NY008220  
NY008230  
NY008240  
NY008250  
NY008260  
NY008270  
NY008280  
NY008290  
NY008300  
NY008310  
NY008320  
NY008330  
NY008340  
NY008350  
NY008360  
NY008370  
NY008380  
NY008390  
NY008400  
NY008410  
NY008420  
NY008430  
NY008440  
NY008450  
NY008460  
NY008470  
NY008480  
NY008490  
NY008500  
NY008510  
NY008520  
NY008530  
NY008540  
NY008550  
NY008560  
NY008570  
NY008580  
NY008590  
NY008600  
NY008610  
NY008620  
NY008630  
NY008640

CC

CCCCC



W(34) = 0.0	NY008650
A(35) = 0.1499	NY008660
W(35) = -80.0/1.8	NY008670
SPHERICAL INSTRUMENTS:	
DIVISOR IS THE LENGTH OF THE INSTRUMENT ( METERS)	
(36)-?, (37)-MIT P/T, (38)-RADIC FLOAT	NY008680
A(36) = 0.0	NY008690
W(36) = 0.0	NY008700
A(37) = 0.207	NY008710
W(37) = -18.0/0.4	NY008720
A(38) = 0.26	NY008730
W(38) = 41.0/1.0	NY008740
SPHERES MOUNTED ON 3/8" CHAIN, 1 METER COMPONENT LENGTH	
(39)-16" SPHERE - 17.5" O.D., (40)-17" SPHERE-18.5" O.D.	NY008750
A(39) = 0.25138	NY008760
W(39) = 43.5	NY008770
A(40) = 0.26962	NY008780
W(40) = 53.0	NY008790
UNDEFINED COMPONENTS W/UNIQUE DRAG COEFF.	
(41)-?, (42)-?	NY008800
A(41) = 0.0	NY008810
W(41) = 0.0	NY008820
A(42) = 0.0	NY008830
W(42) = 0.0	NY008840
DRAG COEFFICIENTS - CDN-NORMAL, CDT - TANGENTIAL	
(1)-WIRE,(2)-LINE & CHAIN,(3)-INSTRUMENTS,	NY008850
(4)-SPHERES,(5)-UNSPECIFIED	NY008860
CDN(1) = 1.3	NY008870
CDT(1) = 0.007	NY008880
CDN(2) = 1.3	NY008890
CDT(2) = 0.007	NY008900
CDN(3) = 1.2	NY008910
CDT(3) = 0.9	NY008920
CDN(4) = 0.5	NY008930
CDT(4) = 0.5/PI	NY008940
CDN(5) = 0.0	NY008950
CDT(5) = 0.0	NY008960
STRETCH CHARACTERISTICS	
	NY008970
	NY008980
	NY008990
	NY009000
	NY009010
	NY009020
	NY009030
	NY009040
	NY009050
	NY009060
	NY009070
	NY009080
	NY009090
	NY009100
	NY009110
	NY009120



C C (1-4)-W IRE, (5-8)-DACRON, (9-12)-NYLON, (13-16)-UNSPECIFIED

E(1) = 1.42E571E-03

E(2) = 20.5E+06

E(3) = 0.0

E(4) = 0.0

E(5) = 2.81E+06

E(6) = 0.607

E(7) = 3.83E+06

E(8) = 0.74

E(9) = 1.56E+05

E(10) = 0.516

E(11) = 1.3262E+05

E(12) = 0.535

E(13) = 0.0

E(14) = 0.0

E(15) = 0.0

E(16) = 0.0

CONTINUE

RETURN

END

SUBROUTINE TRANS

END

SUBROUTINE TRANS - FOR USE WITH PROG. NOYFB, REV. 9.1  
CALCULATE AND STORE IN ARRAY (TW), PEAK TENSION ON EACH  
COMPONENT EXPERIENCED DURING ANCHOR LAST LAUNCH  
INITIATES EACH RUN BY DETERMINING TOTAL RELAXED LENGTH (TL)  
AND RESERVE BUOYANCY AT COMPONENT NO.1 (QBKUP)  
CALC. TERMINAL VELOCITY OF FREE FALL ANCHOR (TVELO)

COMMON W(42), A(42), RBS(24), AW(5), E(16)

COMMON IT(65), XL(65), TW(66), CDN(5), CDT(5)

COMMON DCP(20), CP(20), RCP(20), IHOG(36)

COMMON T1(65), T2(65), T3(65), T4(65), T5(65), T6(65), T7(65)

COMMON IDA(10), IDM(10), DB(10)

COMMON OPT, QN200, D200, FPM, L11, LQ, LC3, NEW, TERMW, TERML

COMMON RD, P1, TL, IC, IV, FUDG, RC, SEG, DONE

COMMON TCLIN, TTENS, TROTN, QBKUP, TVELC, ANCR, ANCRA

X = 0.0

J = 0

DO 420 I = 1, IC

401 I1 = IT(I)

SUM INPUT LENGTHS (TL) AND COMP. BUOYANCIES (T1(I))

T1(I) IS USED FOR TEMPORARY STORAGE

TL = TL + XL(I) + TERML

X = X + TERMW + W(I1) \* XL(I)

NY009130  
NY009140  
NY009150  
NY009160  
NY009170  
NY009180  
NY009190  
NY009200  
NY009210  
NY009220  
NY009230  
NY009240  
NY009250  
NY009260  
NY009270  
NY009280  
NY009290  
NY009300  
NY009310  
NY009320  
NY009330  
NY009340  
NY009350  
NY009360  
NY009370  
NY009380  
NY009390  
NY009400  
NY009410  
NY009420  
NY009430  
NY009440  
NY009450  
NY009460  
NY009470  
NY009480  
NY009490  
NY009500  
NY009510  
NY009520  
NY009530  
NY009540  
NY009550  
NY009560  
NY009570  
NY009580  
NY009590  
NY009600



```

C      T1(I) = X
C      RESERVE BUOYANCY (QBKUP) AT TOP COMP. = SUM OF WEIGHTS
C      OF RELEASE (35) AND ALL COMPONENTS ABOVE IT
C
C      IF(J-1) 402,405,405
C      402 QBKUP = X
C      IF(I1-35) 405,403,405
C      403 J = 2
C
C      DETERMINE DRAG COEFF. SUBSCRIPT (J1), FUNCT. JCD
C      405 J1 = JCD(I1)
C
C      CALC. TOTAL AREA*CD FOR EACH COMPONENT AND SUM
C      GO TO 410 - WIRE, LINE AND CHAIN
C      412 - INSTRUMENTS (CYLINDERS)
C      410 - SPHERES
C      412 - UNSPECIFIED
C      Y = CDT*SURFACE AREA
C
C      Y = CDT (J1)*PI*A(I1)*XL(I1)*FPM**2
C      GO TO (410,410,412,410,412), J1
C      TW(I+1) = TW(I)+Y
C      GO TO 420
C      410 TW(I+1) = TW(I)+(CDN(J1)*(PI/4.0)*A(I1)**2
C      412 1*FPM**2)+Y
C      420 CONTINUE
C
C      CALC. VELQ.**2(VSQ), CD*AREA OF ANCHOR APPLIED
C      THEN CALC. TERMINAL VELOCITY AND TRANSIENT PEAK LOAD
C
C      VSQ = (ANCR-T1(IC))/(TW(IC+1)+ANCR*A*FPM**2*1.15)
C      DO 425 I = 1,IC
C      416 TW(I) = T1(I)+TW(I+1)*VSQ
C      425 CONTINUE
C      TVELO = (SQRT(VSQ)/FPM)*60.0
C      RETURN
C      END
C      SUBROUTINE PTIO
C
C      SUBROUTINE PTIO - FOR USE WITH PRCG. NOYFB, REV. 9.1
C      INPUT/OUTPUT SUBROUTINE FOR PERMANENT RECORD OF MOORING
C      SPECIFICALLY FOR USE WITH A PAPER TAPE READER AND PAPER TAPE
C      PUNCH BUT USEABLE WITH OTHER I/O DEVICES
C      READS/Writes NUMBER, TYPE AND LENGTH OF MOORING COMPONENTS
C      READS/Writes CONSTANTS AND VARIABLES USED FOR MOORING CONFIG
C      READS/Writes OPERATOR COMMENTS

```

```

NY0009610
NY0009620
NY0009630
NY0009640
NY0009650
NY0009660
NY0009670
NY0009680
NY0009690
NY0009700
NY0009710
NY0009720
NY0009730
NY0009740
NY0009750
NY0009760
NY0009770
NY0009780
NY0009790
NY0009800
NY0009810
NY0009820
NY0009830
NY0009840
NY0009850
NY0009860
NY0009870
NY0009880
NY0009890
NY0009900
NY0009910
NY0009920
NY0009930
NY0009940
NY0009950
NY0009960
NY0009970
NY0009980
NY0009990
NY0100000
NY0100010
NY0100020
NY0100030
NY0100040
NY0100050
NY0100060
NY0100070
NY0100080

```





C

```

COMMON W(42), A(42), RBS(24), AW(5), E(16)
COMMON IT(65), XL(65), TW(66), CDN(5), CDT(5)
COMMON DCP(20), CP(20), RCP(20), IHDG(36)
COMMON T1(65), T2(65), T3(65), T4(65), T5(65), T6(65), T7(65)
COMMON IDA(10), IDM(10), DB(10)
COMMON DPT, QN200, D200, FPM, L11, LC, LC3, NEW, TERMW, TERML
COMMON RD, PI, TL, IC, IV, FUDG, RC, SEG, DONE
COMMON TCLIN, TTENS, TROTN, QBKUP, TVELO, ANCR, ANCR
36 FORMAT (4F13.7)
37 FORMAT (12, F13.7)
38 FORMAT (36(A2))
39 IF (NEW-15) 1010, 1006, 1000
1000 WRITE (L03, 39) DPT, QN200, D200, ANCR, ANCR, TERMW, TERML, IC
      WRITE (L03, 36) (W(I), A(I), I = 1, 42), (RBS(I), I = 1, 24),
1(AW(I), I = 1, 5), (E(I), I = 1, 16), (CDN(I), CDT(I), I = 1, 5)
      DO 1002 I = 1, IC
      WRITE (L03, 37) IT(I), XL(I)
1002 CCNTINUE
      WRITE (L03, 38) (IHDG(I), I = 1, 36)
      GO TO 1010
1006 READ (L11, 39) DPT, QN200, D200, ANCR, ANCR, TERMW, TERML, IC
      READ (L11, 36) (W(I), A(I), I = 1, 42), (RBS(I), I = 1, 24),
1(AW(I), I = 1, 5), (E(I), I = 1, 16), (CDN(I), CDT(I), I = 1, 5)
      DO 1008 I = 1, IC
      READ (L11, 37) IT(I), XL(I)
1008 CCNTINUE
1010 READ (L11, 38) (IHDG(I), I = 1, 36)
      CCNTINUE
      RETURN
      END
      FUNCTION JCD(I1)
      FUNCTION ION JCD - USE WITH PRG. NUYFB, REV 9.1
      SET SUBSCRIPT FOR COMPONENT DRAG COEFFICIENTS (J1)
      (5) - 41 AND 42
      (4) - 36 THRU 40
      (3) - 25 35
      (2) - 6 24
      (1) - 1 5
      JCD = 5
      IF (I1-41) 483, 487, 487
483 JCD = 4
      IF (I1-36) 484, 487, 487
484 JCD = 3
      IF (I1-25) 485, 487, 487

```

C C C C C C C C

NY01 0090  
 NY01 0100  
 NY01 0110  
 NY01 0120  
 NY01 0130  
 NY01 0140  
 NY01 0150  
 NY01 0160  
 NY01 0170  
 NY01 0180  
 NY01 0190  
 NY01 0200  
 NY01 0210  
 NY01 0220  
 NY01 0230  
 NY01 0240  
 NY01 0250  
 NY01 0260  
 NY01 0270  
 NY01 0280  
 NY01 0290  
 NY01 0300  
 NY01 0310  
 NY01 0320  
 NY01 0330  
 NY01 0340  
 NY01 0350  
 NY01 0360  
 NY01 0370  
 NY01 0380  
 NY01 0390  
 NY01 0400  
 NY01 0410  
 NY01 0420  
 NY01 0430  
 NY01 0440  
 NY01 0450  
 NY01 0460  
 NY01 0470  
 NY01 0480  
 NY01 0490  
 NY01 0500  
 NY01 0510  
 NY01 0520  
 NY01 0530  
 NY01 0540  
 NY01 0550  
 NY01 0560



```
485 JCD = 2      486,487,487
    IF (I1-6)
486 JCD = 1
487 CONTINUE
    RETURN
    END
```

```
NY010570
NY010580
NY010590
NY010600
NY010610
NY010620
```



# APPENDIX B

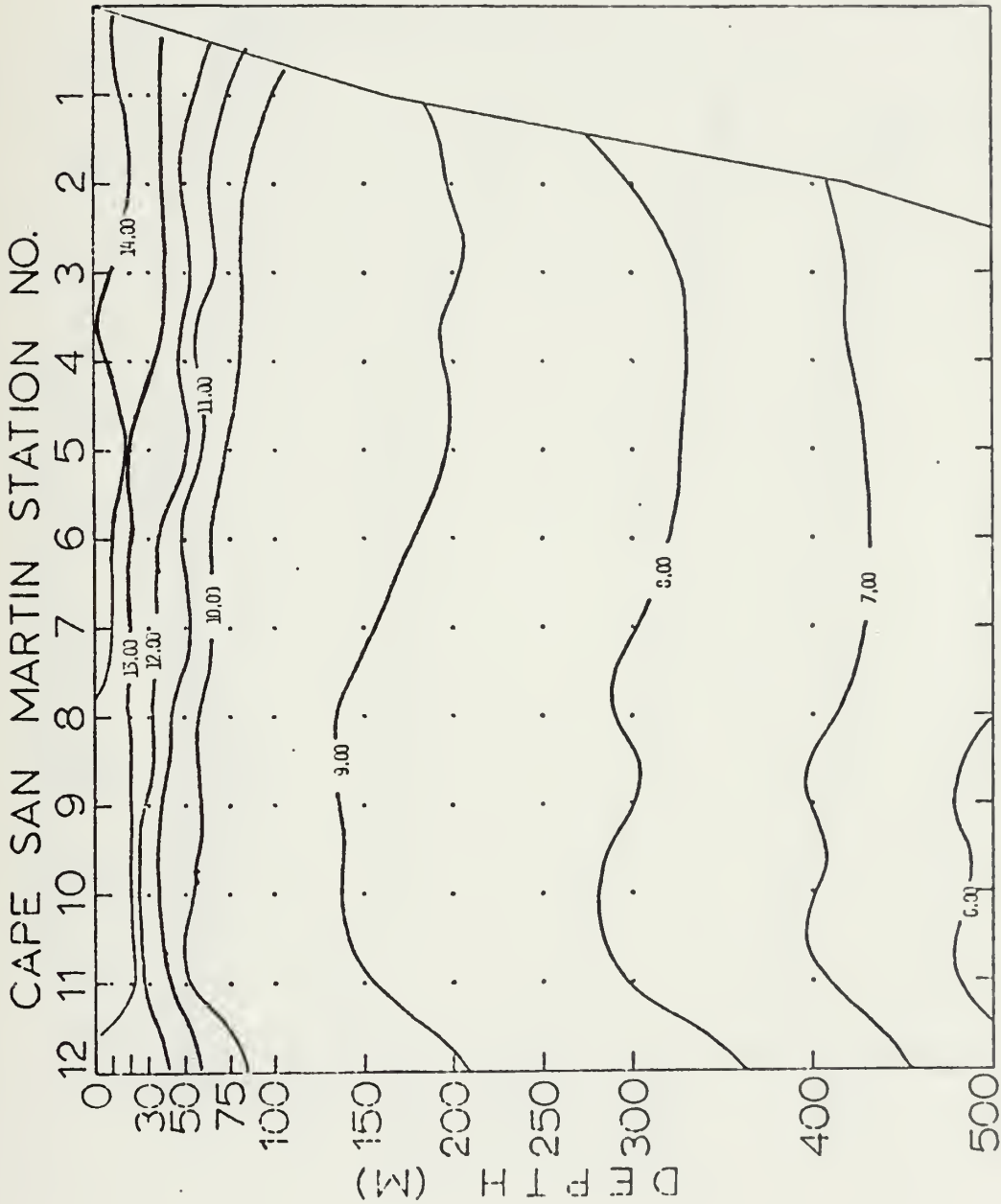


Figure 43. Temperature (°C) on a vertical section for the Cape San Martin line on 27-28 November 1978.



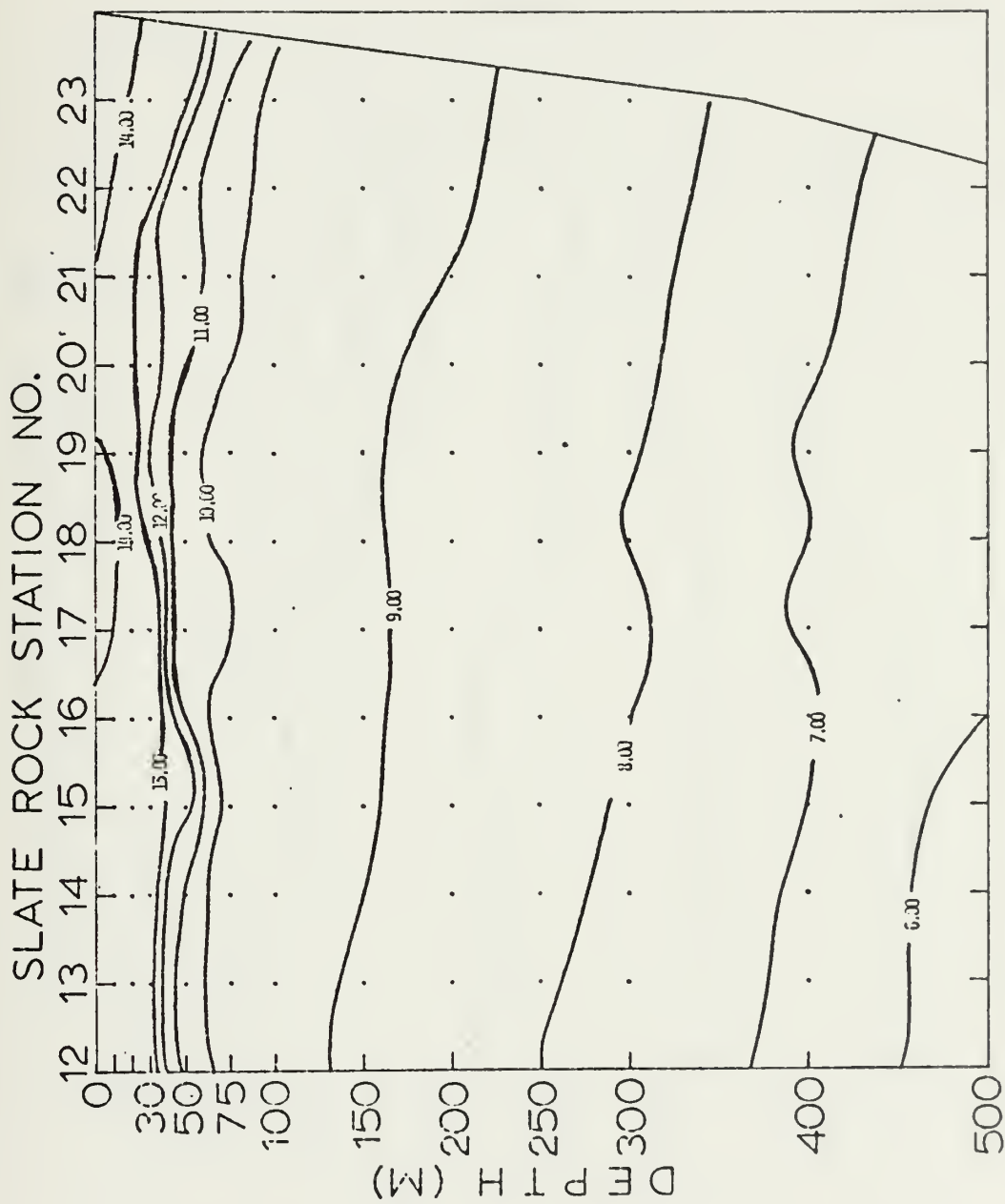


Figure 44. Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Slate Rock line on 27-28 November 1978.





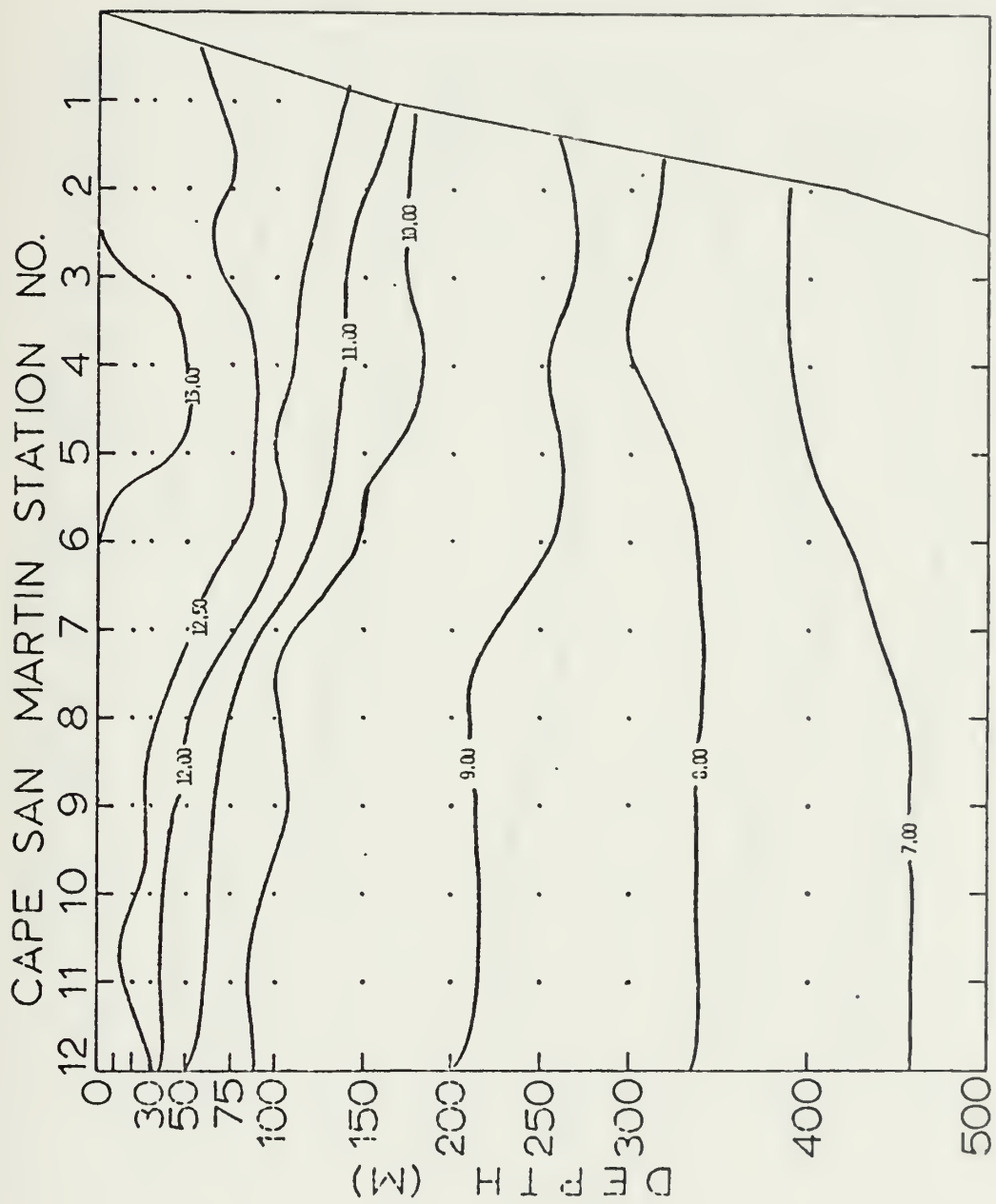


Figure 45. Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Cape San Martin line on 8-9 January 1979.



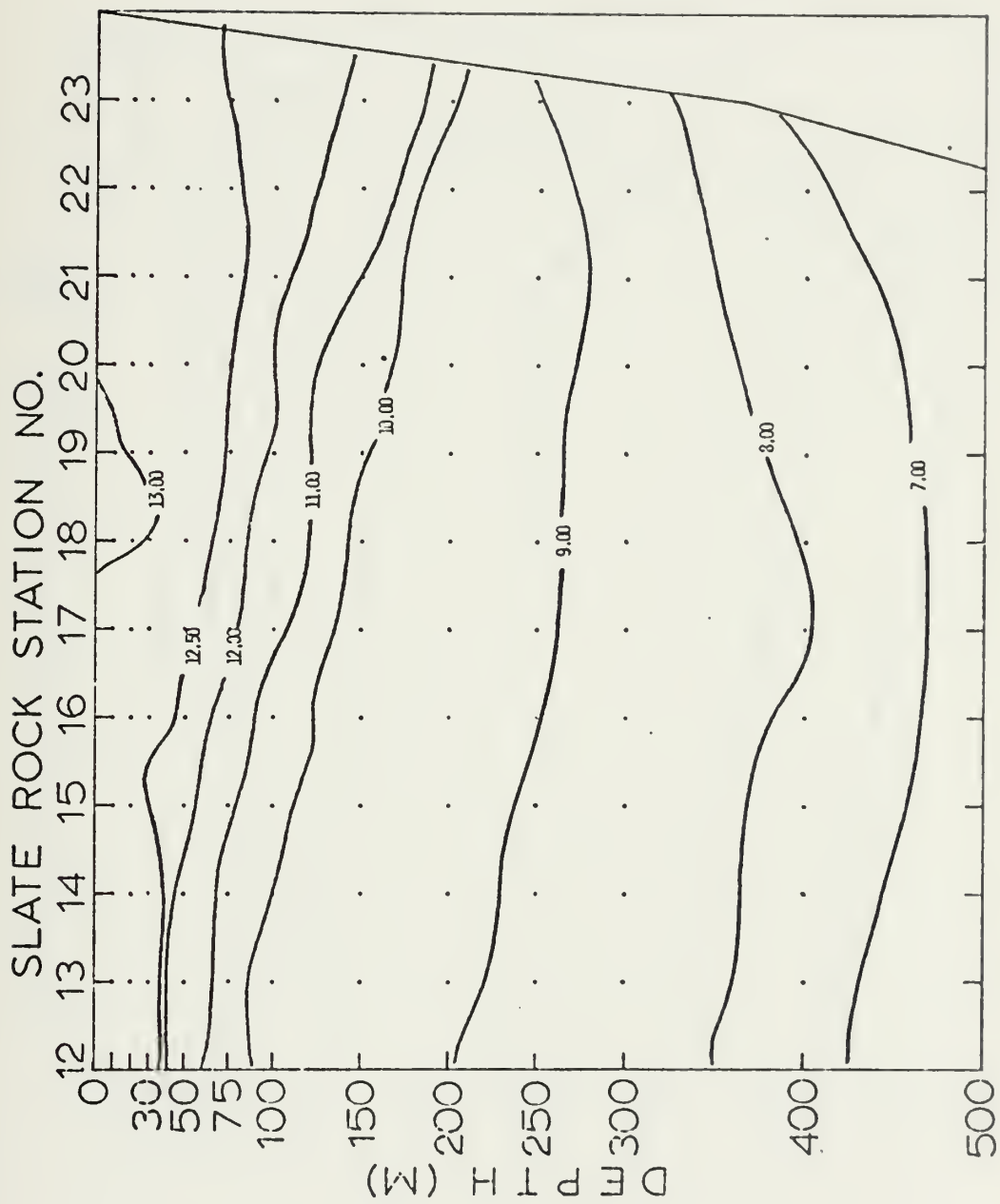


Figure 46. Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Slate Rock line on 8-9 January 1979.



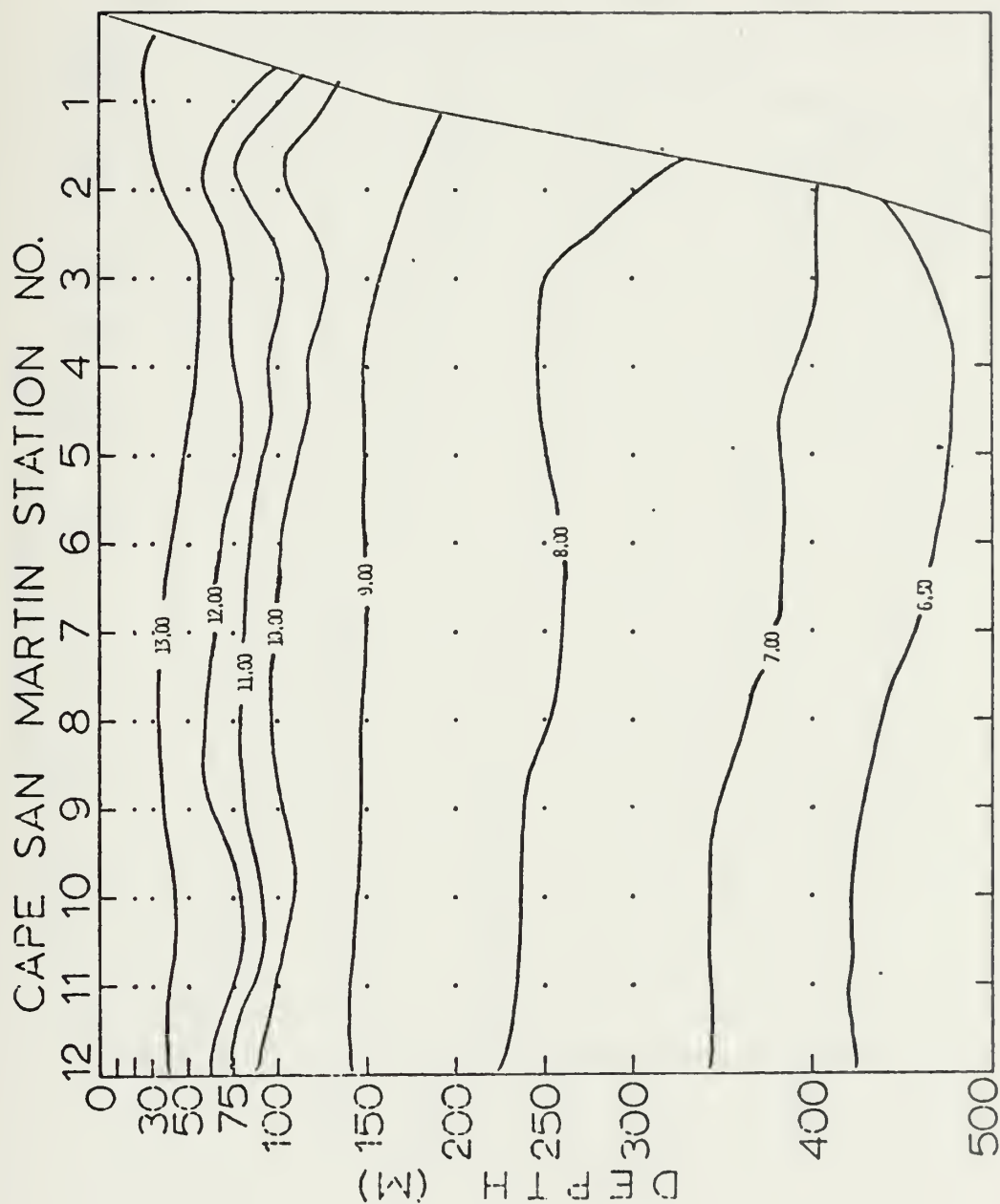


Figure 47. Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Cape San Martin line on 22-23 January 1979.



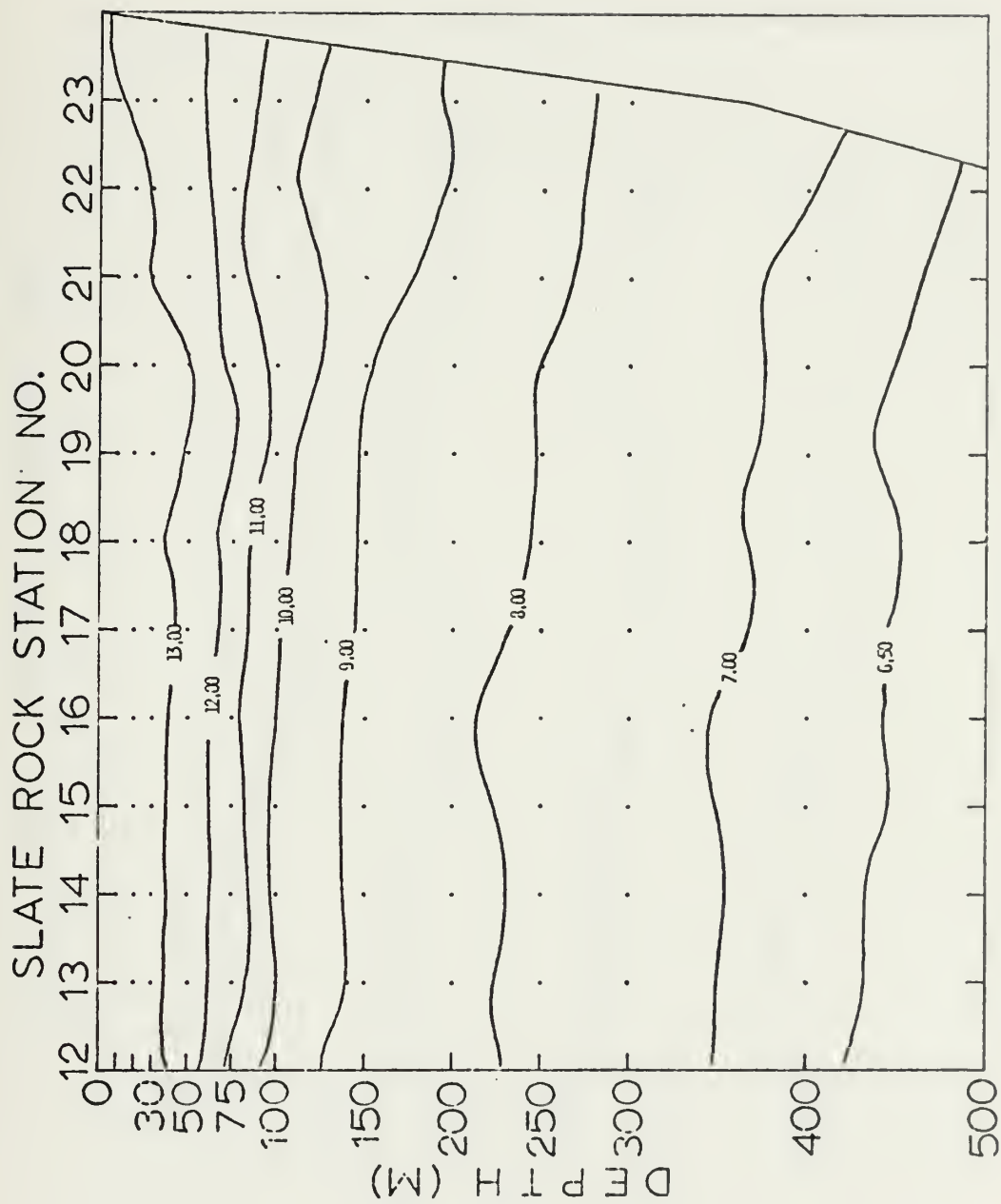


Figure 48. Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Slate Rock line on 22-23 January 1979.





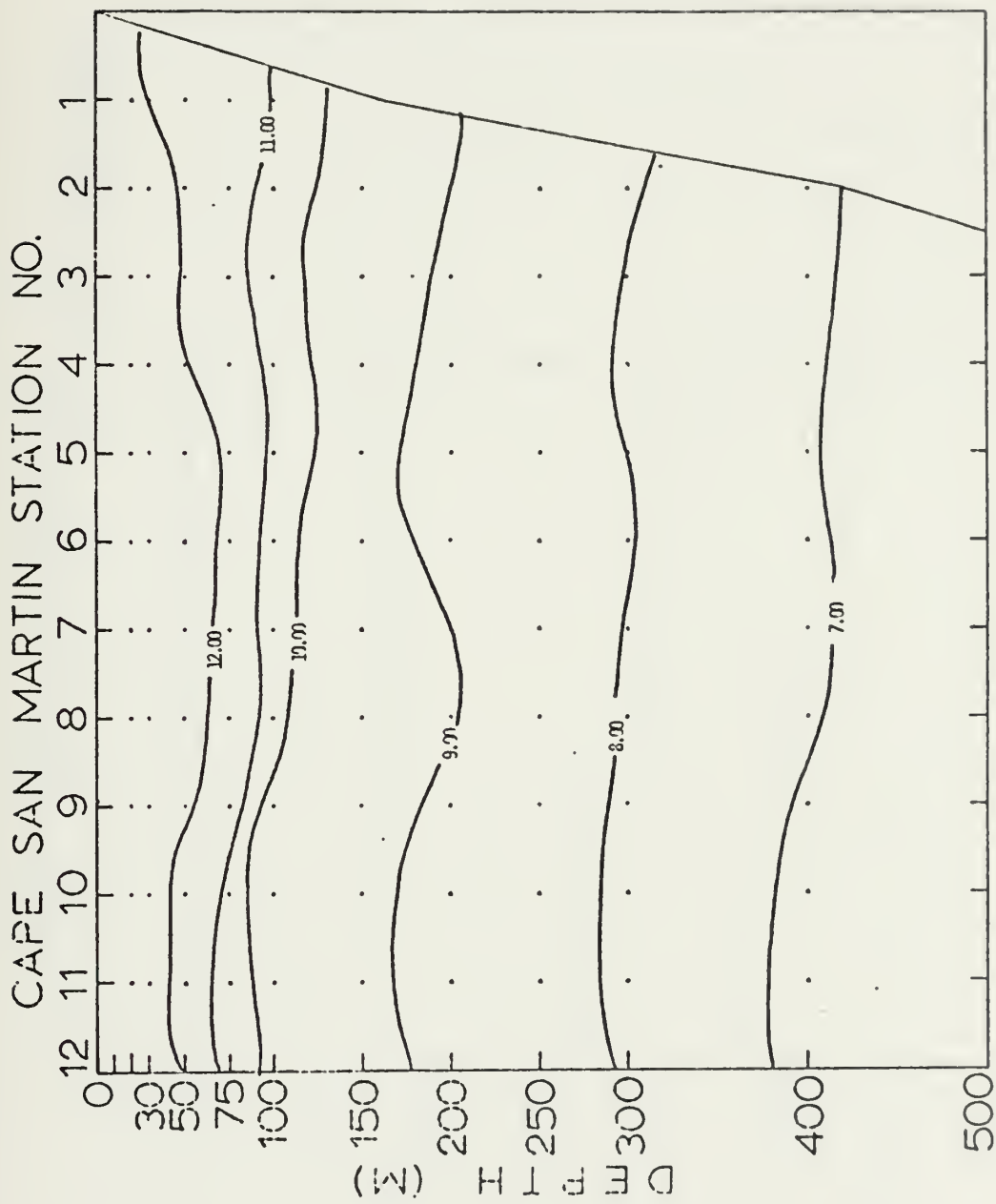


Figure 49. Temperature ( $^{\circ}\text{C}$ ) on a vertical section for the Cape San Martin line on 21-22 February 1979.



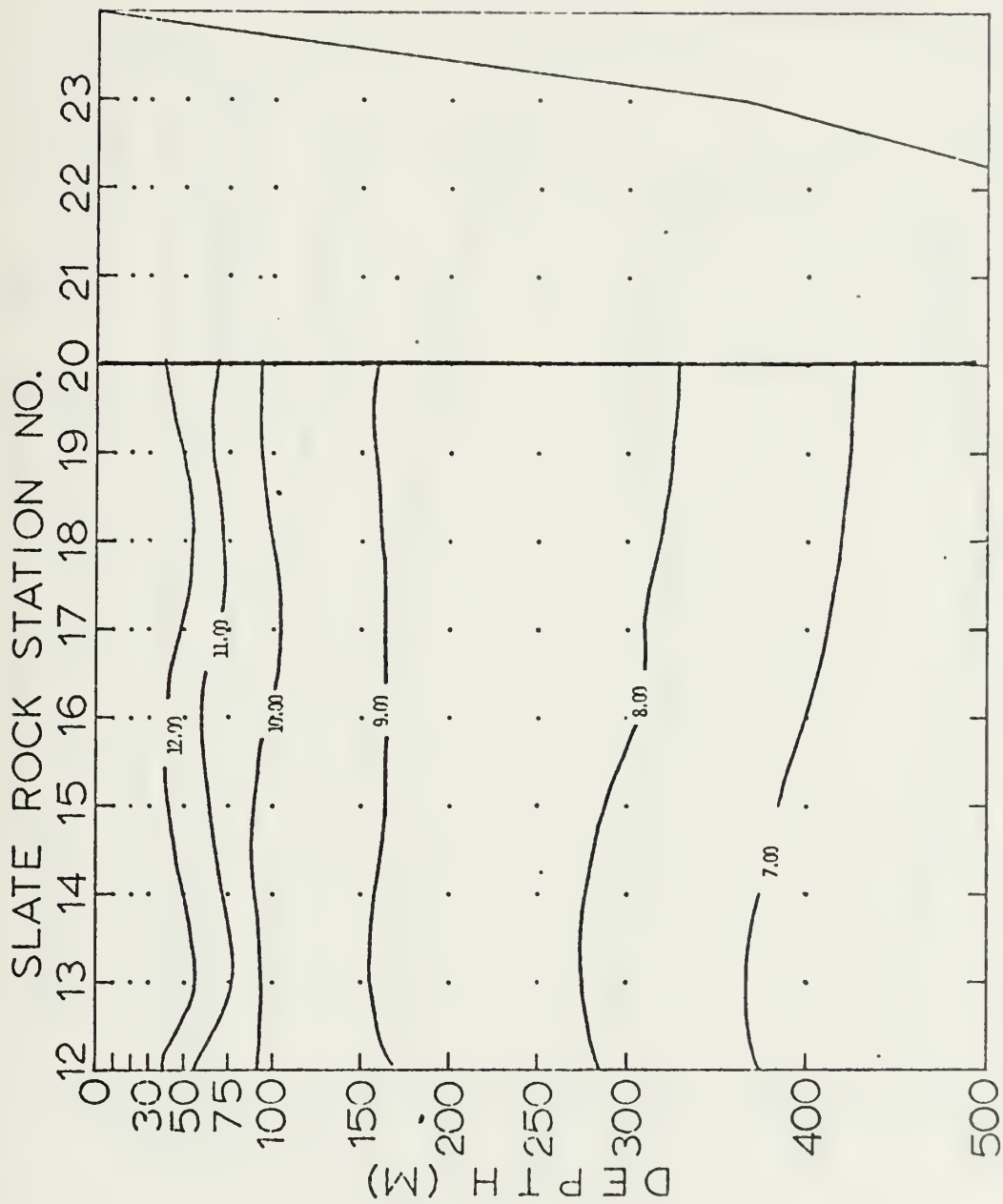


Figure 50. Temperature (°C) on a vertical section for the Slate Rock line on 21-22 February 1979.



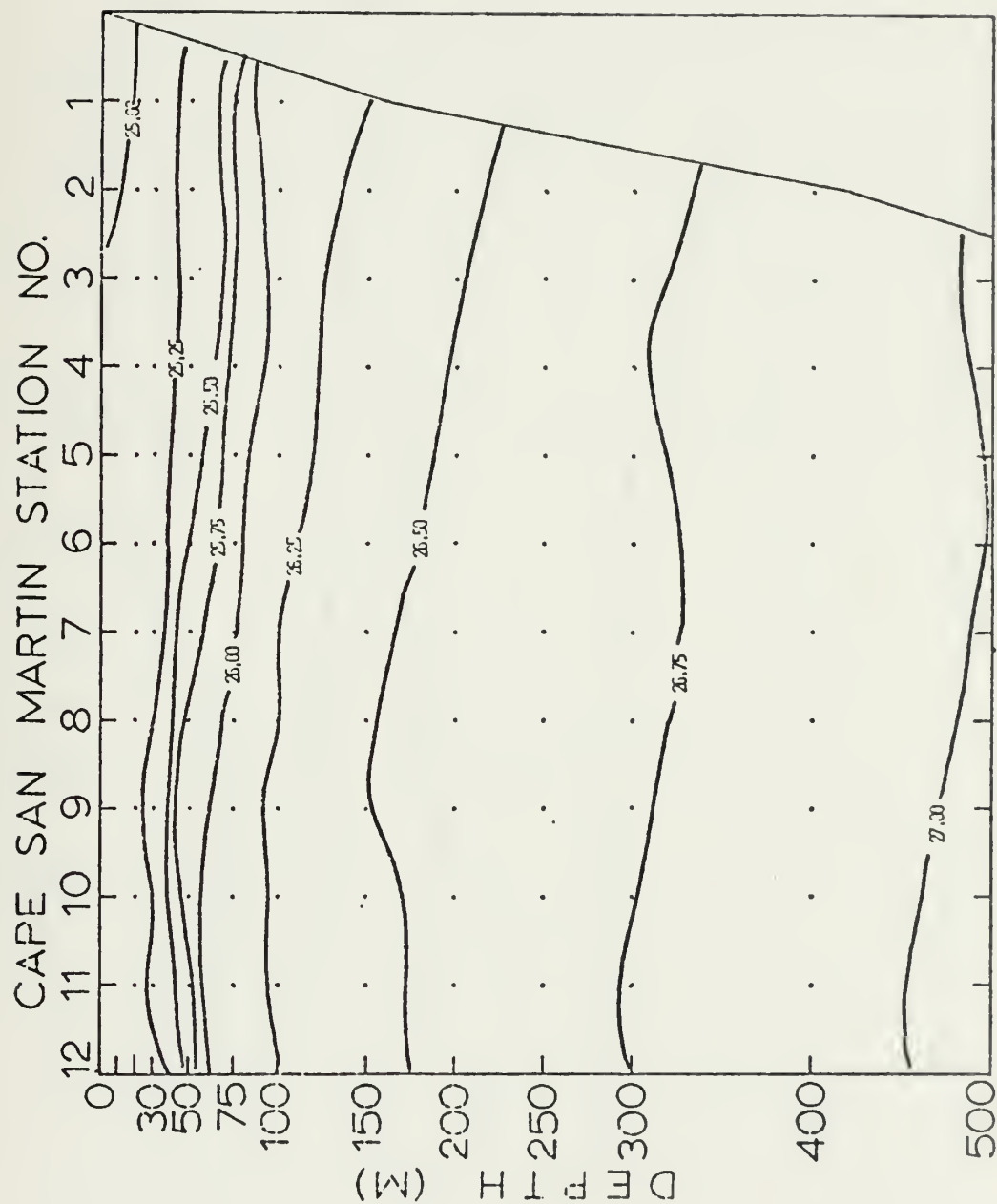


Figure 51. Sigma-t on a vertical section for the Cape San Martin line on 27-28 November 1978.



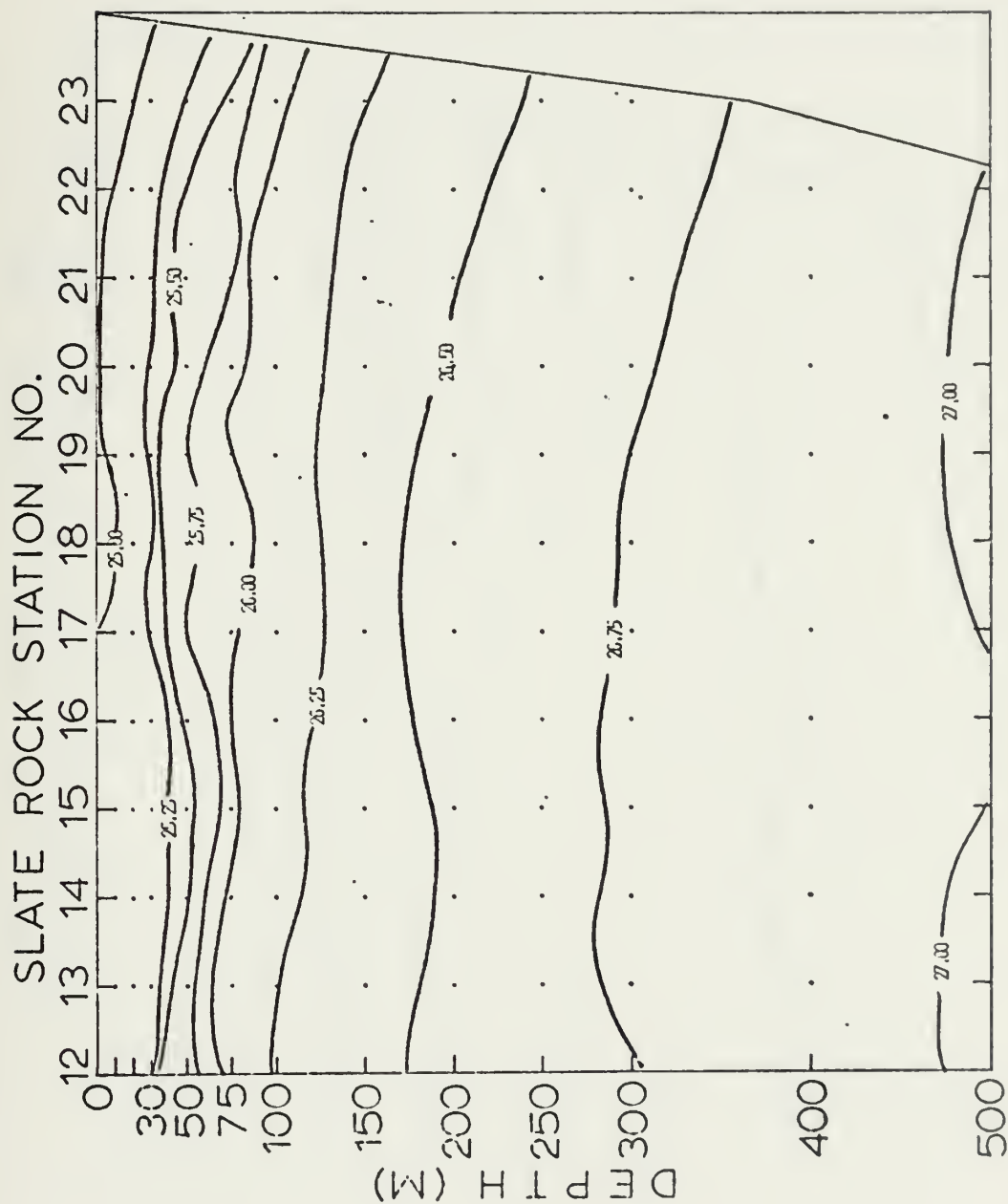


Figure 52, Sigma-t on a vertical section for the Slate Rock line on 27-28 November 1979.





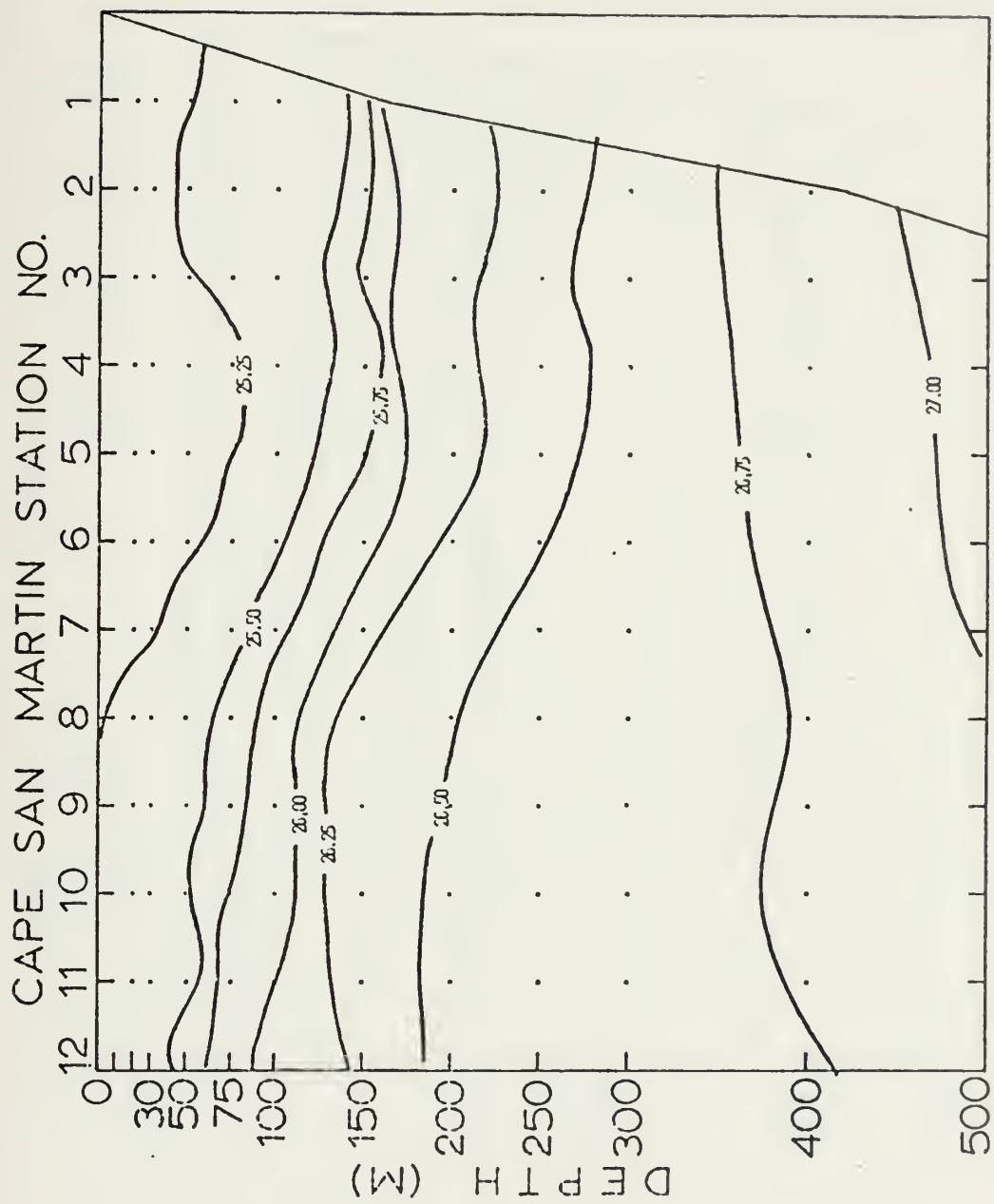


Figure 53. Sigma-t on a vertical section for the Cape San Martin line on 8-9 January 1979.



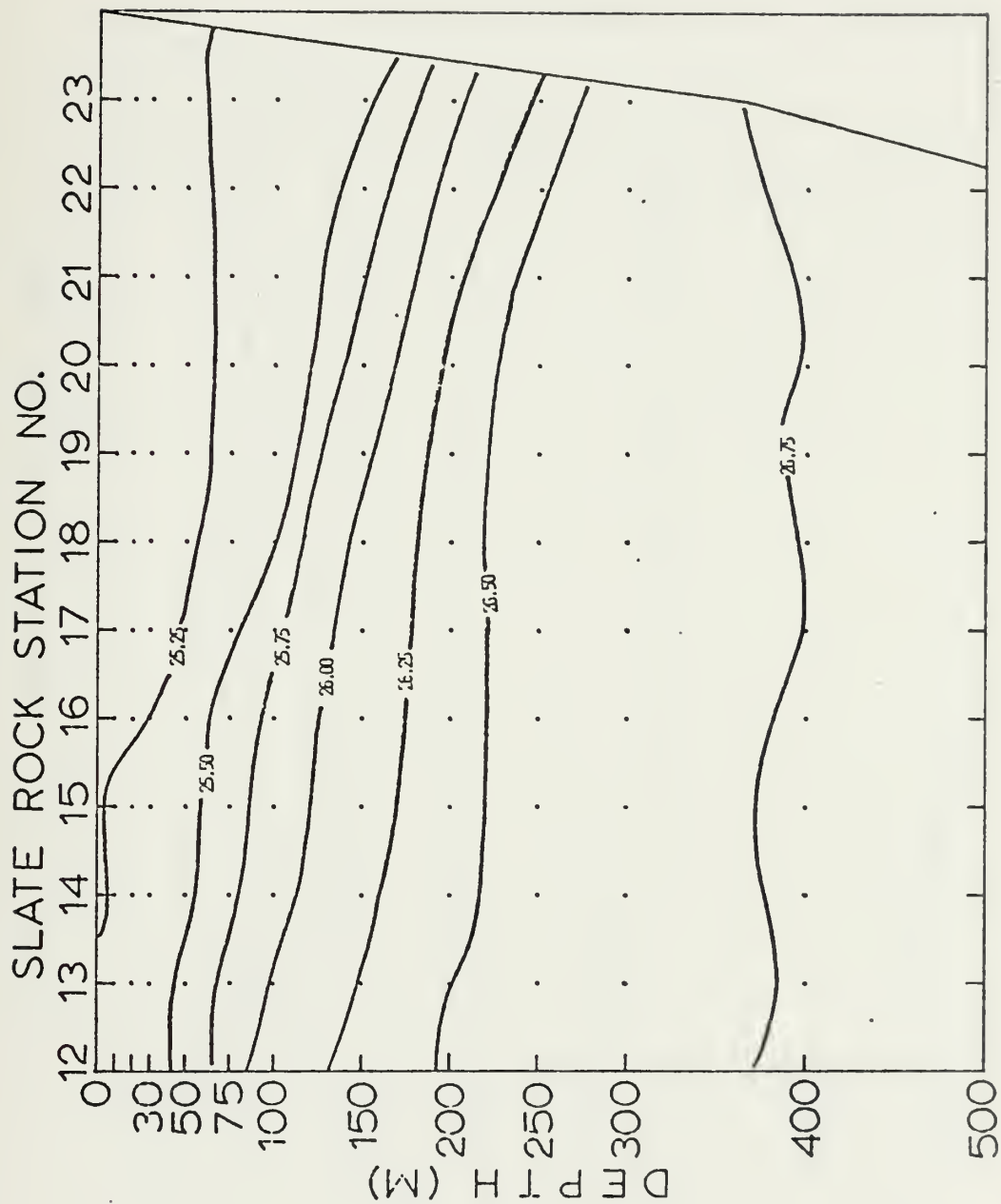


Figure 54. Sigma-t on a vertical section for the Slate Rock line on 8-9 January 1979.



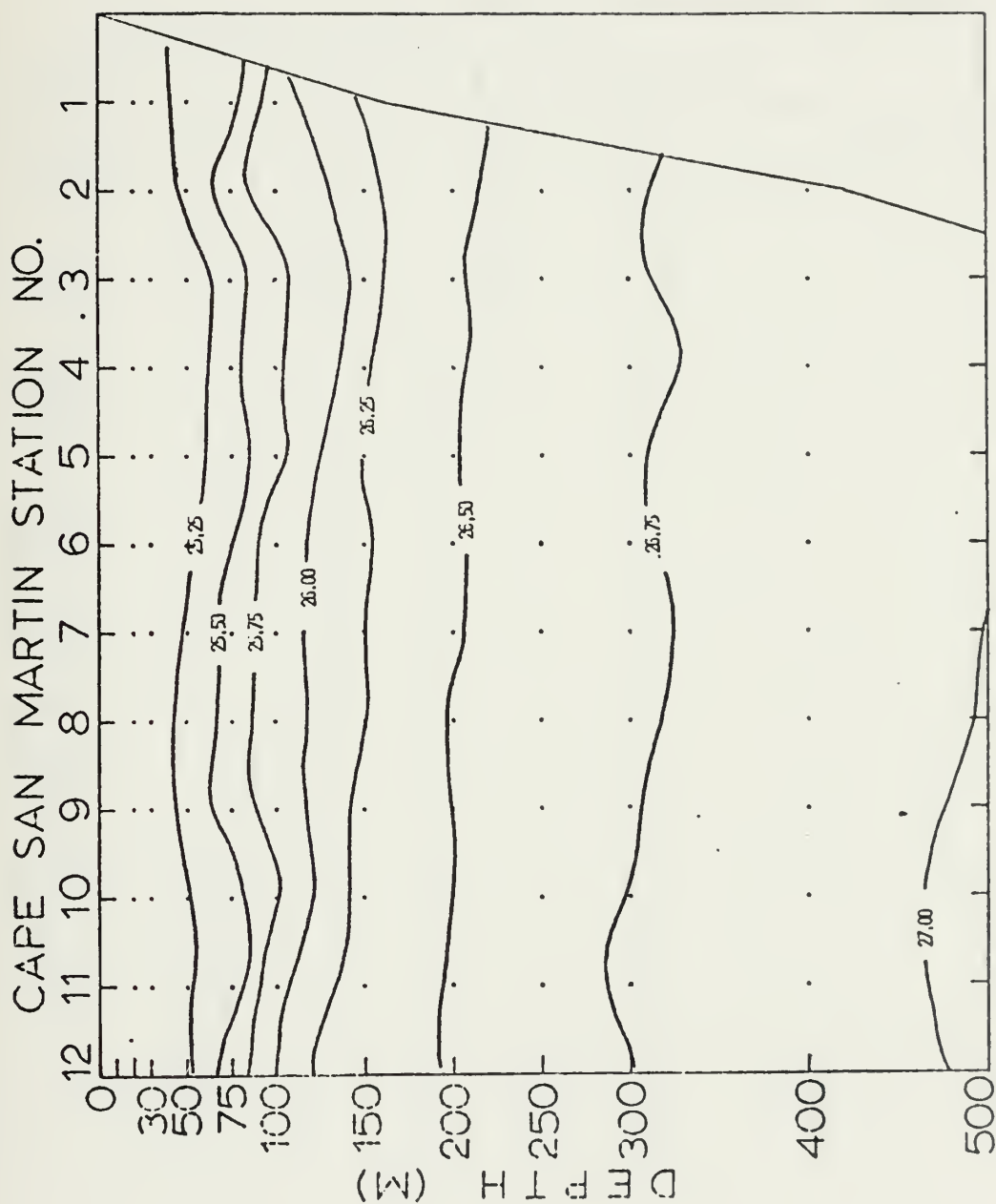


Figure 55. Sigma-t on a vertical section for the Cape San Martin line on 22-23 January 1979.



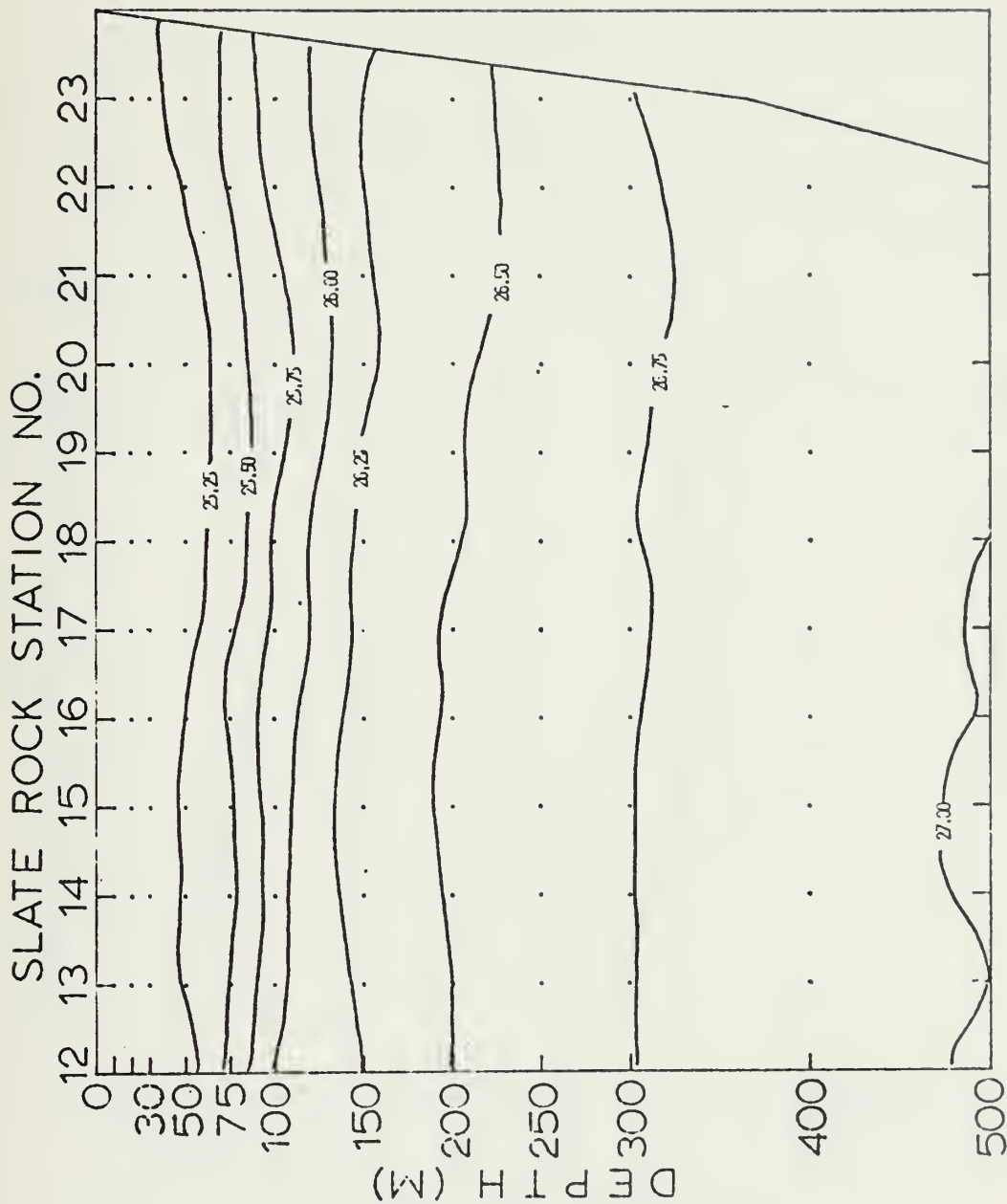


Figure 56. Sigma-t on a vertical section for the Slate Rock line on 22-23 January 1979.





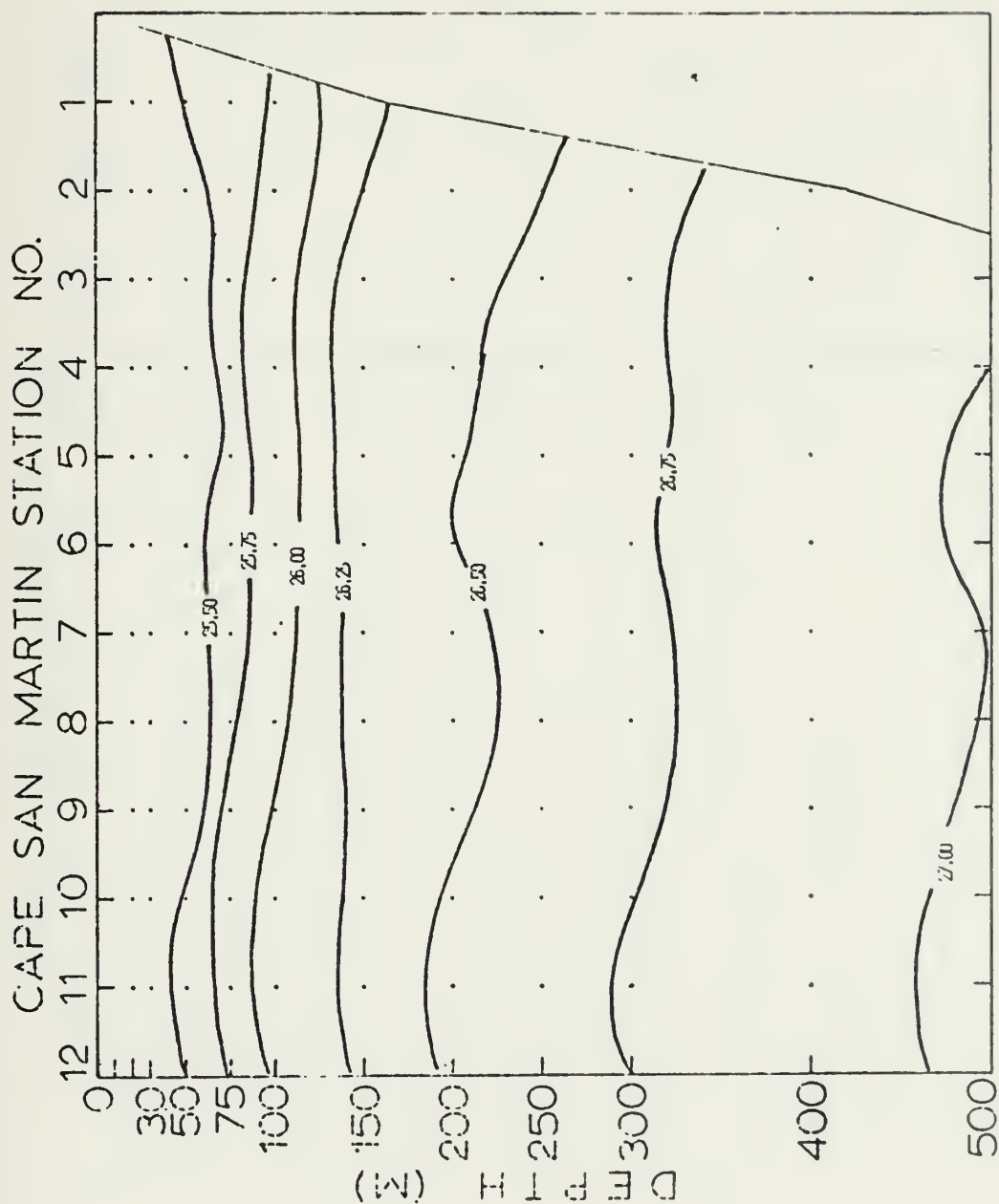


Figure 57. Sigma-t on a vertical section for the Cape San Martin line on 21-22 February 1979.



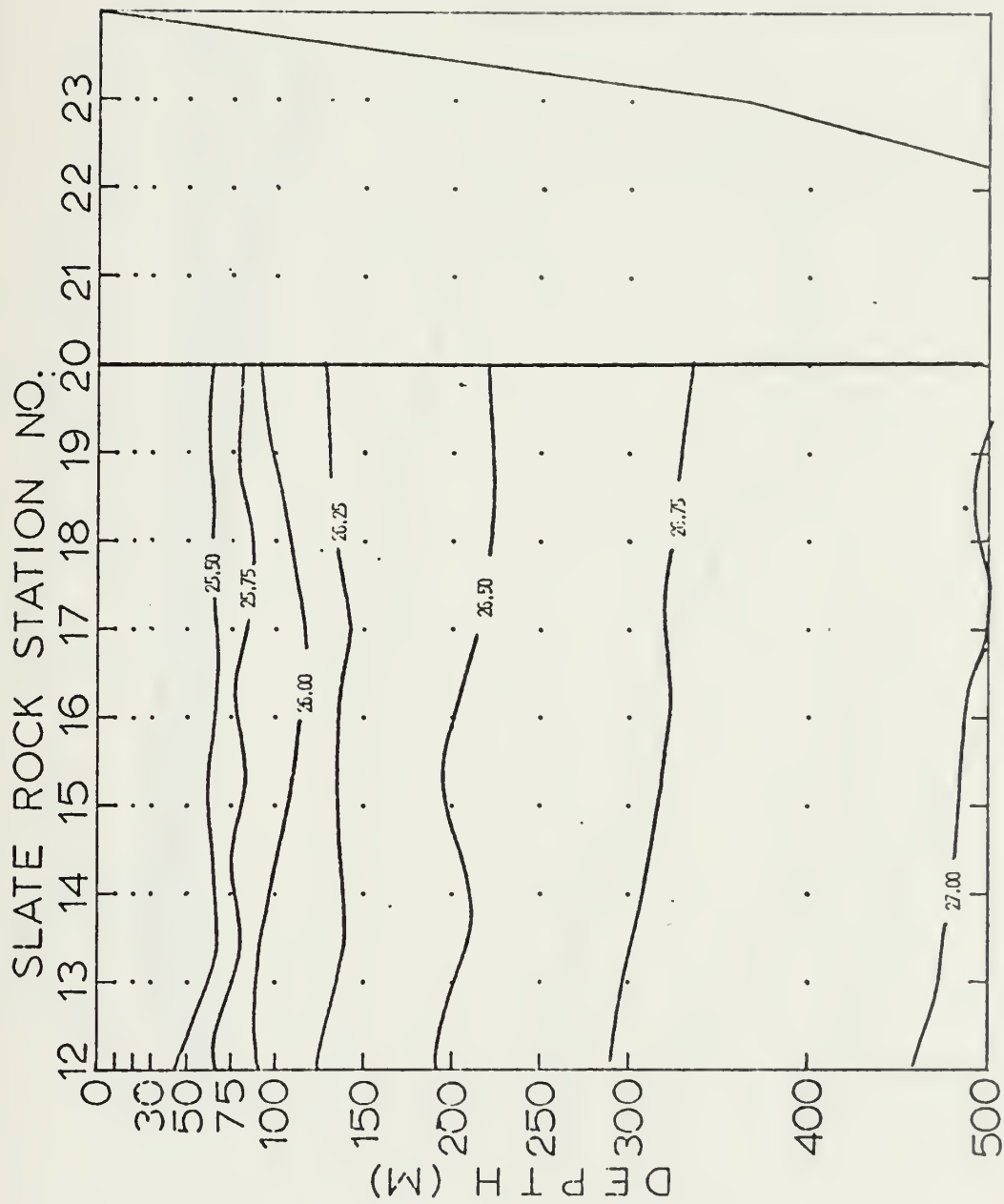


Figure 58. Sigma-t on a vertical section for the Slate Rock line on 21-22 February 1979.







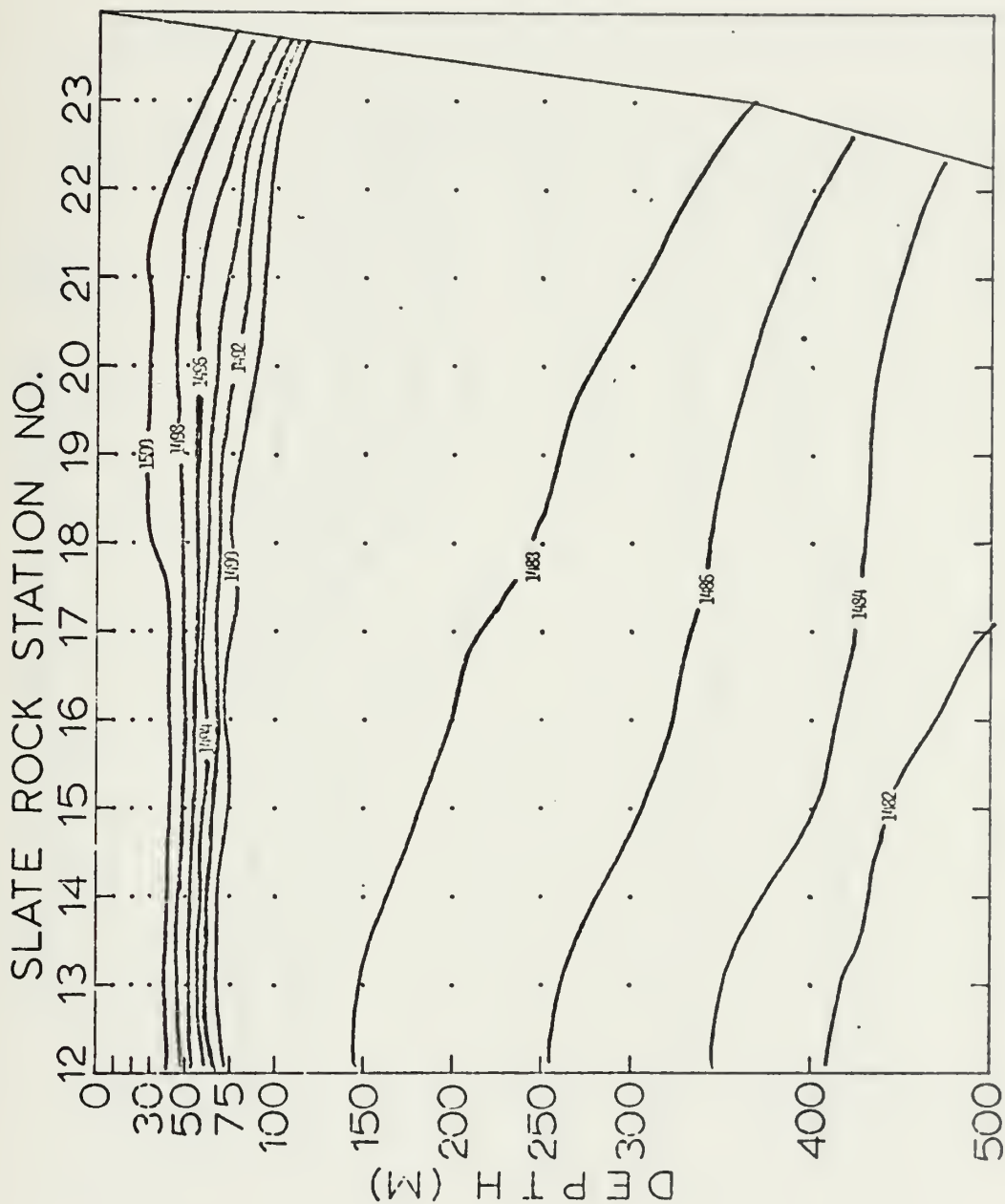


Figure 60. Sound speed (m/sec) on a vertical section for the Slate Rock line on 27-28 November 1978.









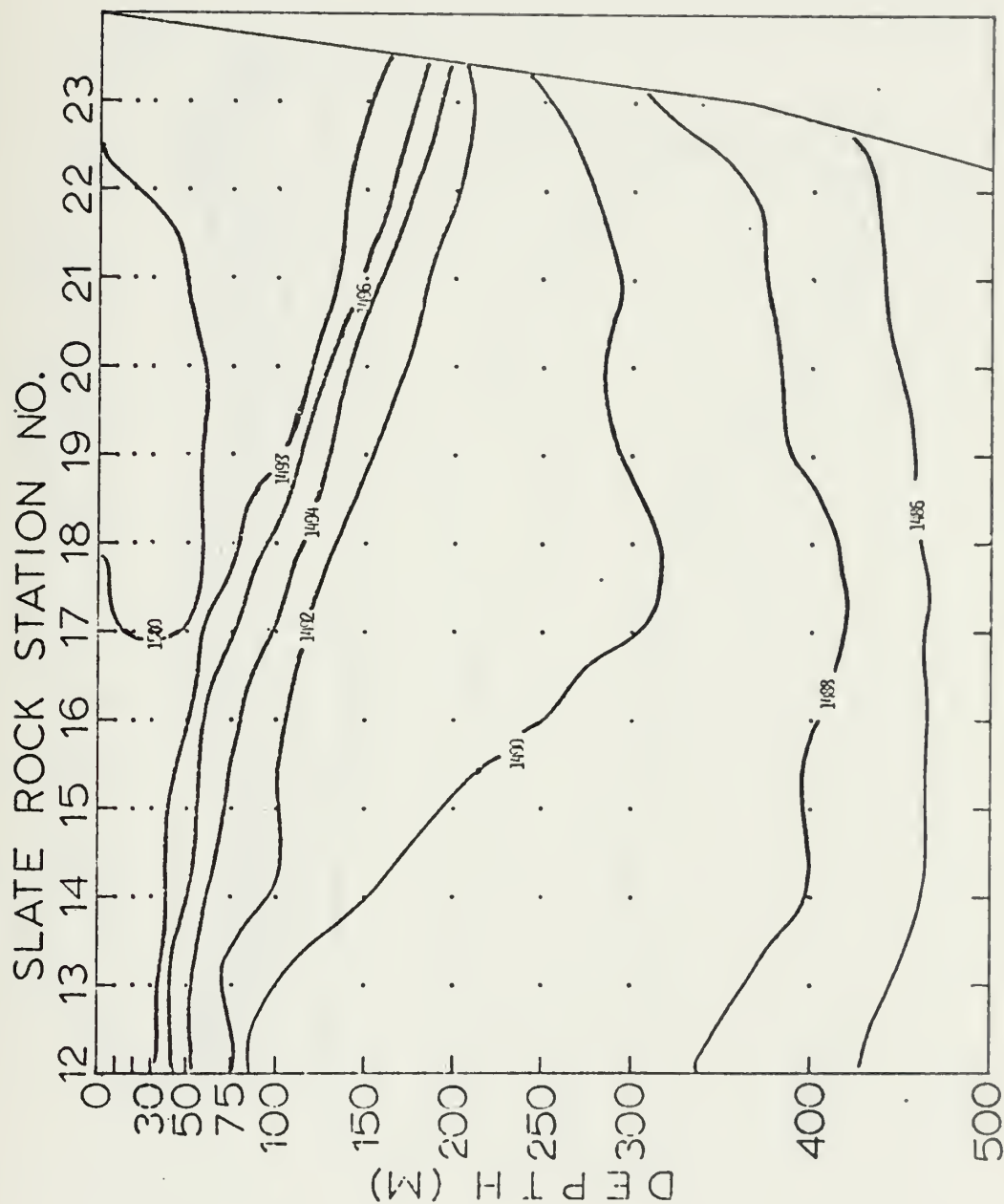


Figure 62. Sound speed (m/sec) on a vertical section for the Slate Rock line on 8-9 January 1979.







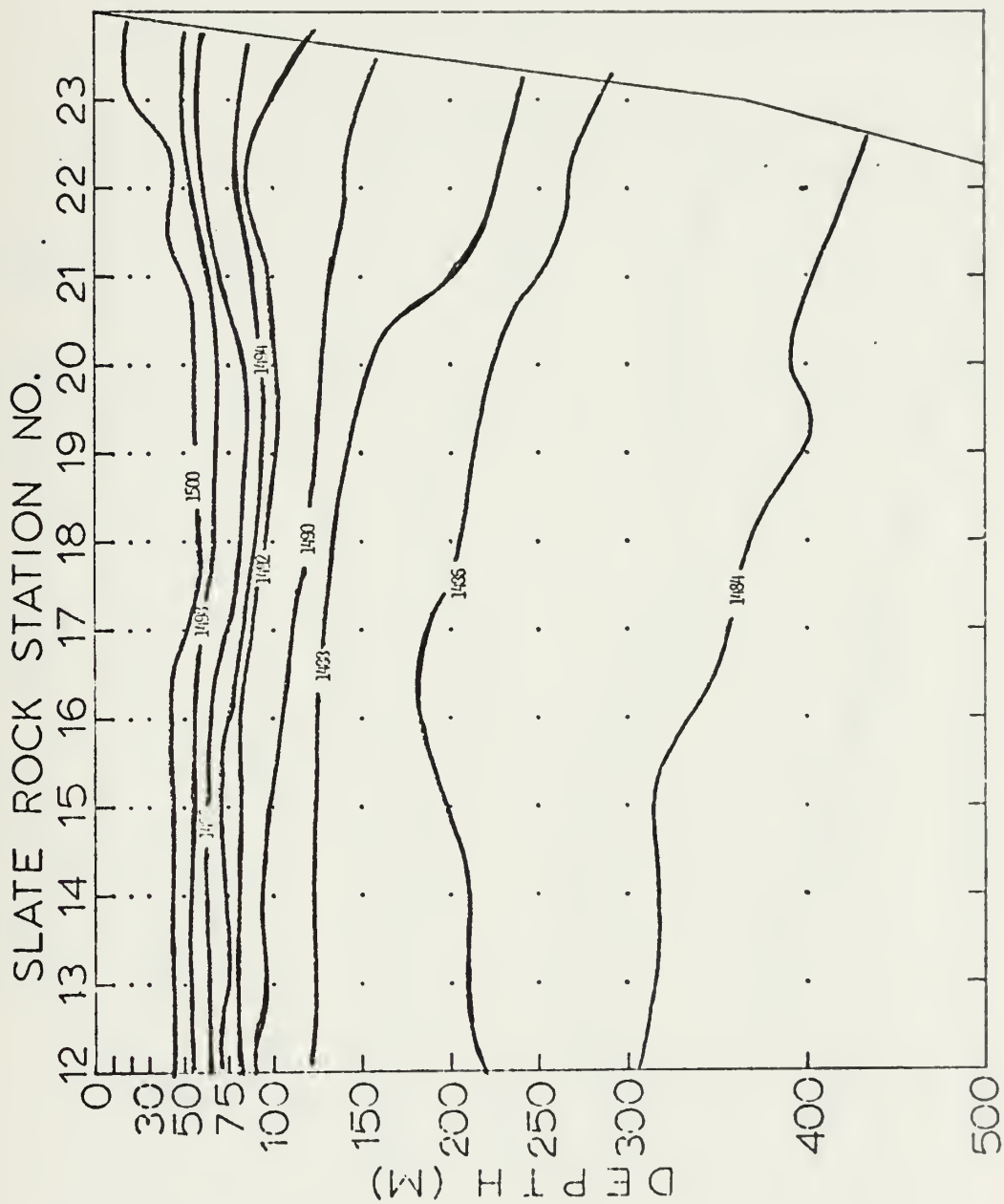


Figure 64. Sound speed (m/sec) on a vertical section for the Slate Rock line on 22-23 January 1979.









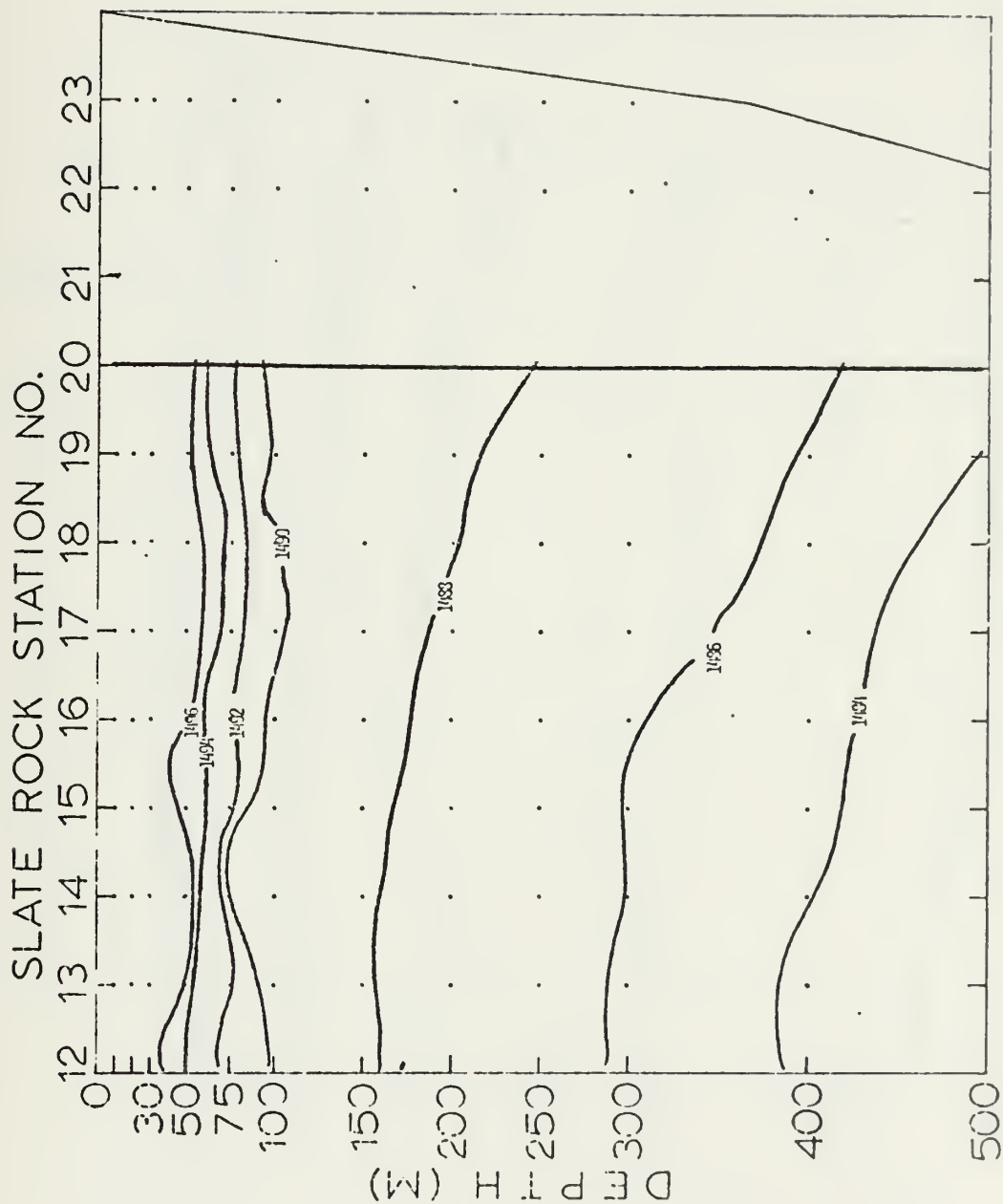


Figure 66. Sound speed (m/sec) on a vertical section for the Slate Rock line on 21-22 February 1979.



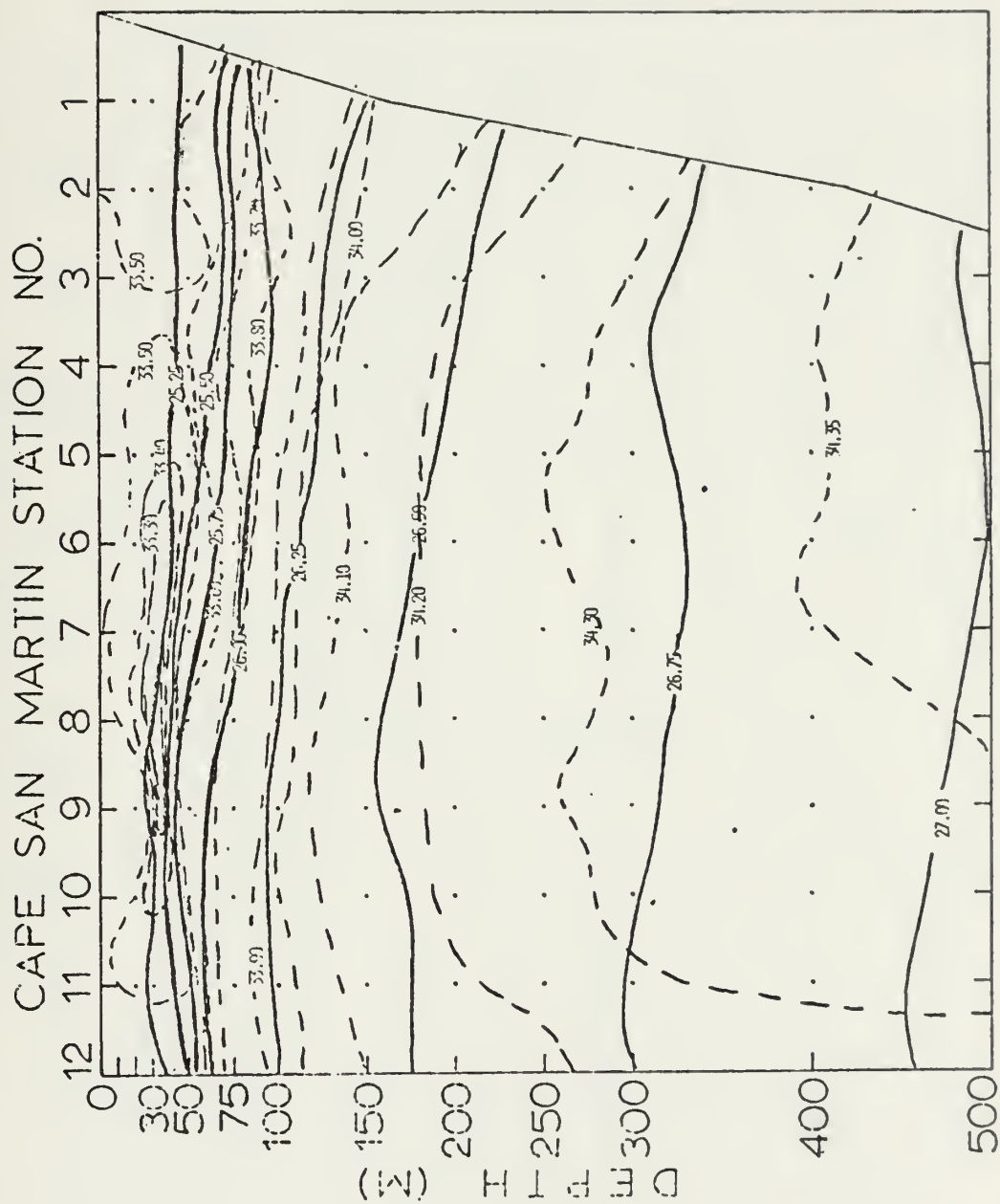
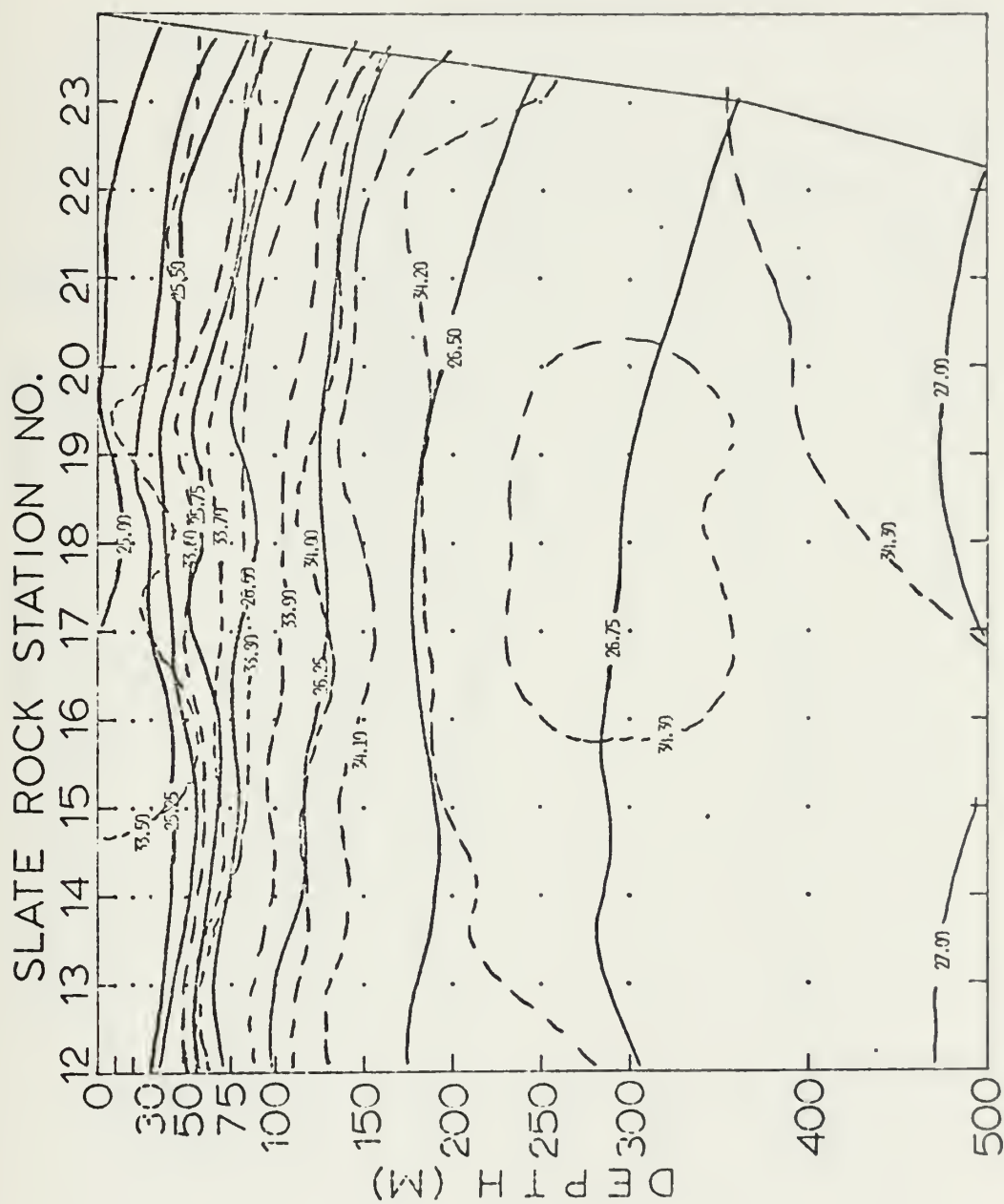


Figure 67. Sigma-t, solid line, and Salinity (‰), dashed line, superimposed on a vertical section for the Cape San Martin line on 27-28 November 1978.









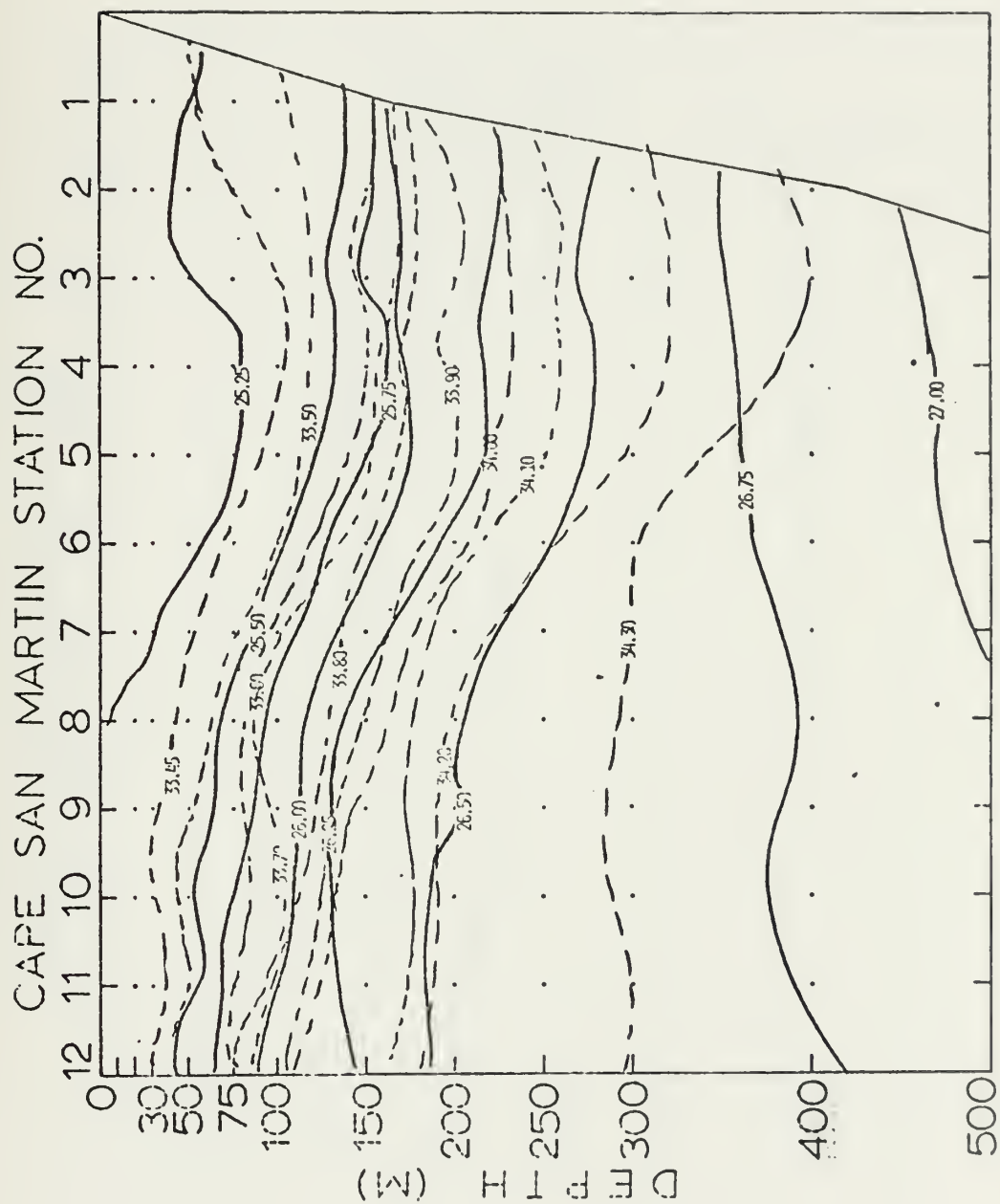


Figure 69. Sigma-t, solid line, and Salinity (‰), dashed line, superimposed on a vertical section for the Cape San Martin line on 8-9 January 1979.



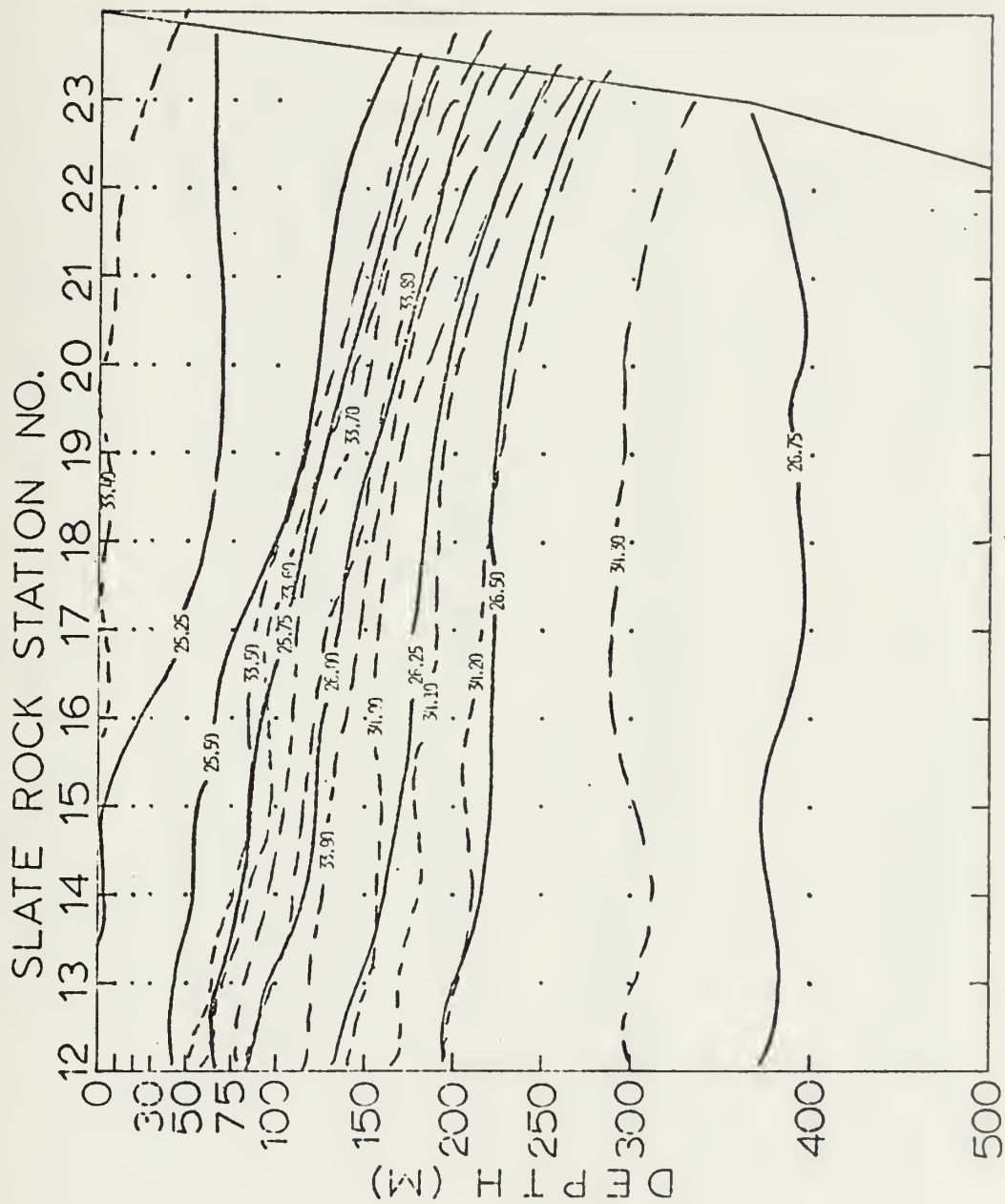


Figure 70. Sigma-t, solid line, and Salinity ( $\text{‰}$ ), dashed line, superimposed on a vertical section for the Slate Rock line on 8-9 January 1979.



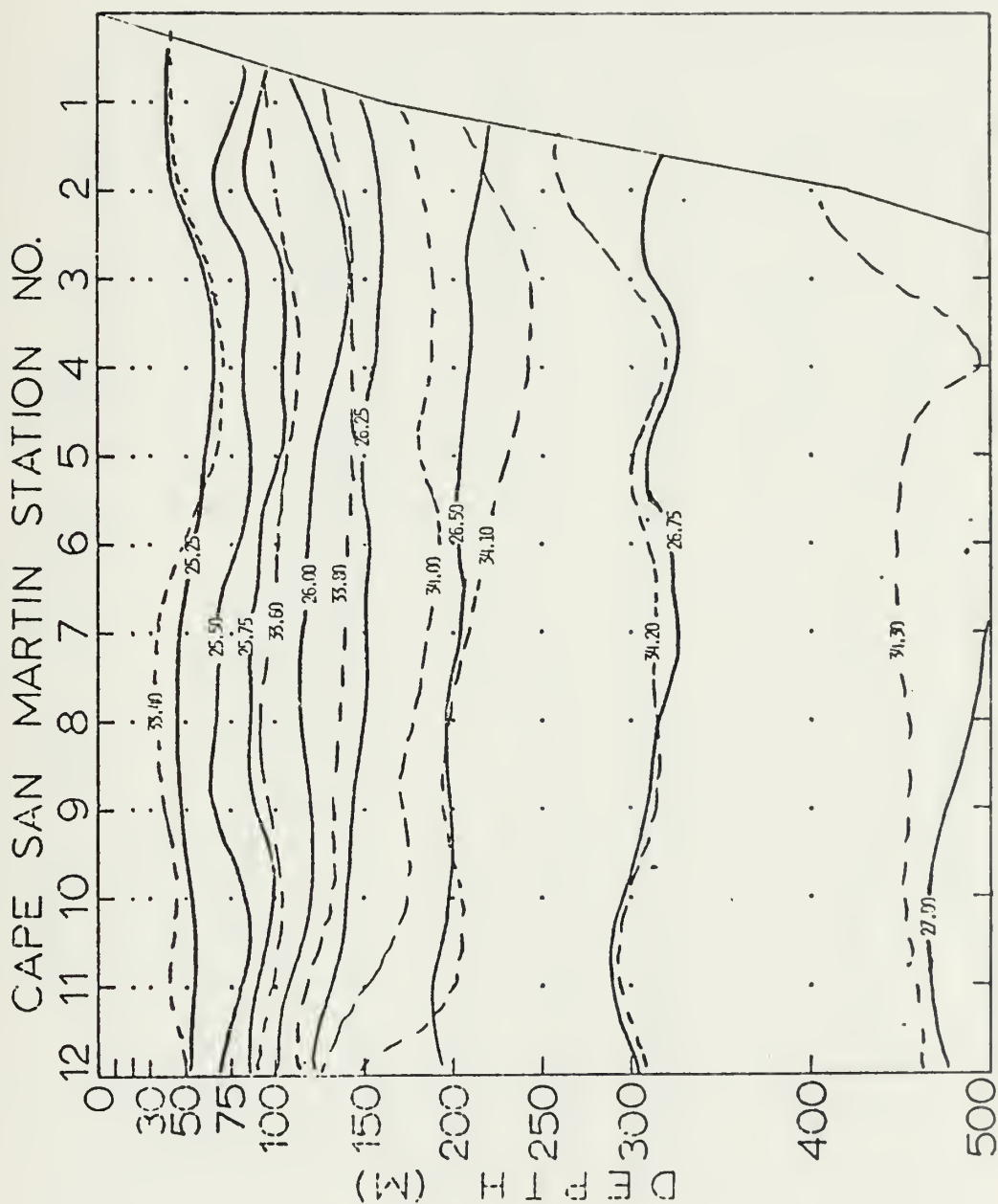


Figure 71. Sigma-t, solid line, and Salinity (‰), dashed line, superimposed on a vertical section for the Cape San Martin line on 22-23 January 1979.



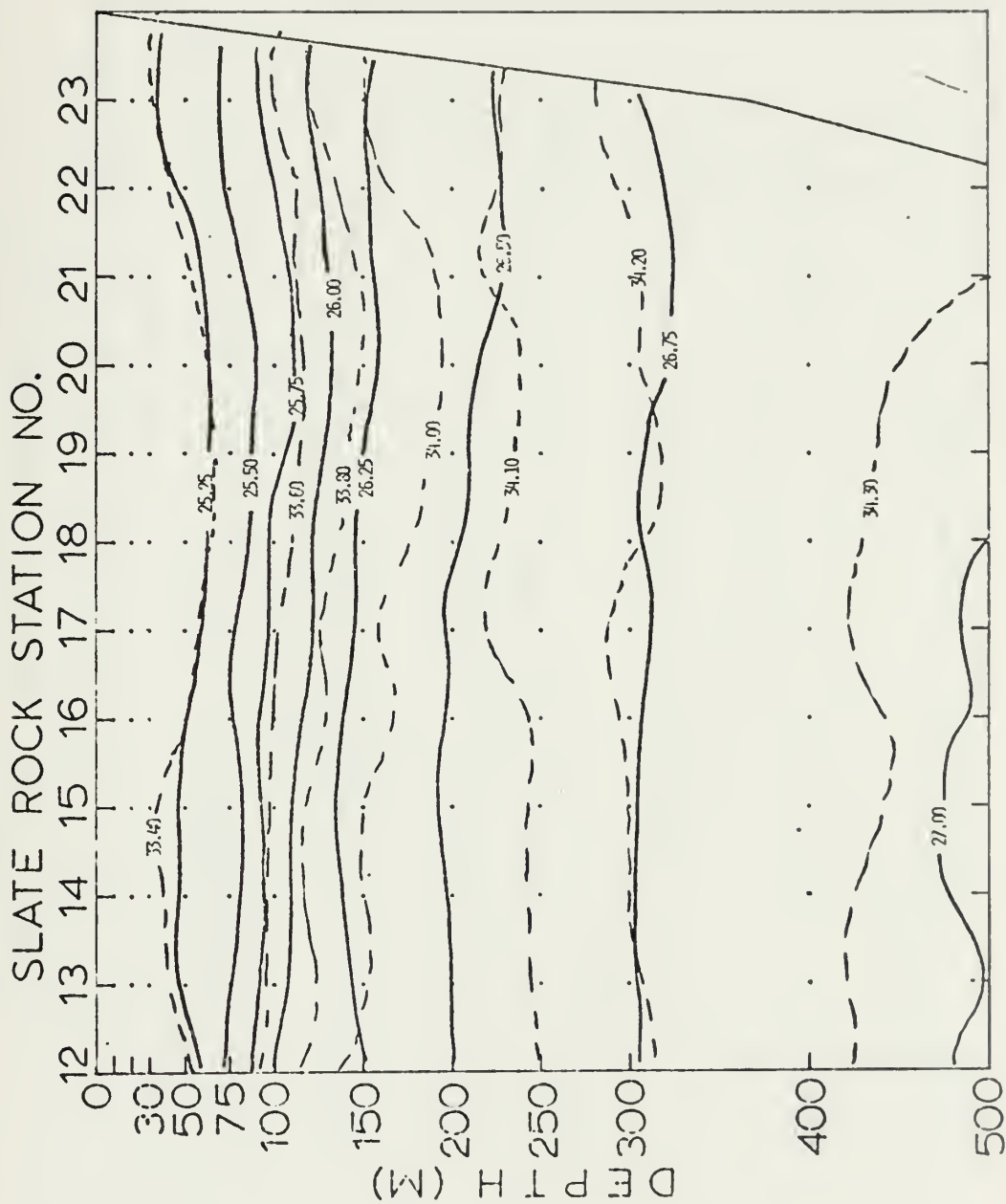


Figure 72. Sigma-t, solid line, and Salinity (‰), dashed line, superimposed on a vertical section for the Slate Rock line on 22-23 January 1979.





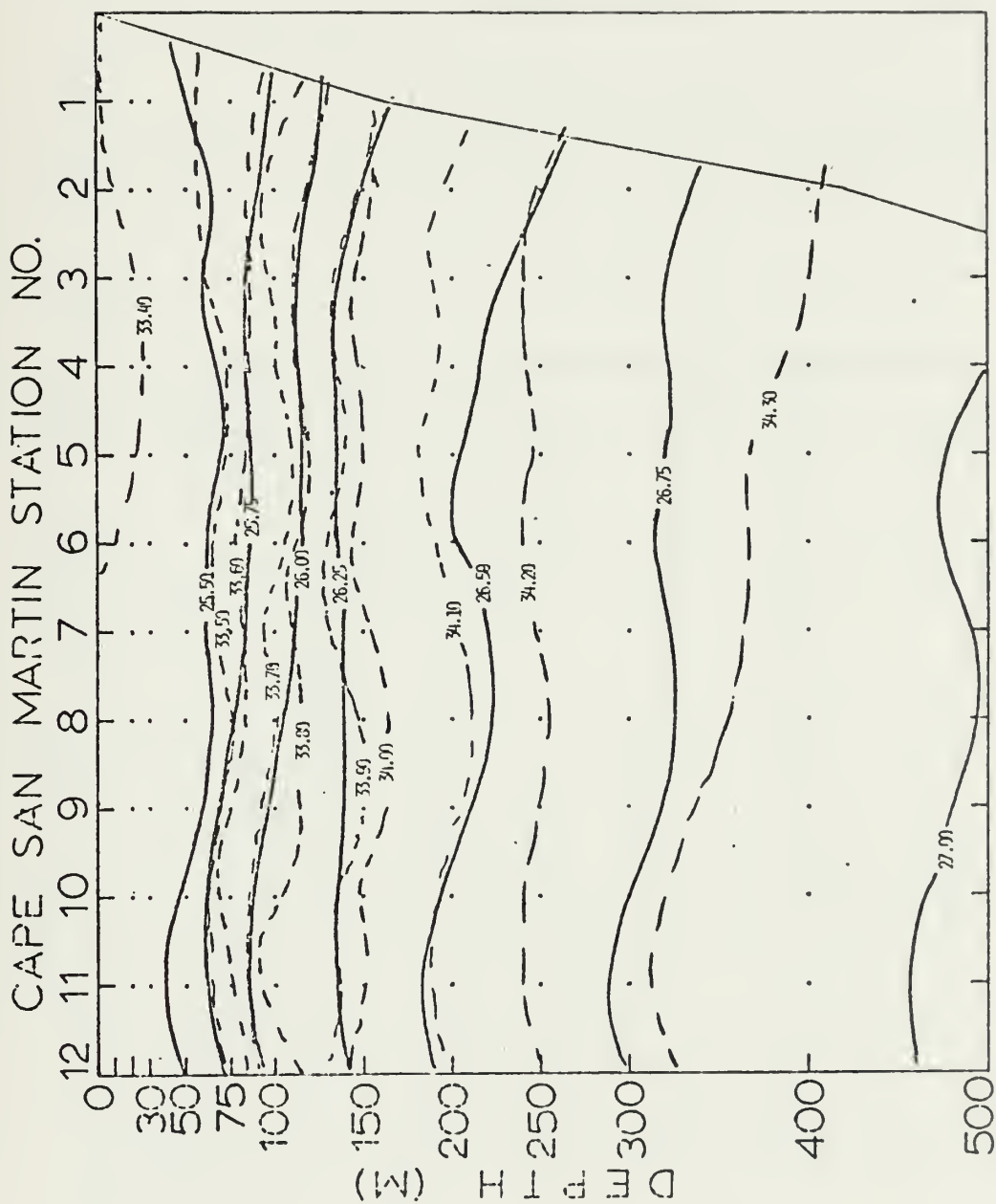


Figure 73. Sigma-t, solid line, and Salinity ( $\text{‰}$ ), dashed line, superimposed on a vertical section for the Cape San Martin line on 21-22 February 1979.



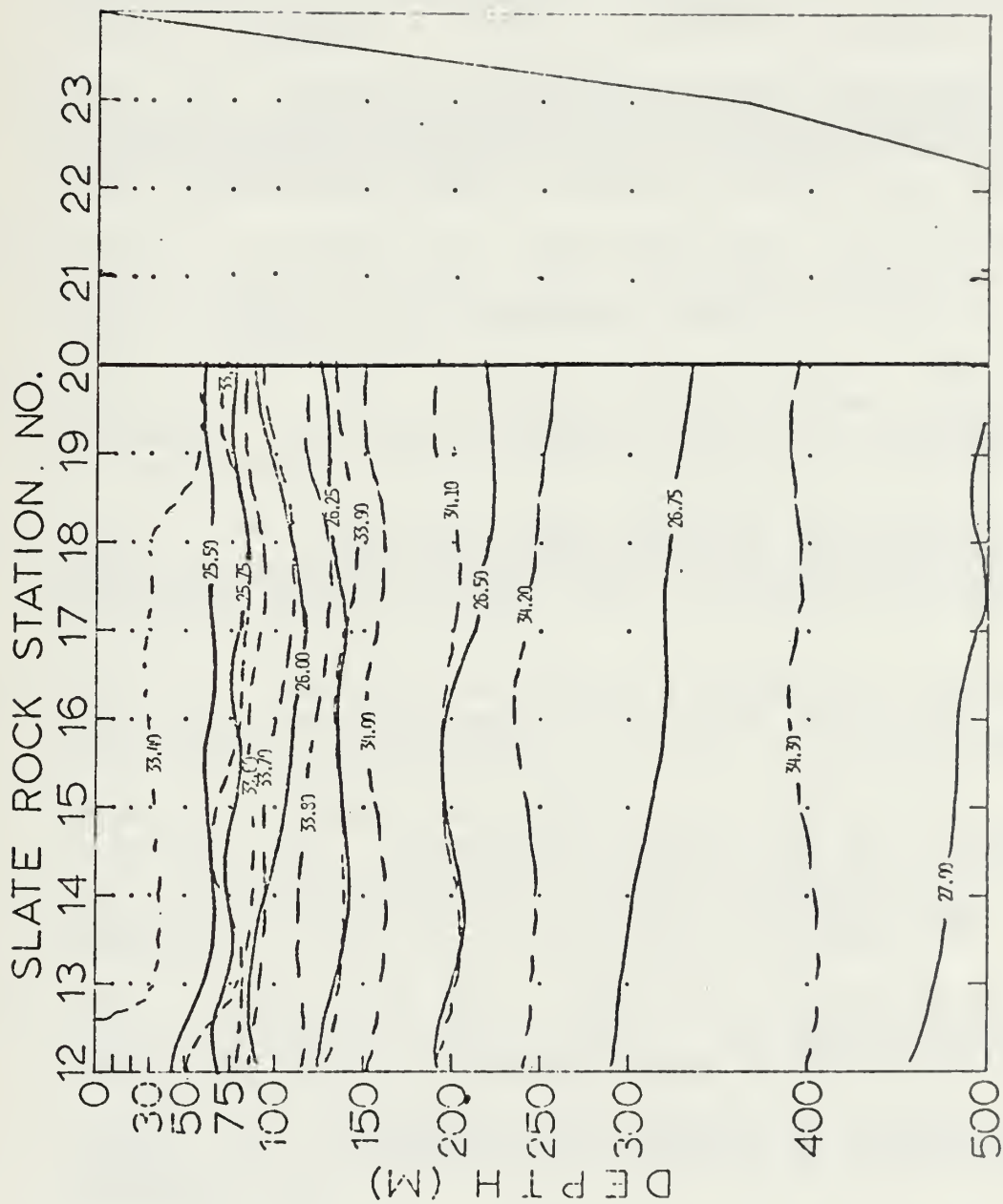


Figure 74. Sigma-t, solid line, and Salinity (‰), dashed line, superimposed on a vertical section for the Slate Rock line on 21-22 February 1979.



## BIBLIOGRAPHY

1. Bernstein, R.L., Breaker L., and Whritner, R., "California Current Eddy Formation: Ship, Air, and Satellite Results," Science, v. 195, p. 353-359, 28 January 1977.
2. Greer, R.E., Mesoscale Components of the Geostrophic Flow and its Temporal and Spatial Variability in the California Current off Monterey Bay in 1973-74, Masters Thesis, Naval Postgraduate School, Monterey, California, 1975.
3. Halpern, D., Smith, R.L., and Reed, R.K., "On the California Undercurrent Over the Continental Slope Off Oregon," Journal of Geophysical Research, v. 83, p. 1366-1372, 20 March 1978.
4. Hickey, B.M., "The California Current System - Hypothesis and Facts," Contribution Number 1038 of the Department of Oceanography, University of Washington, 24 April 1978.
5. Hughes, J.G., The Spatial and Temporal Variation of Sound Speed in the California Current System Off Monterey, California, Masters Thesis, Naval Postgraduate School, Monterey, California, 1975.
6. Huyer, A., "Seasonal Variation in Temperature, Salinity, and Density Over the Continental Shelf Off Oregon," Limnology and Oceanography, v. 22(3), p. 442-445, May 1977.
7. Huyer, A., Hickey, B.M., Smith, J.D., Smith, R.L., and Pillsbury, R.D., "Alongshore Coherence at Low Frequencies in Currents Observed Over the Continental Shelf Off Oregon and Washington," Journal of Geophysical Research, v. 80(24), p. 3495-3505, 20 August 1975.
8. Huyer, A. and Smith, R.L., "Physical Characteristics of Pacific Northwest Coastal Waters," The Marine Plant Biomass of the Pacific Northwest Coast, p. 37-47, 1978.
9. McCreary, J.P., Eastern Ocean Response to Changing Wind Systems, Ph.D Dissertation, University of California, San Diego, 1977.
10. Molnar, D.L., California Undercurrent Reconnaissance Between Monterey and Santa Barbara, Masters Thesis, Naval Postgraduate School, Monterey, California, 1972.



11. Mooers, C.N.K., Collins, C.A., and Smith, R.L., "The Dynamic Structure of the Frontal Zone in the Coastal Upwelling Region Off Oregon," Journal of Physical Oceanography, v. 6, p. 3-21, January 1976.
12. Mysak, L.A., "On the Stability of the California Undercurrent Off Vancouver Island," Journal of Physical Oceanography, v. 7, p. 904-917, November 1977.
13. NOAA Technical Report NMFS SSRF-718, Surface Currents as Determined by Drift Card Releases Over the Continental Shelf Off Central and Southern California, by J.L. Squire, Jr., p. 11, December 1977.
14. Pavlova, Yu. V., "Seasonal Variations of the California Current," Oceanology, v. 13, p. 806-814, August 1967.
15. Reid, J.L., Jr., "Measurements of the California Counter-current at a Depth of 250 Meters," Journal of Marine Research, v. 20, p. 134-137, 15 July 1962.
16. Reid, J.L., Jr., "Measurements of the California Counter-current Off Baja California," Journal of Geophysical Research, v. 68, p. 4819-4822, 15 August 1963.
17. Reid, J.L., Jr., Roden, G.I., and Wyllie, J.G., "Studies of the California Current System," California Cooperative Oceanic Fisheries Investigation Progress Report, p. 27-56, 1958.
18. Reid, J.L., Jr., and Schwatzlose, R.A., "Direct Measurements of the Davidson Current Off Central California," Journal of Geophysical Research, v. 67, p. 559-565, June 1962.
19. Smith, R.L., Pattullo, J.G., and Lane, R.K., "An Investigation of the Early Stage of Upwelling Along the Oregon Coast," Journal of Geophysical Research, v. 71(4), p. 1135-1140, 15 February 1966.
20. Sverdrup, H.U. and Fleming, R.H., "The Waters Off the Coast of Southern California," Scripps Institute of Oceanography Bulletin, v. 4, p. 261-375, 9 October 1941.
21. Sverdrup, H.U., Johnson, M.W., and Fleming, R.H., The Oceans: Their Physics, Chemistry, and General Biology, 23rd ed., Prentice-Hall, Inc., 1942.
22. Wickham, J.B., "Observations of the California Undercurrent," Journal of Marine Research, v. 33, p. 325-340, September 1975.







23. Wilson, W.D., "Speed of Sound in Sea Water as a Function of Temperature, Pressure, and Salinity," Journal of Acoustical Society of America, v. 32, p. 1357, 1960.
24. Woods Hole Oceanographic Institution Technical Report WHOI-73-71, Details of Woods Hole Moorings, by R.H. Heinmiller, and R.G. Walden, October 1973.
25. Woods Hole Oceanographic Institution Report WHOI-76-59, A Computer Program for the Design of a Single-Point Subsurface Mooring Systems: NOYFB, by D.A. Moller, June 1976.
26. Wooster, W.S. and Jones, J.H., "California Undercurrent Off Northern Baja California," Journal of Marine Research, v. 28, p. 235-250, 15 May 1970.
27. Wyllie, J.G., "Geostrophic Flow of the California Current at the Surface and at 200 Meters," California Cooperative Oceanic Fisheries Investigation, Atlas No. 4, December, 1966.

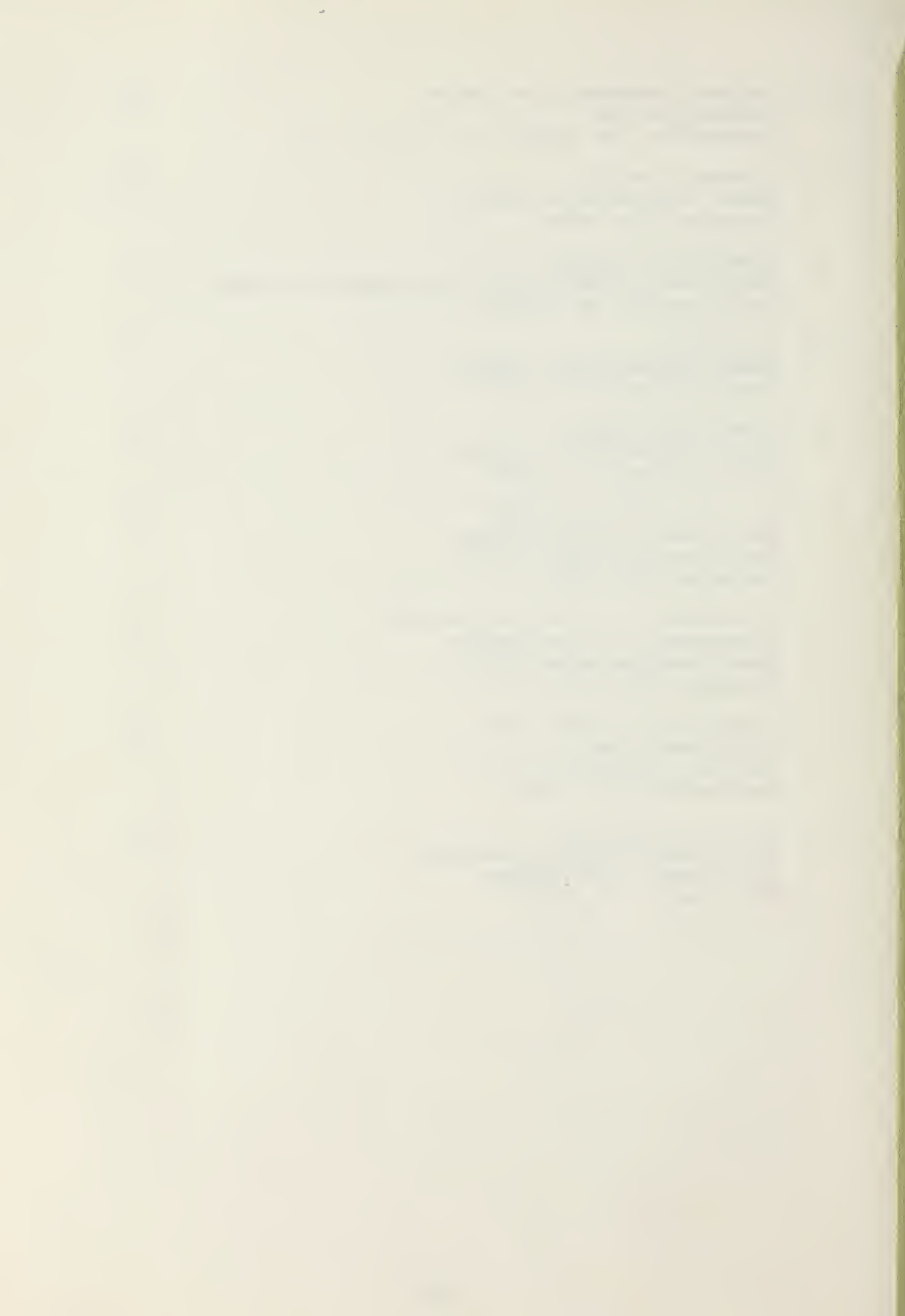


INITIAL DISTRIBUTION LIST

	No. Copies
1. Chairman Code 68 Department of Oceanography Naval Postgraduate School Monterey, CA 93940	3
2. Director Naval Oceanography Division (OP952) Navy Department Washington, DC 20350	1
3. Office of Naval Research Code 480 Naval Ocean Research and Development Activity NSTL Station, MS 39529	1
4. Dr. Robert E. Stevenson Scientific Liaison Office, ONR Scripps Institution of Oceanography La Jolla, CA 92037	1
5. SIO Library University of California, San Diego P.O. Box 2367 La Jolla, CA 92037	1
6. Department of Oceanography Library University of Washington Seattle, WA 98105	1
7. Department of Oceanography Library Oregon State University Corvallis, OR 97331	1
8. Commanding Officer Fleet Numerical Weather Central Monterey, CA 93940	1
9. Commanding Officer Naval Environmental Prediction Research Facility Monterey, CA 93940	1
10. Commander Oceanographic Systems Pacific Box 1390 Pearl Harbor, Hawaii 96860	1



- |     |   |   |
|-----|---|---|
| 11. | Defense Documentation Center<br>Cameron Station<br>Alexandria, VA 22314   | 2 |
| 12. | Library Code 0142<br>Naval Postgraduate School<br>Monterey, CA 93940  | 2 |
| 13. | Commanding Officer<br>Naval Ocean Research and Development Activity<br>NSTL Station, MS 39529                     | 1 |
| 14. | Commander<br>Naval Oceanography Command<br>NSTL Station, MS 39529   | 1 |
| 15. | Commanding Officer<br>Naval Oceanographic Office<br>NSTL Station, MS 39529  | 1 |
| 16. | Dr. S.P. Tucker Code 68Tx<br>Department of Oceanography<br>Naval Postgraduate School<br>Monterey, CA 93940        | 4 |
| 17. | Professor J.B. Wickham Code 68Wk<br>Department of Oceanography<br>Naval Postgraduate School<br>Monterey, CA 93940 | 4 |
| 18. | Commandant (G-PTE-1/72)<br>U.S. Coast Guard<br>400 7th Street S.W.<br>Washington, DC 20590                        | 2 |
| 19. | Lt. K. Coddington<br>Department of Ocean Sciences<br>U.S. Coast Guard Academy<br>New London, CT 06320             | 5 |



000000

28657

Thesis 185670  
C53025 Coddington  
c.1 Measurement of the  
California counter-  
current.

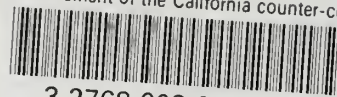
000000

28657

Thesis 185670  
C53025 Coddington  
c.1 Measurment of the  
California counter-  
current.

thesC53025

Measurement of the California counter-cu



3 2768 002 09452 6  
DUDLEY KNOX LIBRARY